



1 Determination of **black carbon** mass concentration from aerosol light

2 absorption using variable mass absorption cross-section

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8 Abstract. Atmospheric black carbon (BC) is the strongest visible solar radiative absorber in the atmosphere, exerting significant 9 influences on the earth's radiation budget. The mass absorption cross-section (MAC) is a crucial parameter for converting light 10 absorption coefficient (σ_{ab}) to mass equivalent BC concentration (m_{BC}). Traditional filter-based instrument, such as AE33, uses a 11 constant MAC of 7.77 m^2/g to derive m_{BC} , which may lead to uncertainty in m_{BC} . In this paper, a new method of converting light 12 absorption coefficient to BC mass concentration is proposed by incorporating the variations of MAC attributed to the influences of 13 aerosol coating state. Mie simulation showed that MAC varied dramatically with different core-shell structures. We compared our 14 new method with traditional method during a field measurement at a site of North China Plain. The results showed that the MAC 15 was smaller (larger) than 7.77 m²/g for particle smaller (larger) than 280 nm, resulting in BC mass size distribution derived from 16 new method was higher (lower) than traditional method for particle smaller (larger) than 280 nm. Size-integrated BC mass 17 concentration derived from new method was 16% higher than traditional method. Sensitivity analysis indicated that the uncertainty 18 in mBC caused by refractive index (RI) was with in 35% and the imaginary part of RI had dominant influence on the derived mBC. 19 This study emphasizes the necessity to take variations of MAC into account when deriving m_{BC} from σ_{ab} and can help constrain the 20 uncertainty in m_{BC} measurements.

21 1 Introduction

22 Black carbon (BC) is an important component of ambient aerosol particles. Because of its highly absorbing properties in the visible 23 spectral region, BC is considered to have a significant influence on global warming. The warming effects of BC is only second to 24 that of carbon dioxide (Ramanathan and Carmichael, 2008). Despite the importance of BC to climate, the global mean direct 25 radiative forcing of BC particles still spans over a poorly constrained range of 0.2 - 1 W/m² (Chung et al., 2012;Bond et al., 26 2013;Boucher et al., 2013). The large uncertainty of BC radiative forcing is partially attributed to the lack of reliable measurements 27 of BC mass concentration in the atmosphere (Arnott et al., 2005; Boucher et al., 2013). Furthermore, BC aerosols can serve as cloud 28 condensation nuclei or ice nucleation particles and change atmospheric convection by heating aerosol layer and influencing the 29 regional precipitation patterns and cloud lifetime (Wild, 2012;Ramanathan and Carmichael, 2008). To fully evaluate the influences 30 of BC particles on solar radiation or precipitation, more precise measurements of BC mass loading in the atmosphere are required. 31 A variety of techniques have been developed to measure real-time BC mass concentrations. Aethalometer (Hansen et al., 1984), 32 Particle Soot Absorption Photometer (PSAP) (Bond et al., 1999), and Multiple-Angle Absorption Photometer (MAAP) (Petzold





33 and Schonlinner, 2004) are based on filter-based attenuation, while the Single Particle Soot Photometer (SP2) is a light-induced 34 incandescent instrument (Stephens et al., 2003;Schwarz et al., 2006). Other instruments that use photo-acoustic methods such as 35 Photoacoustic Spectrometer (PAS) (Truex and Anderson, 1979) or Photo-Acoustic Soot Spectrometer (PASS) have also been 36 introduced. The aethalometer AE33 (model 33, Magee, USA), a convenient and rapid instrument, is commonly used for routine BC 37 observations or dedicated campaigns (Castagna et al., 2019;Sandradewi et al., 2008;Helin et al., 2018). It measures real-time BC 38 concentrations by converting the absorption coefficient (σ_{ab}) into mass equivalent BC concentrations (m_{BC}) through a constant mass 39 absorption cross-section (MAC), which provides the BC absorption per unit mass. 40 However, it has been reported that the MAC of BC is substantially affected by the process through which BC mixes with other 41 aerosol components (Gunter et al., 1993; Doran et al., 2007; Lack and Cappa, 2010; Peng et al., 2016). Field measurements have 42 indicated that fresh BC particles are generally subject to several coating processes while being transported in the atmosphere and 43 tend to be covered in layers of other organic or inorganic components (Shiraiwa et al., 2007;Cappa et al., 2019;Bond et al., 2006). 44 The gathered shell that builds up on the BC core, acting as a lens to focus additional incident light on the enclosed BC core, can enhance BC light absorption (Fuller et al., 1999) and has significant influences on the BC radiative forcing (Jacobson, 2001). This 45 46 light absorption enhancement has been termed as "lensing effect" of the BC particles. 47 For typical core-coating mixed BC containing particles, this lensing effect was found to enhance BC absorption by 50-100% (Bond 48 et al., 2006). Schwarz et al. (2008) found that fresh soot particles internally mixed with sulfates and organics during transportation, 49 and the lensing effect enhanced the light absorption by a factor of 1.3-1.5. Some controlled laboratory studies also confirmed the 50 occurrence of absorption enhancement and their conclusions were consistent with the model calculation (Adler et al., 2010;Brem et 51 al., 2012; Shiraiwa et al., 2010). Meanwhile, other field studies demonstrated a wide range of this lensing effect (Cappa et al., 2019). 52 In contrast, some field observations showed a slight absorption enhancement (Cappa et al., 2012; Nakayama et al., 2014). A wide 53 range of MAC (2-25 m²/g) has been reported in previous studies (Bond and Bergstrom, 2006;Sharma et al., 2002;Schwarz et al., 54 2008).

Some studies suggested using site-specific MAC values for converting σ_{ab} into m_{BC} (Martins et al., 1998;Schmid et al., 2006). However, field measurements indicated that MAC showed both large temporal and spatial variability (Bond and Bergstrom, 2006;Lack et al., 2012;Cappa et al., 2012;Ram and Sarin, 2009). Bond and Bergstrom (2006) suggested using consistent MAC and refractive index (RI) values for the BC measurements. In addition to the mixing state, the degree of MAC also relies on diameter of the BC core (D_{BC}), RI, coating thickness, and the location of the BC core (Bond and Bergstrom, 2006;Fuller et al., 1999;Lack and Cappa, 2010). To better determine the current atmospheric BC mass loading, a more reliable MAC application is imperative to infer BC mass from measured light attenuation.

