Anonymous Referee #1

Received and published: 6 August 2019

Authors' response: We are grateful to Anonymous Referee #1 for the time and useful comments. We have made point-to-point responses to each question in blue text below. Please note that line numbers correspond to the revised manuscript.

(I apologize in advance for not figuring out the LaTeX commands for "del" and "per mil" symbols for my comments.)

This paper discusses the results of measurements of the isotopic composition of NOx, HONO, and particle nitrate emission from laboratory biomass burns done as part of the FIREX campaign. This represents the first measurements of the isotopic composition of HONO from biomass burning, a critical component of the smoke radical budget in the ambient atmosphere. They find that the isotopic composition of the HONO and NOx is well corelated, suggesting they are formed via similar pathways.

The paper is very well-written and while I am not competent to evaluate the analytical chemistry methods used, they are described thoroughly in the paper and it is clear the authors put a lot of thought into their experimental setup. The results are clearly explained and the conclusions follow from the evidence presented in the paper. Overall, this is a very good manuscript and should be published after minor revisions to address my minor concerns listed below.

L20: Missing and "and" before "proton-transfer-reaction"

Authors' response:

Corrected. Thank you!

L31: "connected via formation pathways" is a very vague phrase. Can you be more specific about how they are linked here?

Authors' response:

Modified. Please see L23-24, "NOx and HONO are connected via formation pathways" is changed to "HONO is directly formed via subsequent chain reactions of NO_x emitted from biomass combustion"

L33: It seems odd to mention the importance of your particle nitrate measurements here, since you didn't think they were important enough to mention the quantitative results in the previous paragraph. Can you please add those results to the abstract?

Authors' response:

We added the isotopic composition results for particulate nitrate in the abstract. Please see L24-26.

L38-39: Why are the del18O measurements for each species expected to track with the influence of ozone, photochemistry and nighttime chemistry? That isn't obvious to me.

Authors' response:

O₃ has uniquely high δ^{18} O value (~110‰) that can be distinguished from other oxidants including O₂, OH, and HO₂ (< ~20‰), therefore formation pathways of secondary products such as NO₂, HONO, HNO₃ and particulate nitrate may be distinguished with δ^{18} O value of these products. For example, NO₂ formed from NO+O₃ will exhibit much higher δ^{18} O-NO₂ than that formed from NO+HO₂/RO₂. Nighttime HNO₃ is mainly formed from NO+O₃ \rightarrow NO₂, NO₂+O₃ \rightarrow NO₃, NO₂+ NO₃ \rightarrow N₂O₅, and N₂O₅+H₂O \rightarrow 2HNO₃. Therefore, HNO₃ formed during the night has higher δ^{18} O than during the day.

L49: While everyone knows what cooking and heating are, "prescribed" is an uncommon term and should be defined. I'd also refer to "prescribed burning" to match the "ing"s in the other terms.

Authors' response:

Corrected and "prescribed burning" is defined in the text (L49). Thanks!

L93-96: This sentence is hard to understand, with many "and"s stringing things together. Can you please rephrase, maybe splitting into two sentences to make the meaning clearer?

Authors' response:

This sentence has been split in to two sentences as shown in L94-97. Thanks!

L152-154: I have no idea how these acronyms were made – how did "subalpine fir" become "ABLA"? Not critical as you define them, but it was a strange choice that left me constantly flipping back to understand what was burning.

Authors' response:

This is a good point. The acronyms are defined based on Latin names of the vegetation species, as you can see in this example (<u>https://plants.usda.gov/core/profile?symbol=PICO</u>)

For readers' convenience, all the acronyms have been replaced by corresponding names.

L236-237: Please describe how you corrected the NOx for HONO interference, as you say on L235 that HONO is partially, not completely, converted on the catalyst.

Authors' response:

Thank you for asking this question. "Partially" was used in the original manuscript to discuss the NOy interference species in a qualitative and conservative way. The catalytic

efficiency of NOy species other than NO2 to NO were not measured for this particular Thermo NOx analyzer, however a number of previous works found the HONO (e.g. Febo et al 1995) conversion efficiency is high and deemed the conversion efficiency 100%. We correct the NOx concentration by subtracting mean HONO concentration during each period from NOx concentration, and this provided an approximate lower limit of NOx concentration with upper limit of HONO concentration.

L341-342: I'm assuming the bacteria preserves the isotope ratios and/or you can correct for the effect of the bacteria, but that isn't obvious to me, so you might want to clarify that here.

Authors' response:

The bacteria method can preserve the isotopic signature of NO_3^- and/or NO_2^- so I replace "quantitative" with "complete" in the original sentence (L321). Thanks for the suggestion.

L388-394: Your explanation for the low HONO values in Fire #12 being linked to smoldering combustion and lower MCE (0.87) isn't very convincing, as Fire #15, which you note on L385-387 had the highest HONO, had an MCE of 0.89, as did Fire #17. My guess is you don't have enough information to really explain why Fire 12 was anomalously low in HONO, but in any case, you need to revise this section to include the caveat about the MCE of Fire #15.

Authors' response:

This is a really great question and thank you!

We agree with the reviewer's comment on MCE and we note for fire #12 the strong smoldering resulted in clogging of the inlet filter of the MC/IC that significantly impact the measured HONO concentration by MC/IC. Accordingly, we added the following (L369-372 and L378-382 respectively):

"We note that fire no. 12 has the smallest MCE value of 0.868 (FIREX, 2016), and an abnormal flow rate (less than half of the typical flow rate during all other measurements) due to the inlet filter clogging from extraordinarily large particulate loadings."

"Although fires no. 15 and no. 17 have relatively low MCE (~0.89), the pulse of HONO in first 5-10 minutes suggest an active flaming phase followed by longer smoldering phase. This indicates both fires had combustion conditions that consisted of a mixture of flaming and smoldering, and thus significant HONO was still produced."

L435 and elsewhere: Since you have both positive and negative values, I'd use the word "to" instead of the n-dash symbol to connect the ranges. i.e. "-4.3 per mil to +7.0 per mil" instead of "-4.3 per mil - +7.0 per mil."

Authors' response:

All "-"s have been replaced by "to" when isotopic ranges are discussed. Thanks!

Data availability: Consider using the CERN Zenodo archive (zenodo.org) or similar free service to store the data in a public repository with a unique DOI.

Authors' response:

Thank you for this information! As expected by the funding of our project, our data are archived in the NOAA database, which is also public.

Figure 1 caption: Mention that the HONO and HNO3 in these plots was measured using the MC/IC method from the text

Authors' response:

Corrected. Thanks!

Table S1 caption: Please make a more descriptive caption.

Authors' response:

A new descriptive caption has been added to Table S1. Thank you!

Table S2 caption: Please explain why some data are missing (below detection limit? Instrument error?)

Authors' response:

The missing data are results of instrumental issues and this has been clarified in the caption of Table S2. Thank you!

Figure S1 caption: Please clarify what the p values are for, e.g., the slope of a linear correlation?

Authors' response:

The p values are for the slope of a linear correlation. This has been clarified in the caption of Figure S1 (now Figure S2). Thank you!

Figure S2 caption: The caption is hard to understand. Try "Linear regression between (a) del15N-HONO and del15N-biomass (equation) and (b). . ." instead.

Authors' response:

The caption of Figure S2 (now Figure S3) has been modified based on the reviewer's suggestion. Thank you!

Isotopic characterization of nitrogen oxides (NO_x), nitrous acid (HONO), and nitrate (NO₃⁻(p)) from laboratory biomass burning during FIREX

Jiajue Chai¹, David J. Miller^{1,a}, Eric Scheuer², Jack Dibb², Vanessa Selimovic³, Robert Yokelson³, Kyle J. Zarzana^{4,5,b}, Steven S. Brown^{4,6}, Abigail R. Koss^{4,5,6,c}, Carsten Warneke^{5,6}, Meredith Hastings¹

- 1. Department of Earth, Environmental and Planetary Sciences, and Institute at Brown for Environment and Society, Brown University, Providence, RI, USA
- 2. Institute for the Study of Earth, Ocean and Space, University of New Hampshire, Durham, NH, USA
- 3. Department of Chemistry, University of Montana, Missoula, USA
- 4. Chemical Sciences Division, NOAA Earth System Research Laboratory, Boulder, CO, USA
- 5. Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, CO, USA
- 6. Department of Chemistry, University of Colorado, Boulder, CO, USA
- a. Now at: Environmental Defense Fund, Boston, MA, USA
- b. Now at: Department of Chemistry, University of Colorado, Boulder, CO, USA
- c. Department of Civil and Environmental Engineering, Massachusetts Institute of Technology, Cambridge, MA, USA

Correspondence: Jiajue Chai (jiajue_chai@brown.edu)

1 Abstract

2

3 New techniques have recently been developed and applied to capture reactive nitrogen 4 species including nitrogen oxides ($NO_x = NO + NO_2$), nitrous acid (HONO), nitric acid 5 (HNO₃) and particulate nitrate (NO₃(p)), for accurate measurement of their isotopic composition. Here, we report - for the first time - the isotopic composition of HONO 6 7 from biomass burning (BB) emissions collected during the Fire Influence on Regional to 8 Global Environments Experiment (FIREX, later evolved into FIREX-AQ) laboratory 9 experiments at the Missoula Fire Science Laboratory in the fall of 2016. Using our newly 10 developed annular denuder system (ADS), which was verified to completely capture 11 HONO associated with BB via comparison with 4 other high time resolution 12 concentration measurement techniques, including mist chamber/ion chromatography 13 (MC/IC), open-path Fourier transform infrared spectroscopy (OP-FTIR), cavity enhanced 14 spectroscopy (CES), and proton-transfer-reaction time-of-flight mass spectrometer (PTR-15 ToF).