62 The hypothetical BC mixing state affects the corresponding absorption properties. It is critical to propose a method to infer m_{BC} 63 from light attenuation measurements considering aerosol size and the process by which BC aerosols mix with other aerosol 64 components. A simplified core-shell configuration has been introduced to illustrate the structure of BC-containing particles and 65 calculate the relevant optical properties. Several studies have demonstrated that it is appropriate to use the core-shell configuration





- 66 for aged aerosol (Majdi et al., 2020;Liu et al., 2019;Li et al., 2019).
- 67 With the objective of improving the reliability of m_{BC} inferred from AE33, the Mie model incorporated with core-shell configuration 68 hypothesis was applied in this study to assess the limitation of the constant conversion factor used for MAC. Based on the detailed 69 analysis of the relationship among MAC, D_{BC}, and coating thickness (T_{shell}), a modified approach has been proposed for filter-based 70 instruments to derive m_{BC} from σ_{ab} . This modified method measures size-resolvoled m_{BC} accurately and improves the evaluation of 71 BC radiative forcing. 72 2 Dataset and instrumentation 73 The measured BC particle mass size distribution (BCPMSD) was obtained from the field campaign conducted at the Zhangqiu 74 Meteorology Station (36°42'N, 117°30'E), Shandong Province. This field campaign lasted for about 1 month, from July 23, 2017 75 to August 24, 2017. The Zhangqiu observation site is located in the North China Plain (NCP) and is surrounded by farmland and 76 residential areas, representing regional background conditions of the NCP. The DMA (Differential Mobility Analyzer)-SP2 system

77 measurements to determine the number fraction of BC-containing aerosols and to compare AE33 and the three-wavelength

78 photoacoustic soot spectrometer (PASS-3) were conducted in Taizhou (119°57' E, 32°35' N). The suburban measurement site

- 79 Taizhou lies at the south end of the Jianghuai Plain in the East of China. This industrial area between the two megacities of Nanjing
- and Shanghai has experienced severe pollution during the past thirty years. The measurements were conducted from May 24, 2018

to June 18, 2018. The measurements for comparing AE33 and PASS-3 were also conducted from March 20, 2018 to April 30, 2018
and from October 10, 2018 to October 19, 2018 in Peking University (39°59' N, 116°18' E). This site is located at the northwest of

83 Beijing, a megacity experiencing severe and complex urban pollution. Meanwhile, from March 21, 2017 to April 9, 2017 at the

Peking University site, the results from simultaneous measurements from AE51 (model 51, microAeth, USA) and AE33 were compared.

86 All the measurements in the three sites were conducted in containers where ambient temperature was controlled within 24 ± 2 °C 87 with a particle pre-impactor to remove particles larger than 10 µm from the input air stream. The drying systems in the three sites 88 were configured with a Nafion dryer to keep the relative humidity of sample flow below 40%. This type of dryer performs good in 89 reducing aerosol losses. The transmission efficiency of the Nafion dryer is up to 90% for particles smaller than 10 nm and rises up 90 to 100% for particles larger than 30 nm (The performance details of the Nafion dryer can be accessed at http://www.permapure.com). 91 During the field campaign at the Zhangqiu site, the particle number size distribution (PNSD) as well as BCPMSD were 92 simultaneously determined using the measurement system developed by (Ning et al., 2013). The instrument setup was further 93 improved by Zhao et al. (2019b). The polydisperse aerosol sample flow was first drawn into DMA (Model 3080, TSI, USA) to 94 select relatively monodispersed aerosol sub-populations with diameters ranging from 97 to 602 nm. Sheath and sample flows were 95 set as 3 and 0.5 L/min, respectively. The selected monodispersed aerosol populations were further divided into two paths. One path 96 (0.2 L/min) was drawn into AE51 for m_{BC} measurements. The other path (0.3 L/min) was analyzed using CPC (model 3772, TSI, 97 USA) for number concentration measurements. As the standard sample flow for CPC is 1 L/min, a cleaned airflow of 0.7 L/min 98 was added for compensation. A BCPMSD cycle measured here required 5 min and we averaged the data with a temporal resolution





- 99 of 2 hours. The dry aerosol scattering coefficients at 525 nm were measured simultaneously by an integrated nephelometer (Ecotech 100 Pty Ltd., Aurora 3000) with a flow rate of 3 L/min. The temporal resolution was 1 min. Similar to the measured BCPMSD, aerosol 101 scattering coefficients that were used to represent air pollution conditions were also averaged with a temporal resolution of 2 hours. 102 While observing BCPMSD at the Beijing site, Zhao et al. (2019b) added AE33 (3 L/min) simultaneously to measure the bulk m_{BC}. 103 The bulk mBC from AE33 and from the integrated BCPMSD measured by AE51 were then compared. For AE51, the influence of 104 loading effect was resolved by using $\sigma_{ab,corrected} = (1 + k \cdot ATN)\sigma_{ab,uncorrected}$. $\sigma_{ab,corrected}$ and $\sigma_{ab,uncorrected}$ are the 105 corrected and uncorrected σ_{ab} , respectively. Factor k was set as 0.004 and ATN is the measured light attenuation when particles load 106 on the fiber filter of AE51. A recommended compensation of 2.6 was introduced here to mitigate the multiple scattering problem 107 (Zhang et al., 2018). Results showed that the variation trends and magnitudes of m_{BC} measured by AE33 and AE51 were in good 108 consistence (Zhao et al., 2019b). Therefore, in this study, the BCPMSD measured by AE51 was regarded as the measurement results 109 of AE33, and the size-resolved σ_{ab} were retrieved by the constant MAC value of 7.77 m²/g used in AE33. 110 For the coupling DMA-SP2 system measurement, the dried sample flow was drawn into DMA to select particles with diameters 111 ranging from 200 to 450 nm. Then, the selected monodispersed aerosol samples were analyzed in SP2 (0.12 L/min) to identify the 112 BC containing particles and in CPC (0.28 L/min) to count the total number of particles. The size-resolved number fraction of BCcontaining particles was then derived. As the total flowrate was 0.4 L/min from DMA, the sheath flow of DMA was 4 L/min. 113 114 Detailed configuration of the DMA-SP2 system has been demonstrated in a previous study (Zhao et al., 2019a). According to the
- 115 measurements from Taizhou, only 17% of the ambient particles that contained BC averagely for bulk aerosol populations.
- We adjusted the measured wavelengths of AE33 to the measured wavelengths of PASS-3 (405 nm, 532 nm, and 781 nm). Measurement results from Taizhou and Beijing showed that all the ratios of σ_{ab} measured by AE33 with a measurement flowrate of 118 3 L/min and PASS-3 with a measurement flowrate of 1 L/min at the three wavelengths varied slightly over the East and North China Plain (± 0.04), with the average value at 2.9. Therefore, as the measurement results between AE33 and AE51 were consistent, all the size-resolved σ_{ab} from AE51 adopted in this study were corrected through the mean ratio of 2.9. All the measurement systems at the three sites are shown in Fig. S1 in the supplement.
- 122 3 Method
- For current filter-based intruments, m_{BC} are generally derived from σ_{ab} through a constant MAC value. However, the MAC values are enhanced by different degrees when BC particles are mixed with other weakly-absorbing materials, leading to large uncertainties on BC mass retrieval and further evaluations of BC atmospheric optical effects. In order to gain more accurate atmospheric BC mass loading, it is critical to consider the discrepancies in MAC caused by variations in the coating process, BC sizes, etc. Among with the core-shell configuration hypothesis, developing the relationship between MAC, D_{BC} , and T_{shell} is a new approach to correlate
- 128 m_{BC} with σ_{ab} .
- 129 3.1 Core-shell geometry of aerosol particles
- 130 To evaluate the theoretical discrepancies in MAC values caused by the corresponding impact factors, an appropriate model
- 131 simulation is needed for representing a single BC particle's optical properties. There are three widely employed mixing states that





132 are used to represent the structure of BC-containing aerosols: internal, external, and core-shell model (Ma et al., 2011; China et al., 133 2015). Generally, newly-emitted BC particles are chain-like aggregates composed of small spheres. During the coating process, the 134 chain-like BC aggregates become more compact as they collapse and are coated as a core by organic and inorganic materials (Bond 135 and Bergstrom, 2006). Therefore, core-shell configuration is more plausible (Jacobson, 2000). Ma et al. (2012) also indicated that 136 the core-shell assumption can provide a better performance in optical closure than the internal or external models. Furthermore, 137 Moffet et al. (2016) studied particle mixing state and morphology using scanning transmission X-ray microscopy and highlighted 138 that core-shell structure dominated the mixing state of ambient aerosol particles. As aerosols are assumed to be core-shell mixed, 139 with a spherical BC core in the center of the coating sphere, the spherical core and shell favor the Mie model. Therefore, the Mie 140 model was used in this study to simulate the optical properties of BC particles with core-shell mixing state. The consistency in 141 observed and theoretical values obtained using Mie and core-shell morphology support the suitability of this method (Cappa et al., 142 2012).

143 3.2 Mie modeled MAC of BC particles

Many optical simulations for BC particles with concentric sphere geometry have been reported and the corresponding results show that the absorption of a pure BC particle will be enhanced when a shell composed of non-absorbing material deposits on this pure BC particle. Since we focused on the optical properties rather than chemical compositions of the mixed aerosols, a simplified hypothesis of BC/sulfate mixtures, which is frequent in the atmosphere (Khalizov et al., 2009), was introduced in the algorithm for calculating m_{BC}.

149 Since the filter-based instruments (AE33) use σ_{ab} at the wavelength of 880 nm to determine m_{BC}, the MAC distribution for a wide 150 range of core and coating sizes at the wavelength of 880 nm, calculated using the Mie theory, has been presented. The refractive 151 index (RI), reported to vary with incident light wavelength, is an important parameter to determine aerosol optical properties. 152 However, as BC particles can be emitted from different fuels and conditions, RI cannot be observed directly, with both real and 153 imaginary part of RI varying over a significantly wide range. Liu et al. (2018) summarized RI values for specific wavelengths and 154 showed that the real part is generally in the range of 1.5 to 2.0 while the imaginary part usually varies from 0.5 to 1.1 (Sorensen, 155 2001;Bond and Bergstrom, 2006). Therefore, the real part and imaginary part of RI were set to change from 1.5 to 2.0 and from 0.5 156 to 1.1, respectively, with a step increase of 0.01. Meanwhile, the RI of sulfate was set as 1.55-1.0⁻⁶ i and the density of BC was set 157 as 1.8 g/cm³, similar to Bond et al. (2006). A total of 3111 values were obtained, and the averaged values are illustrated in Fig. 1. 158 The D_{BC} and total aerosol particle diameter ($D_{particle}$, D_{BC} + T_{shell}) ranged from 10 to 700 nm. 159 Figure 1 presents several features of the variation pattern of MAC. MAC values varied significantly with D_{BC} and the thickness of 160 non-absorbing coating. When the D_{BC} was less than 100 nm, the thickness of the coating dominated the variation of MAC values,

- 161 and MAC values increased with T_{shell}. As the value T_{shell} increased, the lensing effect became more significant. MAC value can
- 162 increase from 4 m^2/g to about 17 m^2/g when the total aerosol size reached up to 700 nm. When the D_{BC} was larger than about 100
- 163 nm, both T_{shell} and D_{BC} determined MAC values and D_{BC} played a more important role. MAC increased with T_{shell} and decreased
- 164 with the D_{BC}. Moreover, even for pure BC particles, MAC values varied significantly with the size of BC particles. For smaller





- 165 particles, the MAC values increased slightly with BC size until D_{BC} reached 220 nm. Then, MAC decreased with an increase in D_{BC}.
- 166 Therefore, the constant MAC value of 7.77 m^2 /g used in AE33 is only appropriate for a very limited condition.
- 167 **3.3** New method to retrieve m_{BC} by considering the variation of MAC
- 168 In this subsection, we introduce a new method to determine m_{BC} from the measurement of the σ_{ab} at a given diameter. For a given
- 169 $D_{\text{particle,}}$ if D_{BC} is assumed, the corresponding T_{shell} is fixed. Combining the simultaneously measured PNSD and the percentage of
- 170 particles containing BC, the number of BC-containing particles is then determined at D_{particle}. Corresponding absorption properties
- 171 at the D_{particle} with fixed D_{BC} and T_{shell} can be calculated using the Mie model. Hence, if the number concentration of BC-containing
- 172 particles and σ_{ab} at a given $D_{particle}$ are measured, we can infer the D_{BC} by closing the measured and the calculated σ_{ab} . Then, the m_{BC}
- 173 can be obtained from D_{BC} for every D_{particle}. Finally, the BCPMSD is derived.
- The detailed iterative procedure is illustrated in Fig. 2. As the absorption properties of BC particles in different coating states have been evaluated with the Mie model, as represented in Fig. 1, a simplified algorithm for deriving BCPMSD was proposed by considering Fig. 1 as a look-up table. For every specific $D_{particle}$, if a D_{BC} is assumed, the corresponding MAC of the particle can be derived from the look-up table. Then, the σ_{ab} can be derived from the MAC, the assumed BC density (1.8 g/cm³ in this study), and the number of BC-containing particles (17% of the total number for every $D_{particle}$). We adjusted the guessed D_{BC} until the difference between calculated and measured σ_{ab} was within an acceptable range (0.1%). Consequently, the D_{BC} and thus the m_{BC} at a given
- 180 $D_{particle}$ was determined. The m_{BC} at different aerosol sizes were derived separately. Finally, the size-resolved m_{BC} and the bulk m_{BC}
- 181 were obtained.
- 182 It should be pointed out that the retrieval algorithm of BCPMSD is based on the assumption that BC-containing particles of a fixed

183 diameter are all core-shell mixed and the corresponding D_{BC} for a specific D_{particle} is same. Moreover, a constant number percentage 184 (17%) of BC-containing particles was adopted in this study. However, the BC-containing particle fraction varied with the primary 185 source, time, coagulation, and extent of atmospheric process. The influence attributed to the constant fraction of BC-containing 186 particles has been discussed in section 2 of the supplement. Additionally, Bond et al. (2013) summarized the density for different 187 graphitic materials. The density values are 1.8 - 2.1 g/cm³ for pure graphite, 1.8 - 1.9 g/cm³ for pressed pellets of BC, and 1.718 g/cm3 for fullerene soot. A constant density (1.8 g/cm3) for BC was briefly used to calculate MAC and BC mass from the volume 188 189 of particles with a diameter of D_{BC} . Therefore, the uncertainty of derived m_{BC} in this study simply depends on the ratio of 1.8 g/cm³ 190 and the real density. Finally, the MAC values in the look-up table were the averaged values for different RI and the corresponding 191 effects have been discussed in section 5.