16

In 20 "stack" fires (direct emission within ~5 seconds of production by the fire) that 17 burned various biomass materials from the Western U.S., δ^{15} N-NO_x ranges from -4.3 ‰ 18 19 to +7.0 %, falling near the middle of the range reported in previous work. The first measurements of δ^{15} N-HONO and δ^{18} O-HONO in biomass burning smoke reveal a range 20 of -5.3‰ to +5.8‰ and +5.2‰ to +15.2‰, respectively. Both HONO and NO_x are 21 sourced from N in the biomass fuel and δ^{15} N-HONO and δ^{15} N-NO_x are strongly 22 23 correlated ($R^2 = 0.89$, p<0.001), suggesting HONO is directly formed via subsequent 24 chain reactions of NO_x emitted from biomass combustion. Only 5 of 20 NO₃ (p) samples had a sufficient amount for isotopic analysis, and showed $\delta^{15}N$ and $\delta^{18}O$ of NO₃ (p) range 25 26 from -10.6‰ to -7.4 ‰ and +11.5‰ to +14.8‰ respectively.

27

Our δ^{15} N of NO_x, HONO and NO₃(p) ranges can serve as important biomass burning 28 29 source signatures, useful for constraining emissions of these species in environmental 30 applications. The δ^{18} O of HONO and NO₃⁻ obtained here verify our method is capable of determining the oxygen isotopic composition in BB plumes. The δ^{18} O for both of these 31 32 species reflect laboratory conditions (i.e. a lack of photochemistry), and would be 33 expected to track with the influence of different oxidation pathways in real environments. 34 The methods used in this study will be further applied in future field studies to 35 quantitatively track reactive nitrogen cycling in fresh and aged Western US wildfire 36 plumes.

- 37
- 38
- 39 40
- 41
- 42
- 43
- 44
- 45

47 **1 Introduction**

48 Biomass burning (BB), which occurs in both anthropogenic processes (e.g. cooking, 49 heating, and prescribed burning that is human controlled burning for management 50 purpose) and natural wildfire (lightning ignited vegetation burning), is a significant 51 source of atmospheric reactive nitrogen species, including nitrogen oxides (NO_x = NO + 52 NO₂), nitrous acid (HONO), nitric acid (HNO₃), particulate nitrate (NO₃⁻(p)), organic 53 nitrates, peroxyacyl nitrate (PAN) and ammonia (NH_3) that have major impacts on air 54 quality and climate from regional to global scales (Crutzen and Andreae, 1990). Globally, 55 biomass burning emits ~6 Tg of nitrogen oxides (NO_x = NO + NO₂) per year, 56 contributing at least 14% to total NO_x emissions (Jaeglé et al., 2005), with large 57 interannual and seasonal variation due to fire frequency and intensity (Jaffe and Briggs, 58 2012). Primarily emitted NO_x plays an important role in the photo-oxidation of volatile 59 and semi volatile organic compounds, which are present in high concentrations in BB 60 plumes, and strongly influences production of tropospheric ozone (O_3) and secondary 61 aerosols (Alvarado et al., 2015). In BB plumes, NO_x can be converted to PAN, which can 62 be transported long distances (100s to 1000s of km) in lofted plumes before rereleasing NO_x . Therefore, BB emitted NO_x could widely influence air quality downwind for days 63 64 to weeks (Val Martín et al., 2006; Ye et al., 2016). In addition, NO_x is also the major photochemical precursor of HNO_3 and $NO_3^{-}(p)$, which can be transported downwind, mix 65 66 with anthropogenic emissions, and impact air quality and ecosystem health (Hastings et 67 al., 2013).

68 HONO has been observed in BB plumes in both laboratory and field experiments, with 69 HONO mixing ratios in the range of \sim 5-33% of observed NO_x (Akagi et al., 2012, 2013; 70 Burling et al., 2010, 2011: Keene et al., 2006: Liu et al., 2016: Roberts et al., 2010: 71 Selimovic et al., 2018; Yokelson et al., 2007, 2009). Photolysis of HONO is a major OH 72 precursor in the daytime; therefore HONO plays an important role in photochemical 73 aging of BB plumes and atmospheric oxidation capacity at regional scales (Alvarado and 74 Prinn, 2009; Liu et al., 2016; Tkacik et al., 2017; Trentmann et al., 2005). HONO has 75 been proposed as a significant OH source in BB plumes and the inclusion of HONO in 76 photochemical models could explain much of the uncertainty in the modeled O₃ 77 (Alvarado et al., 2009; Alvarado and Prinn, 2009; Cook et al., 2007; Travis et al., 2016; 78 Trentmann et al., 2005).