- 192 4 Results and discussion
- Figure 3 provides a comprehensive overview of the variations in measured and retrieved size-resolved parameters during the campaign. As evident from Fig. 3(a), for the BCPMSD derived by the new method, two modes were found, similar to the results of AE33. Figure 4(a) shows the averaged BCPMSD derived from the new method and AE33 during the campaign. The finer mode was located between 97 – 240 nm while the coarser mode was located between 240 – 602 nm. Figure 3(b) represents the relative deviations between the BCPMSD derived from the new proposed method and those derived from a constant MAC value of 7.77





- 198 m²/g. The results indicate that with the boundary of 280 nm, two opposite deviation tendencies exist. For aerosol particles larger 199 than 280 nm, the m_{BC} derived by the new method were mostly lower than those derived with the constant MAC value of 7.77 m²/g. 200 In contrast, when aerosol particles were smaller than 280 nm, the m_{BC} from the new method were significantly higher than those 201 calculated by the constant MAC, as shown in Fig. 3(c). Figure 3(c) shows the time series of size-resolved MAC during the derivation 202 process of BCPMSD. According to Fig. 3(c), for aerosol particles smaller than 280 nm, the corresponding MAC was almost lower than 7.77 m²/g. This is because the MAC values of particles smaller than 280 nm are mostly lower than 7.77 m²/g, as represented 203 204 in Fig. 1. A smaller MAC implies a weaker absorption ability, which means that the same measured σ_{ab} will correspond to an 205 increased m_{BC}. Therefore, more BC mass loadings were derived from the new method. For aerosol particles larger than 280 nm, in 206 order to match the measured σ_{ab} , the corresponding D_{BC} were generally found to be in those regions of look-up table where the 207 MAC values were larger than 7.77 m²/g (Fig. 3(c)). Thus, the BC mass loadings for particles larger than 280 nm were found to be 208 less than those calculated with the constant MAC value of 7.77 m²/g. The simultaneously measured scattering coefficients at 525 209 nm were introduced here to represent air pollution. As shown in Fig. 3(d), the observation station experienced different levels of 210 pollution. Deviations of m_{BC} derived from the newly proposed method and the constant MAC at different aerosol sizes did not show 211 dependencies on pollution conditions.
- Figure 3(e) shows the time series of m_{BC} at finer and coarser modes. Compared to the results of AE33, the m_{BC} were more concentrated in the finer mode as compared to the coarser mode. The m_{BC} at finer mode were found to be higher than those at the coarser mode for 73% of the experiment duration. The variation trends of bulk m_{BC} calculated by considering the variations of MAC and a constant MAC were similar (Fig. 3(f)). The bulk m_{BC} calculated by the new method were higher than those derived by the constant MAC in 83% of the experiment duration.
- 217 The m_{BC} calculated from the new method and AE33 for different aerosol size ranges were statistically analyzed. As shown in Fig. 218 4, for all m_{BC} of aerosols ranging between 97 - 602 nm and 97 - 280 nm derived from new method and AE33, strong linear 219 relationships were observed with correlation coefficients of 0.99 and 1.00, respectively. The ratios between the m_{BC} derived from 220 AE33 and the new method for aerosol diameter ranges of 97 - 602 nm and 97 - 280 nm were 0.84 and 0.69, respectively, indicating 221 that the m_{BC} obtained from AE33 was 16% lower for bulk aerosol particles and 31% lower for aerosols smaller than 280 nm. For 222 the diameter range of 280 - 602 nm, MAC values varied significantly and the deviations in m_{BC} derived from the new method and 223 AE33 were divided into two types with a boundary of $0.7 \ \mu g/m^3$. If the m_{BC} derived from AE33 was lower than $0.7 \ \mu g/m^3$, there 224 was a relatively consistent ratio of 1.13 between the m_{BC} from the new method and AE33, with a correlation coefficient of 0.95. 225 Therefore, BC mass loading from the AE33 algorithm was 13% higher for aerosol particles larger than 280 nm and m_{BC} lower than 226 $0.7 \,\mu\text{g/m}^3$. However, when the m_{BC} derived from AE33 was larger than 0.7 $\mu\text{g/m}^3$, data points become discrete, and the relationship between the m_{BC} derived from AE33 and the new method could be expressed through an equation (y = 0.29 + 0.48x). However, 227 228 these comparisons for aerosols at different size ranges were obtained based on the measurements in the NCP. Additionally, the 229 number of samples where m_{BC} of 280 – 602 nm were larger than 0.7 μ g/m³ was too small. Further studies on BCPMSD in conjunction 230 with the PNSD measurements at different sites need to be carried out.





231 5 Influences of RI on MAC

232 As the RI of BC is still reported to vary over a wide range and the MAC used in this study was a mean value, it is critical to assess 233 the impact caused by the real and imaginary parts on the calculated MAC and the derived BC mass concentrations. For aerosol 234 particles with fixed D_{BC} and T_{shell}, we calculated the MAC of BC with the real part of RI ranging from 1.5 to 2.0 and imaginary part 235 ranging from 0.5 to 1.1. The step increase of both real and imaginary parts was 0.01 and there were 3111 MAC values for every 236 aerosol particle with fixed BC core size and T_{shell}. The ratio of standard deviation to the mean value for these 3111 MAC values 237 have been presented to demonstrate the uncertainty in MAC due to the uncertainty of BC RI. 238 Figure 5(a) shows the uncertainties in MAC along different values of D_{BC} and T_{shell}. It shows that aerosol particles with a small BC 239 core have larger uncertainties and all the uncertainties were below 24%. When D_{particle} was fixed, the uncertainties decreased with 240 D_{BC}. When D_{BC} was determined, the uncertainties did not change much with T_{shell}. For pure BC particles, the uncertainties also 241 decreased with BC particle size. Figure 5(b) shows the uncertainties when the imaginary part was fixed at 0.8 and the real part 242 ranged from 1.5 to 2.0 with an interval of 0.01. It can be seen that when the imaginary part of RI was fixed, variations in the real 243 part led to slight uncertainties. All the uncertainties were found to be below 14%. Figure 5(c) demonstrates the uncertainties when 244 the real part was fixed at 1.75 and the imaginary part ranged from 0.5 to 1.1 with an interval of 0.01. Comparing Fig. 5(a) and 5(c),

- 245 we can see that the patterns of MAC uncertainties were similar. Overall, the uncertainties were dominated by the variations of the 246 imaginary part and only slightly affected by variations in the real part.
- 247 The variations in on m_{BC} caused by the uncertainties in RI were further evaluated. As stated in section 3.2, all MACs in the look-up 248 table in Fig. 1 are the mean values as the imaginary part and real part of BC RI varied over a wide range. Therefore, the mean MACs 249 in the look-up table plus corresponding standard deviation (MAC + Std) and minus corresponding standard deviation (MAC - Std) 250 were utilized to show the uncertainties in m_{BC} caused by RI of BC. As we can see from Fig. 6(a), irrespective of the MAC values in 251 look-up table were MAC + Std or MAC – Std, there was no change in the mode of BCPMSD. The derived m_{BC} of all aerosols 252 ranging from 97 - 602 nm increased when the MAC values used in the look-up table were MAC - Std and decreased when MAC + 253 Std values were used in the look-up table. Compared to the bulk mBC retrieved through the look-up table with mean MAC, those 254 derived through the look-up table with MAC - Std were higher within 35% (Fig. 6(b)). The decrease in the magnitude of derived 255 m_{BC} caused by MAC + Std values was significantly less than the increase in the magnitude caused by the MAC – Std values. 256 Similarly, for aerosol particles at both finer and coarser modes, the deviations in m_{BC} caused by MAC + Std or MAC - Std were 257 also within 35% (Fig. 6 (c) and Fig. 6 (d)). Meanwhile, the increase in the magnitude of derived σ_{ab} into m_{BC} caused by the MAC-258 Std values was also significantly higher than the decrease in the magnitude caused by the MAC + Std values. This sensitivity study 259 indicated that the accuracy of the retrieved BCPMSD is sensitive to the accuracy of MAC values in the look-up table, especially 260 when the real MACs are less than the mean MAC values used in the look-up table.
- 261 An idealized concentric core-shell model with a spherical BC core fully coated by sulfate was configured to study the MAC of BC
- $\label{eq:constraint} 262 \qquad \text{aerosols and derive the } m_{BC}. \ \text{However, freshly emitted BC particles were found to normally exist in the form of loose cluster-like}$
- aggregates with numerous spherical primary monomers (Liu et al., 2015). Soon after, these aggregates become coated with other





- components and collapsed to a more compact form during the coating process (Zhang et al., 2008;Peng et al., 2016). Therefore, the
- 265 uncertainty in the idealized core-shell configuration is discussed in section 3 of the supplement.
- 266 6 Conclusions
- 267 There was a significant variability in the MAC values of BC with the size of BC core and the thickness of coating, which exerted a
- 268 significant influence on the optical method for measuring m_{BC}. In this study, a new method was proposed to derive m_{BC} while
- 269 considering the lensing effect of core-shell structure and subsequently the MAC variations of BC.
- 270 A look-up table describing the variations of MAC attributed to the coating state and size of BC core was established theoretically
- 271 using Mie simulation and assuming a core-shell configuration for BC-containing aerosols. Then, the m_{BC} at different aerosol sizes
- 272 were derived by finding an appropriate BC core configured with a MAC value from the look-up table to close the calculated and
- 273 measured σ_{ab} .
- 274 This newly proposed method was applied to a campaign measurement in the NCP. There were two modes for BCPMSD at the
- 275 accumulation mode separated by 240 nm. For 73% of the cases, the m_{BC} of the finer mode were larger than those of the coarser
- 276 mode during the measurement. The m_{BC} derived by the new method were mostly lower than those derived by a constant MAC of
- 277 7.77 m²/g for particles larger than 280 nm, and higher for particles smaller than 280 nm. Similarly, the bulk m_{BC} accumulated from
- 278 BCPMSD derived from the new method were mostly higher than those derived from a constant MAC of $7.77 \text{ m}^2/\text{g}$.
- 279 The uncertainty in derived m_{BC} that was caused due to the wide range of RI of the BC core was also studied. The results indicated
- that the uncertainty of the imaginary part results in larger uncertainties to the MAC as compared to the real part. The relative
- 281 uncertainty of the derived m_{BC} was within 35%.
- 282 This study provides a new way to derive m_{BC} from σ_{ab} for the widely-used filter-based measurements. This research deepens our
- 283 understanding of the uncertainty in measured m_{BC} caused by the utilization of a constant MAC and illustrates the great necessity to
- take the variation of MAC into account. The new method improves the measurements of BCPMSD and is further beneficial to the
- 285 evaluation of BC radiative forcing.
- 286 Data availability
- 287 The measurement data involved in this study are available upon request to the authors.
- 288 Author contributions
- 289 CZ determined the main goal of this study. WZ and WT designed the methods. WZ carried them out and prepared the paper with
- 290 contributions from all co-authors.
- 291 Competing interests
- 292 The authors declare that they have no conflict of interest.
- 293 References
- 294 Adler, G., Riziq, A. A., Erlick, C., and Rudich, Y.: Effect of intrinsic organic carbon on the optical properties of fresh diesel soot,
- Proceedings of the National Academy of Sciences of the United States of America, 107, 6699-6704, 10.1073/pnas.0903311106,
- 296 2010.





- 297 Arnott, W. P., Hamasha, K., Moosmuller, H., Sheridan, P. J., and Ogren, J. A.: Towards aerosol light-absorption measurements with
- 298 a 7-wavelength Aethalometer: Evaluation with a photoacoustic instrument and 3-wavelength nephelometer, Aerosol Science and
- 299 Technology, 39, 17-29, 10.1080/027868290901972, 2005.
- 300 Bond, T. C., Anderson, T. L., and Campbell, D.: Calibration and intercomparison of filter-based measurements of visible light
- 301 absorption by aerosols, Aerosol Science and Technology, 30, 582-600, 10.1080/027868299304435, 1999.
- 302 Bond, T. C., and Bergstrom, R. W.: Light absorption by carbonaceous particles: An investigative review, Aerosol Science and
- 303 Technology, 40, 27-67, 10.1080/02786820500421521, 2006.
- 304 Bond, T. C., Habib, G., and Bergstrom, R. W.: Limitations in the enhancement of visible light absorption due to mixing state, J.
- 305 Geophys. Res.-Atmos., 111, 13, 10.1029/2006jd007315, 2006.
- 306 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J., Flanner, M. G., Ghan, S., Karcher, B., Koch,
- 307 D., Kinne, S., Kondo, Y., Quinn, P. K., Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin,
- 308 N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo,
- 309 T., Warren, S. G., and Zender, C. S.: Bounding the role of black carbon in the climate system: A scientific assessment, J. Geophys.
- 310 Res.-Atmos., 118, 5380-5552, 10.1002/jgrd.50171, 2013.
- 311 Boucher, O., D. Randall, P. Artaxo, C. Bretherton, G. Feingold, P. Forster, V.-M. Kerminen, Y. Kondo, H. Liao, U. Lohmann, P.
- 312 Rasch, S.K. Satheesh, S. Sherwood, B. Stevens, and Zhang, X. Y.: Clouds and Aerosols. In: Climate Change 2013: The Physical
- 313 Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change,
- 314 in, edited by: Stocker, T. F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley,
- 315 Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 571-657, 2013.
- 316 Brem, B. T., Gonzalez, F. C. M., Meyers, S. R., Bond, T. C., and Rood, M. J.: Laboratory-Measured Optical Properties of Inorganic
- 317 and Organic Aerosols at Relative Humidities up to 95%, Aerosol Science and Technology, 46, 178-190,
- 318 10.1080/02786826.2011.617794, 2012.
- 319 Cappa, C. D., Onasch, T. B., Massoli, P., Worsnop, D. R., Bates, T. S., Cross, E. S., Davidovits, P., Hakala, J., Hayden, K. L., Jobson,
- 320 B. T., Kolesar, K. R., Lack, D. A., Lerner, B. M., Li, S. M., Mellon, D., Nuaaman, I., Olfert, J. S., Petaja, T., Quinn, P. K., Song, C.,
- 321 Subramanian, R., Williams, E. J., and Zaveri, R. A.: Radiative Absorption Enhancements Due to the Mixing State of Atmospheric
- 322 Black Carbon, Science, 337, 1078-1081, 10.1126/science.1223447, 2012.
- 323 Cappa, C. D., Zhang, X. L., Russell, L. M., Collier, S., Lee, A. K. Y., Chen, C. L., Betha, R., Chen, S. J., Liu, J., Price, D. J., Sanchez,
- 324 K. J., McMeeking, G. R., Williams, L. R., Onasch, T. B., Worsnop, D. R., Abbatt, J., and Zhang, Q.: Light absorption by ambient
- 325 black and brown carbon and its dependence on black carbon coating state for two California, USA, cities in winter and summer, J.
- 326 Geophys. Res.-Atmos., 124, 1550-1577, 10.1029/2018jd029501, 2019.
- 327 Castagna, J., Calvello, M., Esposito, F., and Pavese, G.: Analysis of equivalent black carbon multi-year data at an oil pre-treatment
- 328 plant: Integration with satellite data to identify black carbon transboundary sources, Remote Sens. Environ., 235, 10,
- 329 10.1016/j.rse.2019.111429, 2019.