79 Direct BB emission factor measurements of HONO and NO_x exhibit significant 80 uncertainties due to limited observations and large spatial and temporal variability of burning conditions, making it challenging to build an accurate inventory of BB emissions 81 82 relative to other major sources (Lapina et al., 2008). Emission factors vary and mainly 83 depend on 1) fuel nitrogen content (0.2 - 4%) by mass), which is a function of vegetation 84 type, and 2) modified combustion efficiency (MCE = Δ [CO₂]/(Δ [CO] + Δ [CO₂]) that is 85 determined by combustion conditions including fuel moisture, fuel load, temperature, 86 relative humidity, wind speed, and other meteorological parameters (Burling et al., 2010; 87 Jaffe and Briggs, 2012; Yokelson et al., 1996). Additionally, the temporal evolution of 88 HONO in BB plumes varies greatly in different fires and relative contributions from 89 direct emission versus NO₂ conversion to HONO remains unclear. For instance, 90 significant concentrations of HONO and correlation between HONO and NO₂ have been 91 observed in aged plumes, indicating the importance of heterogeneous conversion of NO₂-

to-HONO on BB aerosols (Nie et al., 2015). By contrast, no evidence was found for
secondary HONO formation in a BB plume during the Southeast Nexus Experiment
(Neuman et al., 2016). It is important to constrain HONO directly emitted from BB
compared to HONO formed during plume aging. This would reduce uncertainties
associated with the total HONO budget and increase our understanding of HONO
impacts on O₃ and secondary aerosol formation downwind of BB regions.

98

99 In an effort to better understand reactive nitrogen emissions and chemistry, especially for 100 HONO, new techniques have been developed to analyze the isotopic composition of 101 various species. Stable isotopes provide a unique approach of characterizing and tracking 102 various sources and chemistry for a species of interest (Hastings et al., 2013). Fibiger et 103 al. (2014) developed a method to quantitatively collect NO_x in solution as NO_3^- for 104 isotopic analysis, which has been verified to avoid any isotopic fractionation during 105 collection in both lab and field studies. This allows for high-resolution measurement of 106 δ^{15} N-NO_x in minutes to hours depending on ambient NO_x concentrations (δ^{15} N = $[({}^{15}N/{}^{14}N)_{sample}/({}^{15}N/{}^{14}N)_{air-N2} - 1] \times 1000\%$, and $\delta^{18}O = [({}^{18}O/{}^{16}O)_{sample}/({}^{18}O/{}^{16}O)_{VSMOW} - 1]$ 107 1] \times 1000‰ where VSMOW is Vienna Standard Mean Ocean Water). δ^{15} N has also been 108 109 used to track gaseous NO_x from a variety of major sources including emissions from 110 biomass burning (Fibiger and Hastings, 2016), vehicles (Miller et al., 2017), and 111 agricultural soils (Miller et al., 2018). Using this method, Fibiger and Hastings (2016) systematically investigated BB δ^{15} N-NO_x from different types of biomass from around 112 113 the world in a controlled environment during the fourth Fire Lab at Missoula Experiment 114 (FLAME-4). NO_x emissions collected both immediately from the BB source and 1-2 115 hours after the burn in a closed environment ranged from -7 to +12%, and primarily depended on the δ^{15} N of the biomass itself. BB emitted HONO isotopic composition has 116 117 never been measured before. Our recently developed method for HONO isotopic composition analysis (Chai and Hastings, 2018) enables us to not only characterize $\delta^{15}N$ 118 and δ^{18} O of HONO, but also explore the connection between δ^{15} N-NO_x and δ^{15} N-HONO. 119 120

121 The Fire Influence on Regional to Global Environments and Air Quality (FIREX-AQ) 122 investigates the influence of fires in the western U.S. on climate and air quality, via an 123 intensive, multi-platform, campaign. The first phase of FIREX-AQ took place at the US 124 Forest Service Fire Sciences Laboratory (FSL) in Missoula, Montana, in the fall of 2016, where we measured δ^{15} N-NO_x, δ^{15} N-HONO, δ^{18} O-HONO, δ^{15} N-NO₃(p), δ^{18} O-NO₃(p) 125 and δ^{15} N-biomass in 20 "stack burns" of a variety of fuels representative of northwestern 126 127 North America. Here we report on the results and explore relationships between the 128 isotopic composition of these reactive nitrogen species, as well as the corresponding 129 mixing ratios for HONO that were concurrently measured by a variety of techniques. 130 This work offers characterization and quantification of BB source signatures of these 131 species, which can be applied in the interpretation of observations in future field studies.