- 330 China, S., Scarnato, B., Owen, R. C., Zhang, B., Ampadu, M. T., Kumar, S., Dzepina, K., Dziobak, M. P., Fialho, P., Perlinger, J. A.,
- 331 Hueber, J., Helmig, D., Mazzoleni, L. R., and Mazzoleni, C.: Morphology and mixing state of aged soot particles at a remote marine
- free troposphere site: Implications for optical properties, Geophys. Res. Lett., 42, 1243-1250, 10.1002/2014gl062404, 2015.
- 333 Chung, C. E., Ramanathan, V., and Decremer, D.: Observationally constrained estimates of carbonaceous aerosol radiative forcing,
- 334 Proceedings of the National Academy of Sciences of the United States of America, 109, 11624-11629, 10.1073/pnas.1203707109,
- 335 2012.
- 336 Doran, J. C., Barnard, J. C., Arnott, W. P., Cary, R., Coulter, R., Fast, J. D., Kassianov, E. I., Kleinman, L., Laulainen, N. S., Martin,
- 337 T., Paredes-Miranda, G., Pekour, M. S., Shaw, W. J., Smith, D. F., Springston, S. R., and Yu, X. Y.: The T1-T2 study: evolution of
- aerosol properties downwind of Mexico City, Atmospheric Chemistry and Physics, 7, 1585-1598, 10.5194/acp-7-1585-2007, 2007.
- 339 Fuller, K. A., Malm, W. C., and Kreidenweis, S. M.: Effects of mixing on extinction by carbonaceous particles, J. Geophys. Res.-
- 340 Atmos., 104, 15941-15954, 10.1029/1998jd100069, 1999.
- 341 Gunter, R. L., Hansen, A. D. A., Boatman, J. F., Bodhaine, B. A., Schnell, R. C., and Garvey, D. M.: Airborne measurement of
- 342 aerosol optical-properties over south-central New-Mexico, Atmospheric Environment Part a-General Topics, 27, 1363-1368,
- 343 10.1016/0960-1686(93)90262-w, 1993.
- Hansen, A. D. A., Rosen, H., and Novakov, T.: The aethalometer an instrument for the real-time measurement of optical-absorption

345 by aerosol-particles, Sci. Total Environ., 36, 191-196, 10.1016/0048-9697(84)90265-1, 1984.

- 346 Helin, A., Niemi, J. V., Virkkula, A., Pirjola, L., Teinila, K., Backman, J., Aurela, M., Saarikoski, S., Ronkko, T., Asmi, E., and
- 347 Timonen, H.: Characteristics and source apportionment of black carbon in the Helsinki metropolitan area, Finland, Atmospheric
- 348 Environment, 190, 87-98, 10.1016/j.atmosenv.2018.07.022, 2018.
- 349 Jacobson, M. Z.: A physically-based treatment of elemental carbon optics: Implications for global direct forcing of aerosols,
- 350 Geophys. Res. Lett., 27, 217-220, 10.1029/1999gl010968, 2000.
- Jacobson, M. Z.: Strong radiative heating due to the mixing state of black carbon in atmospheric aerosols, Nature, 409, 695-697,
 10.1038/35055518, 2001.
- 353 Khalizov, A. F., Xue, H. X., Wang, L., Zheng, J., and Zhang, R. Y.: Enhanced light absorption and scattering by carbon soot aerosol
- internally mixed with sulfuric acid, Journal of Physical Chemistry A, 113, 1066-1074, 10.1021/jp807531n, 2009.
- 355 Lack, D. A., and Cappa, C. D.: Impact of brown and clear carbon on light absorption enhancement, single scatter albedo and
- absorption wavelength dependence of black carbon, Atmospheric Chemistry and Physics, 10, 4207-4220, 10.5194/acp-10-4207-
- 357 2010, 2010.
- 358 Lack, D. A., Langridge, J. M., Bahreini, R., Cappa, C. D., Middlebrook, A. M., and Schwarz, J. P.: Brown carbon and internal mixing
- in biomass burning particles, Proceedings of the National Academy of Sciences of the United States of America, 109, 14802-14807,
- 360 10.1073/pnas.1206575109, 2012.
- 361 Li, Z. J., Tan, H. B., Zheng, J., Liu, L., Qin, Y. M., Wang, N., Li, F., Li, Y. J., Cai, M. F., Ma, Y., and Chan, C. K.: Light absorption
- 362 properties and potential sources of particulate brown carbon in the Pearl River Delta region of China, Atmospheric Chemistry and





- 363 Physics, 19, 11669-11685, 10.5194/acp-19-11669-2019, 2019.
- 364 Liu, C., Yin, Y., Hu, F. C., Jin, H. C., and Sorensen, C. M.: The effects of monomer size distribution on the radiative properties of
- 365 black carbon aggregates, Aerosol Science and Technology, 49, 928-940, 10.1080/02786826.2015.1085953, 2015.
- 366 Liu, C., Chung, C. E., Yin, Y., and Schnaiter, M.: The absorption Angstrom exponent of black carbon: from numerical aspects,
- 367 Atmospheric Chemistry and Physics, 18, 6259-6273, 10.5194/acp-18-6259-2018, 2018.
- 1368 Liu, H., Pan, X. L., Wu, Y., Wang, D. W., Tian, Y., Liu, X. Y., Lei, L., Sun, Y. L., Fu, P. Q., and Wang, Z. F.: Effective densities of
- 369 soot particles and their relationships with the mixing state at an urban site in the Beijing megacity in the winter of 2018, Atmospheric
- 370 Chemistry and Physics, 19, 14791-14804, 10.5194/acp-19-14791-2019, 2019.
- 371 Ma, N., Zhao, C. S., Nowak, A., Muller, T., Pfeifer, S., Cheng, Y. F., Deng, Z. Z., Liu, P. F., Xu, W. Y., Ran, L., Yan, P., Gobel, T.,
- 372 Hallbauer, E., Mildenberger, K., Henning, S., Yu, J., Chen, L. L., Zhou, X. J., Stratmann, F., and Wiedensohler, A.: Aerosol optical
- 373 properties in the North China Plain during HaChi campaign: an in-situ optical closure study, Atmospheric Chemistry and Physics,
- 374 11, 5959-5973, 10.5194/acp-11-5959-2011, 2011.
- 375 Ma, N., Zhao, C. S., Muller, T., Cheng, Y. F., Liu, P. F., Deng, Z. Z., Xu, W. Y., Ran, L., Nekat, B., van Pinxteren, D., Gnauk, T.,
- 376 Mueller, K., Herrmann, H., Yan, P., Zhou, X. J., and Wiedensohler, A.: A new method to determine the mixing state of light absorbing
- 377 carbonaceous using the measured aerosol optical properties and number size distributions, Atmospheric Chemistry and Physics, 12,
- 378 2381-2397, 10.5194/acp-12-2381-2012, 2012.
- 379 Majdi, M., Kim, Y., Turquety, S., and Sartelet, K.: Impact of mixing state on aerosol optical properties during severe wildfires over
- 380 the Euro-Mediterranean region, Atmospheric Environment, 220, 11, 10.1016/j.atmosenv.2019.117042, 2020.
- 381 Martins, J. V., Artaxo, P., Liousse, C., Reid, J. S., Hobbs, P. V., and Kaufman, Y. J.: Effects of black carbon content, particle size,
- 382 and mixing on light absorption by aerosols from biomass burning in Brazil, J. Geophys. Res.-Atmos., 103, 32041-32050,
- 383 10.1029/98jd02593, 1998.
- 384 Moffet, R. C., O'Brien, R. E., Alpert, P. A., Kelly, S. T., Pham, D. Q., Gilles, M. K., Knopf, D. A., and Laskin, A.: Morphology and
- 385 mixing of black carbon particles collected in central California during the CARES field study, Atmospheric Chemistry and Physics,
- 386 16, 14515-14525, 10.5194/acp-16-14515-2016, 2016.
- 387 Nakayama, T., Ikeda, Y., Sawada, Y., Setoguchi, Y., Ogawa, S., Kawana, K., Mochida, M., Ikemori, F., Matsumoto, K., and Matsumi,
- 388 Y.: Properties of light-absorbing aerosols in the Nagoya urban area, Japan, in August 2011 and January 2012: Contributions of brown
- 389 carbon and lensing effect, J. Geophys. Res.-Atmos., 119, 12721-12739, 10.1002/2014jd021744, 2014.
- 390 Ning, Z., Chan, K. L., Wong, K. C., Westerdahl, D., Mocnik, G., Zhou, J. H., and Cheung, C. S.: Black carbon mass size distributions
- 391 of diesel exhaust and urban aerosols measured using differential mobility analyzer in tandem with Aethalometer, Atmospheric
- 392 Environment, 80, 31-40, 10.1016/j.atmosenv.2013.07.037, 2013.
- 393 Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., Shang, D. J., Zamora, M. L., Zeng, L. M., Shao, M., Wu, Y. S., Zheng, J., Wang,
- 394 Y., Glen, C. R., Collins, D. R., Molina, M. J., and Zhang, R. Y.: Markedly enhanced absorption and direct radiative forcing of black
- carbon under polluted urban environments, Proceedings of the National Academy of Sciences of the United States of America, 113,





- 396 4266-4271, 10.1073/pnas.1602310113, 2016.
- 397 Petzold, A., and Schonlinner, M.: Multi-angle absorption photometry a new method for the measurement of aerosol light absorption
- and atmospheric black carbon, Journal of Aerosol Science, 35, 421-441, 10.1016/j.jaerosci.2003.09.005, 2004.
- 399 Ram, K., and Sarin, M. M.: Absorption Coefficient and Site-Specific Mass Absorption Efficiency of Elemental Carbon in Aerosols
- 400 over Urban, Rural, and High-Altitude Sites in India, Environmental Science & Technology, 43, 8233-8239, 10.1021/es9011542,
- 401 2009.
- 402 Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to black carbon, Nature Geoscience, 1, 221-227,
- 403 10.1038/ngeo156, 2008.
- 404 Sandradewi, J., Prevot, A. S. H., Szidat, S., Perron, N., Alfarra, M. R., Lanz, V. A., Weingartner, E., and Baltensperger, U.: Using
- 405 aerosol light absorption measurements for the quantitative determination of wood burning and traffic emission contributions to
- 406 particulate matter, Environmental Science & Technology, 42, 3316-3323, 10.1021/es702253m, 2008.
- 407 Schmid, O., Artaxo, P., Arnott, W. P., Chand, D., Gatti, L. V., Frank, G. P., Hoffer, A., Schnaiter, M., and Andreae, M. O.: Spectral
- 408 light absorption by ambient aerosols influenced by biomass burning in the Amazon Basin. I: Comparison and field calibration of
- 409 absorption measurement techniques, Atmospheric Chemistry and Physics, 6, 3443-3462, 10.5194/acp-6-3443-2006, 2006.
- 410 Schwarz, J. P., Gao, R. S., Fahey, D. W., Thomson, D. S., Watts, L. A., Wilson, J. C., Reeves, J. M., Darbeheshti, M., Baumgardner,
- 411 D. G., Kok, G. L., Chung, S. H., Schulz, M., Hendricks, J., Lauer, A., Karcher, B., Slowik, J. G., Rosenlof, K. H., Thompson, T. L.,
- 412 Langford, A. O., Loewenstein, M., and Aikin, K. C.: Single-particle measurements of midlatitude black carbon and light-scattering
- 413 aerosols from the boundary layer to the lower stratosphere, J. Geophys. Res.-Atmos., 111, 15, 10.1029/2006jd007076, 2006.
- 414 Schwarz, J. P., Spackman, J. R., Fahey, D. W., Gao, R. S., Lohmann, U., Stier, P., Watts, L. A., Thomson, D. S., Lack, D. A., Pfister,
- 415 L., Mahoney, M. J., Baumgardner, D., Wilson, J. C., and Reeves, J. M.: Coatings and their enhancement of black carbon light
- 416 absorption in the tropical atmosphere, J. Geophys. Res.-Atmos., 113, 10, 10.1029/2007jd009042, 2008.
- 417 Sharma, S., Brook, J. R., Cachier, H., Chow, J., Gaudenzi, A., and Lu, G.: Light absorption and thermal measurements of black
- 418 carbon in different regions of Canada, J. Geophys. Res.-Atmos., 107, 11, 10.1029/2002jd002496, 2002.
- 419 Shiraiwa, M., Kondo, Y., Moteki, N., Takegawa, N., Miyazaki, Y., and Blake, D. R.: Evolution of mixing state of black carbon in
- 420 polluted air from Tokyo, Geophys. Res. Lett., 34, 5, 10.1029/2007gl029819, 2007.
- 421 Shiraiwa, M., Kondo, Y., Iwamoto, T., and Kita, K.: Amplification of Light Absorption of Black Carbon by Organic Coating, Aerosol
- 422 Science and Technology, 44, 46-54, 10.1080/02786820903357686, 2010.
- 423 Sorensen, C. M.: Light scattering by fractal aggregates: A review, Aerosol Science and Technology, 35, 648-687,
- 424 10.1080/027868201316900007, 2001.
- 425 Stephens, M., Turner, N., and Sandberg, J.: Particle identification by laser-induced incandescence in a solid-state laser cavity, Appl.
- 426 Optics, 42, 3726-3736, 10.1364/ao.42.003726, 2003.
- 427 Truex, T. J., and Anderson, J. E.: Mass monitoring of carbonaceous aerosols with a spectrophone, Atmospheric Environment, 13,
- 428 507-509, 10.1016/0004-6981(79)90143-4, 1979.