- 132
- 133
- 134
- 135
- 136
- 137

138 2. Experimental details

139 2.1 FIREX Fire Science Laboratory design

140

141 The room for controlled BB experiments is 12.5×12.5 m $\times 22$ m, with a continuously 142 weighed fuel bed at the center of the room. The combustion exhaust was vented at a constant flow rate (~3.3 m s⁻¹) through a 3.6 m diameter inverted funnel followed by a 143 144 1.6 m diameter stack, and collected at a platform 17 m above the fuel bed via sampling 145 ports that surround the stack, resulting in a transport time of ~ 5 s. Further details have 146 been described in the literature (Stockwell et al., 2014). All of our instruments for 147 sampling and online measurements were placed on the platform, which can accommodate 148 up to 1820 kg of equipment and operators. Measurements were focused on the "stack 149 burns", for which fires lasted a few minutes up to 40 minutes.

150

151 For this study, we investigated 20 stack fires of vegetation types abundant in the western 152 US, representing coniferous ecosystems, including ponderosa pine (PIPO), lodgepole 153 pine (PICO), Engelmann spruce (PIEN), Douglas-fir (PSME) and subalpine fir (ABLA), 154 with replicate burns for most of these types (Table 1). Some of the fires proceeded by 155 burning of an individual fuel component such as litter, canopy, duff and rotten logs. 156 Other fires simulated actual biomass in the coniferous ecosystem by mixing various fuel 157 components in realistically recreated ecosystem matrices using the first order fire effects 158 model (FOFEM) (Reinhardt et al., 1997).

- 159
- 160

161 **2.2 Instrumentation**162

163 **2.2.1** Collection of HONO, NO_x and nitrate for isotopic analysis

164 HONO was completely collected for isotopic analysis using an annular denuder system 165 (ADS) (Chai and Hastings, 2018). The ADS system deployed in this laboratory experiment consisted of a Teflon particulate filter, a Nylasorb filter to remove HNO₃, 166 167 followed by two annular denuders, each coated with a solution of 10 mL of Na₂CO₃ (1% 168 w/v) + glycerol (1% v/v) + methanol-H₂O solution (1:1 volume ratio) following a 169 standard EPA method. Methanol and glycerol are certified ACS plus with a purity of 170 \geq 99.8% and \geq 99.5%, respectively. After coating, the denuders are dried using zero air 171 and capped immediately. Within 6 hours after each collection, the coating was extracted 172 in 10 mL of ultrapure water (18.2 M Ω) in two sequential 5 mL extractions. The extracted 173 solution with a pH of ~10 was transported to Brown University for concentration and 174 isotopic analysis 3-14 days after the sampling. The timescales for sample extraction and 175 isotopic analysis preserve both the solution concentration and isotopic composition of 176 HONO in the form of nitrite (Chai and Hastings, 2018). The two-denuder set up allows 177 us to minimize the interference for both concentration and isotopic analysis from other N-178 containing species that could be trapped and form nitrite in residual amounts on the 179 denuders, especially NO₂. Our method development study showed NO₂ tends to absorb in 180 the same amount (difference <4%) on the walls of each denuder in a train setup, which is 181 consistent with other studies (Perrino et al., 1990; Zhou et al., 2018). On the basis of this validation, the second denuder extract is used to correct the first denuder extract for both 182 183 concentration and isotopic composition (Chai and Hastings, 2018). Note HONO levels were above the minimum detection limit (0.07 μ M) and the breakthrough amount of HONO threshold is far from being reached given the concentrations (Table 1), flow rate (~ 4 L/min) and collection times (5 - 40 min). The necessary minimum amount of nitrite collected for isotopic analysis is 10 - 20 nmol.

188

189 To avoid scrubbing of HONO, a flow meter (Omega) and the NO_x collection system for 190 analysis of δ^{15} N-NO_x are placed following the ADS (Fibiger et al., 2014; Fibiger and 191 Hastings, 2016; Wojtal et al., 2016). In brief, NO_x is collected in a solution containing 192 0.25 M KMnO₄ and 0.5 M NaOH which oxidizes NO and NO₂ to NO₃⁻ by pumping 193 sampled air through a gas washing bottle with a 65 Watts diaphragm vacuum pump. The 194 flow rate (\sim 4L/min with ±1% uncertainty) is controlled with a critical orifice inserted 195 between the pump and gas stream outlet, and is monitored and recorded with a flow 196 meter placed prior to the NO_x collector. The NO_x trapping solution blanks are also 197 collected every day to quantify background NO_3^- for concentration and isotopic blank 198 corrections. The Omega flow meter was calibrated with another flow meter (Dry Cal Pro) 199 by varying flow rates. Within a day after collection, we stabilized the samples in the wet 200 chemistry lab in the Fire Science Lab by adding 30% w/w H_2O_2 that reduces MnO_4^- to 201 MnO₂ precipitate before being shipped back to Brown University for further processing. 202 This effectively excludes the possible interferences from NH₃ that could be oxidized to 203 NO_3^- by MnO_4^- after a week (Miller et al. (2017) and references therein). The samples 204 were neutralized with 12.1 N HCl in the Brown lab, before concentration measurement 205 and isotopic analyses. NO_3^- on the upstream Millipore filters and HNO₃ from the Nylasorb filters, if there was any, were extracted by sonicating the filters in ~30 mL 206 207 ultrapure H₂O (18.2 MΩ) for 30 minutes. Samples with $[NO_3] > 1 \mu M$ were analyzed for 208 isotopic composition (concentration techniques detailed below).

209

All treated samples from both HONO collection and NO_x collection and their 210 corresponding blanks were analyzed offline for concentrations of NO₂⁻ and NO₃⁻ with a 211 WestCo SmartChem 200 Discrete Analyzer colorimetric system. The reproducibility of 212 the concentration measurement was $\pm 0.3 \ \mu mol \ L^{-1} \ (1\sigma)$ for NO₂⁻ and $\pm 0.4 \ \mu mol \ L^{-1}$ for 213 NO_3^- when a sample was repeatedly measured (n = 30). A detection limit of 0.07 µmol 214 L^{-1} for NO₂ and 0.1 µmol L^{-1} for NO₃ was determined, and no detectable nitrite or 215 nitrate was found in the blank denuder coating solution, whereas blank NO_3^{-1} 216 217 concentrations of $\sim 5 \mu M$ are typical for the NO_x collection method (Fibiger et al., 2014; 218 Wojtal et al., 2016). Note that NO₃⁻ concentration was measured on the ADS solutions to 219 verify whether and to what extent NO_2^- was oxidized to NO_3^- on denuder walls because 220 the denitrifier method converts both NO_3^- and NO_2^- to N_2O for isotopic analysis (see 221 below). In addition, samples collected with a mist chamber/ion chromatography system 222 (described in Sect. 2.2.2) were also tested for their concentrations and only those with 223 sufficient nitrite quantity were further analyzed for isotopic composition.

224

225 **2.2.2 NO_x and HONO online concentration measurement**

NO and NO_x concentrations were measured with a Thermo Scientific Model 42i chemiluminescence NO/NO_x analyzer, which is described in supplemental information. The NO_x measurement verified the concentration of the NO_x collected for isotopic analysis, shown in Table S3 and Figure S1. 230 HONO and HNO₃ concentrations were measured using the University of New 231 Hampshire's dual mist chamber/ion chromatograph system (Scheuer et al., 2003) with the 232 sampling inlet placed right next to that of the ADS. The dual channel IC system is custom 233 built using primarily Dionex analytical components. Briefly, automated syringe pumps 234 are used to move samples and standard solutions in a closed system, which minimizes 235 potential contamination. A concentrator column and 5 ml injections were used to improve 236 sensitivity. Eluents are purged and maintained under a pressurized helium atmosphere. 237 Background signal is minimized using electronic suppression (Dionex-ASRS). The 238 chromatography columns and detectors are maintained at 40 °C to minimize baseline 239 drifting. A tri-fluoro-acetate tracer spiked into the ultra-clean sampling water is used as 240 an internal tracer of sample solution volume, which can decrease due to evaporation in 241 the exhaust flow by 10-20% depending on the ambient conditions and length of the 242 sample integration interval. The spike was analyzed to correct the final mist chamber 243 sampled solution volume with an uncertainty of $\pm 3\%$. This system has been deployed to 244 various field studies for HONO measurement (Dibb et al., 2002; Stutz et al., 2010) and 245 showed reasonable intercomparison with other HONO measurement techniques (within 246 16% uncertainty) during the 2009 SHARP campaign in Houston (Pinto et al., 2014). The 247 detection limits for HNO₃ and HONO are 10 ppt for 5-minute sample integrations. 248 During the experiments, two mist chambers were operated to collect gas samples in 249 parallel, each with an integration interval of 5 minutes. One channel of the IC was 250 utilized for concentration measurement; in the other, the mist chamber's solution was 251 transferred into a sample bottle using the syringe pump, and the collected solution was 252 brought to Brown University for isotopic analysis of HNO₃ if sufficient amount (10-20 253 nmol) was collected for each sample.

254

255 In addition to MC/IC, the HONO mixing ratios were also measured using high time-256 resolution (~1 second) measurement techniques including open-path Fourier transform 257 infrared spectroscopy (OP-FTIR) (Selimovic et al., 2018), cavity enhanced spectrometer 258 (CES) (Min et al., 2016; Zarzana et al., 2018), and proton-transfer-reaction time-of-flight 259 mass spectrometer (PTR-ToF). Inlet ports of CES and PTR-ToF were placed 5' apart 260 from, but at the same height on the platform as those for ADS and MC/IC, while the OP-261 FTIR had an open path cell at the stack. The smoke has been shown to be well-mixed at 262 the sampling platform (Christian et al., 2004) and the mean HONO mixing ratios across 263 each fire obtained from the four techniques were compared with that retrieved from ADS 264 collection. This offers comprehensive verification of complete capture of HONO by ADS 265 that is extremely important for conserving the isotopic composition of HONO.