- 429 Wild, M.: Englightening global dimming and brightening, Bulletin of the American Meteorological Society, 93, 27-37,
- 430 10.1175/bams-d-11-00074.1, 2012.
- 431 Zhang, R. Y., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H. X., and McMurry, P. H.: Variability in morphology, hygroscopicity, and
- 432 optical properties of soot aerosols during atmospheric processing, Proceedings of the National Academy of Sciences of the United
- 433 States of America, 105, 10291-10296, 10.1073/pnas.0804860105, 2008.
- 434 Zhang, Y. X., Zhang, Q., Cheng, Y. F., Su, H., Li, H. Y., Li, M., Zhang, X., Ding, A. J., and He, K. B.: Amplification of light
- 435 absorption of black carbon associated with air pollution, Atmospheric Chemistry and Physics, 18, 9879-9896, 10.5194/acp-18-9879-
- 436 2018, 2018.
- 437 Zhao, G., Tan, T. Y., Zhao, W. L., Guo, S., Tian, P., and Zhao, C. S.: A new parameterization scheme for the real part of the ambient
- 438 urban aerosol refractive index, Atmospheric Chemistry and Physics, 19, 12875-12885, 10.5194/acp-19-12875-2019, 2019a.
- 439 Zhao, G., Tao, J. C., Kuang, Y., Shen, C. Y., Yu, Y. L., and Zhao, C. S.: Role of black carbon mass size distribution in the direct
- 440 aerosol radiative forcing, Atmospheric Chemistry and Physics, 19, 13175-13188, 10.5194/acp-19-13175-2019, 2019b.
- 441

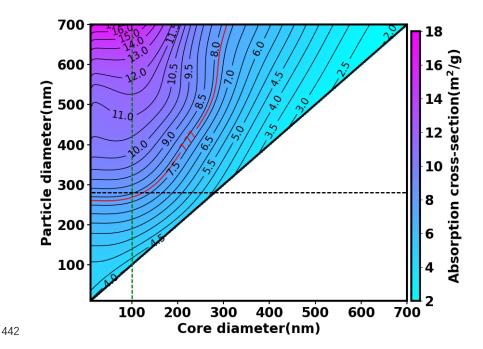
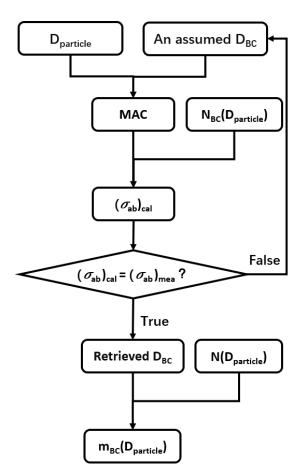


Figure 1. Variations in MAC as a function of D_{BC} and D_{particle}, calculated by the concentric core-shell Mie model at the wavelength of 880 nm. The red solid line is the constant MAC value used in AE33. The bold black solid line is the 1:1 line that presents the variations in MAC for pure BC particles with different D_{BC}. The horizontal black dashed line distinguishes

- 446 particles with a diameter of 280 nm while the vertical green dashed line indicates a D_{BC} of 100 nm.
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449 Figure 2. Schematic diagram of the iterative algorithm for retrieving the mBC at a fixed particle diameter based on the look-

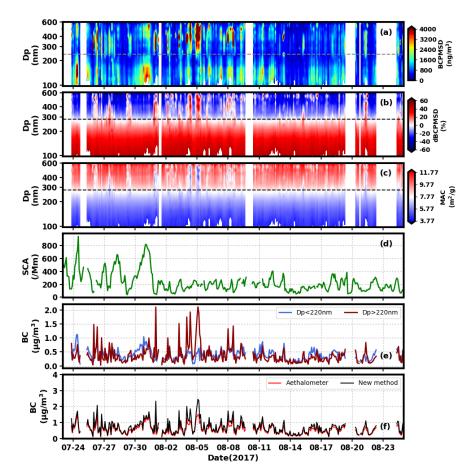
450 up table of MAC, particle size and core size. (σ_{ab})_{cal} and (σ_{ab})_{mea} represent calculated and measured absorption coefficients,

451 respectively. NBC(D_{particle}) indicates the number concentration of particle containing BC at the given D_{particle}.

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Figure 3. Time series of (a) the BCPMSD derived from the newly proposed method (dashed line indicates the particle size of 240 nm); (b) relative deviations between BCPMSD derived from the new method and a constant MAC of 7.77 m²/g (dashed line indicates the particle size of 280 nm); (c) the size-resolved MAC determined during the process of retrieving BCPMSD (dashed line indicates the particle size of 280 nm); (d) the scattering coefficients simultaneously measured with the sizeresolved σ_{ab} ; (e) the m_{BC} for particles smaller than 280 nm and larger than 280 nm; and (f) the m_{BC} determined by the new method and the constant MAC of 7.77 m²/g.



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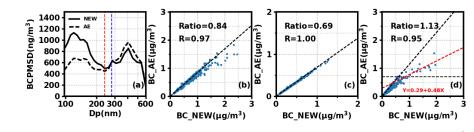


Figure 4. Comparison between the newly proposed method and using a constant MAC of 7.77 m²/g in the derived results of
(a) the BCPMSD (the dashed black line shows the results of AE33 while the solid black line represents the results from the
new method; the dashed red line represents the split line (diameter of 240 nm) between finer mode and coarser mode for





BCPMSD; the dashed blue line indicates the split line (280 nm of diameter) between the opposite tendencies of deviations in the m_{BC} calculated from the new method and the aethalometer); (b) the bulk m_{BC} for particles ranging from 97 nm to 602 nm; (c) the m_{BC} for the finer mode (97 – 280 nm); (d) the m_{BC} for the coarser mode (280 – 606 nm); the dashed black line represents boundary of 0.7 µg/m³ and the red dashed line is the regression line of the m_{BC} derived from AE and the new method when m_{BC} is larger than 0.7 µg/m³.



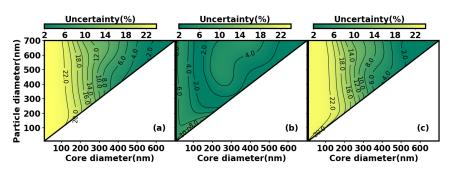


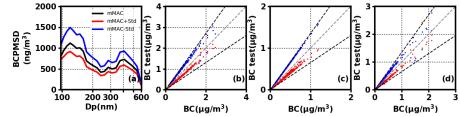
Figure 5. Uncertainty in MAC of BC when (a) real part of RI ranges from 1.5 to 2.0 and imaginary part ranges from 0.5 to

473 1.1; (b) real part of RI ranges from 1.5 to 2.0 and imaginary part is fixed at 0.8 and (c) real part of RI is fixed at 1.75 and
474 imaginary part ranges from 0.5 to 1.1. The bold black solid line is the 1:1 line and presents the uncertainty of MAC for pure

475 **BC particles with different RI.**

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Figure 6. (a) The BCPMSD calculated by using the look up table with mean MAC (black line), mean MAC plus the corresponding standard deviation (red line) and mean MAC minus the corresponding standard deviation (blue line); the m_{BC} derived by the look up table with mean MAC versus those derived by the look up table with mean MAC plus standard deviation (red dots) or mean MAC minus standard deviation (blue dots) for (b) aerosol particles ranging from 97–602 nm; (c) aerosol particles ranging from 97–240 nm (finer mode); and (d) aerosol particles ranging from 240–602 nm (coarser mode). The dashed black line represents the 35% deviation from the 1:1 line (dashed grey lines).

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