266 The details of OP-FTIR are described in previous works (Selimovic et al., 2018; Stockwell et al., 2014). The setup included a Bruker MATRIX-M IR cube spectrometer 267 268 with a mercury cadmium telluride (MCT) liquid-nitrogen-cooled detector interfaced with 269 a 1.6 m base open-path White cell. The white cell was positioned on the platform and its 270 open path spanned the width of the stack. This facilitates direct measurement across the rising emissions. The optical path length was set to 58 m. The IR spectra resolution was 271 0.67 cm⁻¹ from 600–4000 cm⁻¹. Pressure and temperature were continuously recorded 272 273 with a pressure transducer and two temperature sensors respectively, which were placed 274 adjacent to the White cell optical path. They were used for spectral analysis. Time 275 resolution for stack burns was approximately 1.37 s. The OP-FTIR measures CO₂, CO, 276 CH_4 , a series of volatile organic compounds and various reactive nitrogen species 277 (Selimovic et al., 2018). Mixing ratios of HONO were retrieved via multicomponent 278 fitting to a section of the mid-IR transmission spectra with a synthetic calibration 279 nonlinear least-squares method (Griffith, 1996; Yokelson et al., 2007), and both the 280 HITRAN spectral database and reference spectra recorded at the Pacific Northwest 281 National Laboratory (Rothman et al., 2009; Sharpe et al., 2004; Johnson et al., 2010, 282 2013) were used for the fitting. The uncertainty is $\sim 10\%$ for the HONO mixing ratio 283 measurement and the detection limit is no more than a few ppb as reported in previous 284 studies (Stockwell et al., 2014; Veres et al., 2010).

285 HONO measurements by cavity enhanced spectroscopy used the airborne cavity 286 enhanced spectrometer, ACES, recently described by Min et al. (2016). This instrument 287 consists of two channels, one measuring over the spectral range from 438-468 nm where 288 glyoxal (CHOCHO) and NO₂ have structured absorption bands, and one measuring over 289 from 361-389 nm, where HONO has structured absorption. In the HONO channel, light 290 from an LED centered at 368 nm and with an output power of 450 mW and collimated 291 with an off-axis parabolic collector illuminates the input mirror of a 48 cm optical cavity 292 formed from mirrors with a maximum reflectivity R = 99.98% at 375 nm. The effective 293 path length within the optical cavity is > 3 km over the region of greatest reflectivity. The 294 mirror reflectivity (effective path length) was calibrated from the difference in Rayleigh 295 scattering between Helium and zero air to provide an absolute calibration of the 296 instrument response. A fiber optic bundle collects light exiting the optical cavity and 297 transmits it to a grating spectrometer with a CCD detector, where it is spectrally 298 dispersed at a resolution of 0.8 nm. The resulting spectra are fit using DOASIS software 299 (Kraus, 2006) to determine trace gas concentrations, including NO₂, HONO and O₄. 300 Mixing ratios of NO₂ and HONO are reported at 1 s resolution, although the NO₂ 301 precision is higher in the 455 nm channel. The 1 Hz HONO precision is 800 pptv (2σ). 302 (The precision of the HONO instrument in ACES is somewhat degraded by the 303 optimization of 455 nm channel for glyoxal detection, which reduces the photon count 304 rate on the 368 nm channel.) The accuracy of the HONO measurement is 9%. Air was 305 sampled directly from stack at a height of 15 m above the fuel bed through a 1 m length 306 of ¹/₄" O.D. Teflon (FEP) tubing as described in Zarzana et al. (2018). The residence time 307 in the inlet and sample cells was < 1 s. Comparison between the ACES HONO and an 308 open path FTIR agreed to within 13% on average, and ACES HONO was well correlated with 1Hz measurements from a PTR-ToF ($r^2 = 0.95$) (Koss et al., 2018). 309

310 The PTR-ToF instrument used in the FIREX Fire Lab experiment is described in detail in 311 previous studies (Koss et al., 2018; Yuan et al., 2016). The PTR-ToF instrument is a 312 chemical ionization mass spectrometer typically using H_3O^+ reagent ions and a wide 313 range of trace gases can be detected in the range of tens to hundreds of parts per trillion 314 (pptv) for a 1 s measurement time. At the Fire Lab, PTR-ToF detected several inorganic 315 species including HONO with an uncertainty of 15%. HONO is detected at a lower 316 sensitivity than most trace gases in PTR-ToF, but mixing ratios for all fires were well 317 above the detection limit.

318

319 2.2.3 Isotopic composition measurements

The denitrifier method was used to perform nitrogen and oxygen isotope analyses 320 $(^{15}N/^{14}N, ^{18}O/^{16}O)$ of NO₃⁻ and/or NO₂⁻, by complete conversion to N₂O by denitrifying 321 bacteria P. aureofaciens (Casciotti et al., 2002; Sigman et al., 2001). The isotopic 322 323 composition of N₂O is then determined by a Thermo Finnegan Delta V Plus isotope ratio mass spectrometer at m/z 44, 45 and 46 for ${}^{14}N{}^{16}O$, ${}^{14}N{}^{15}N{}^{16}O$ and ${}^{14}N{}^{18}N{}^{18}O$. 324 respectively. Sample analyses were corrected against replicate measurements of the NO3 325 isotopic reference materials USGS34, USGS35, and IAEA-NO-3 (Böhlke et al., 2003). 326 Additional correction was performed for δ^{18} O-HONO following previous studies 327 (Casciotti et al., 2002, 2007; Chai and Hastings, 2018). Precisions for δ^{15} N-HONO, δ^{18} O-328 HONO and δ^{15} N-NO_x isotopic analysis across each of the entire methods are ±0.6‰, 329 $\pm 0.5\%$ and $\pm 1.3\%$, respectively (Chai and Hastings, 2018; Fibiger et al., 2014). δ^{18} O-330 N₂O from the NO_x collection samples was measured but is not reported as δ^{18} O-NO_x 331 because it is greatly impacted by MnO_4 oxidation and does not represent the $\delta^{18}O-NO_x$ in 332 the sample air. The total δ^{15} N of the starting biomass (δ^{15} N-biomass) was measured at the 333 Marine Biological Laboratory Ecosystems Center Stable Isotope Facility. The materials 334 335 measured for δ^{15} N-biomass (Table S1) cover most but not all the biomass types burned in 336 the experiments depending on availability of the leftover materials. Analyses were 337 conducted using a Europa ANCA-SL elemental analyzer-gas chromatograph preparation 338 system interfaced with a Europa 20–20 continuous-flow gas source stable isotope ratio 339 mass spectrometer. Analytical precision was $\pm 0.1\%$, based on replicate analyses of 340 international reference materials.

341 Collection time spanned the whole fire burning (5 min to 40 min) in order to maximize 342 the signal. We chose to report the samples whose concentrations are at least 30% above the 5 µM NO₃⁻ present in the blank KMnO₄ solution upon purchase (Fibiger et al., 2014), 343 344 such that the propagated error through the blank correction does not exceed the analytical precision of $\pm 1.5\%$ for δ^{15} N-NO_x. We found identical concentration and isotopic 345 346 signatures for both Fire Lab and Brown University Lab blanks, which ensures that no 347 additional NO₃⁻ contamination was introduced into the KMnO₄ solutions in the gas-348 washing bottle. In addition, fires with high particulate loading that resulted in >50%349 reduction in flow rate are not considered for isotopic analysis because the low flow rate 350 could induce incomplete collection with potential isotopic fractionation that might not 351 represent BB emissions.

352

353 **3. Results and discussion**

354

355 **3.1 Temporal evolution of HONO and HNO₃ from direct BB emissions**

356

357 The time series of HONO and HNO₃ concentrations measured by MC/IC at 5-minute 358 resolution for majority of the stack burns are shown in Fig. 1, and original data can be 359 found in the NOAA data archive (FIREX, 2016). HNO₃ concentrations were nearly two 360 orders of magnitude lower than typical HONO concentrations. The constant low 361 concentration of HNO₃ from fresh emissions across all fires is consistent with the 362 findings in Keene et al. (2006), confirming HNO₃ is not a primary reactive nitrogen 363 species in fresh smoke. Rather, it is largely produced secondarily in aged smoke and 364 nighttime chemistry. Both HONO and HNO₃ mixing ratios reach their peak in the first 365 five minutes, except for fire no. 12 (Engelmann spruce - duff), from which HONO 366 concentration remains nearly constant over the course of each fire, but much lower than 367 HONO concentration of the rest of the fires. The largest HONO and HNO₃ were emitted 368 from burning subalpine fir-Fish Lake canopy (fire no. 15), integrated concentration of up 369 to 177 ppby and 1.9 ppby in the first 5-minute sample, respectively. We note that fires no. 370 12 has the smallest MCE value 0.868 (FIREX, 2016), and abnormal flow rate (less than 371 half of the typical flow rate during all other measurements) due to the inlet filter clogging 372 from extraordinarily large particulate loadings. In general, the closer the MCE value is to 373 1, the more likely N-oxidation (e.g. NO_x and HONO) dominates over N-reduction (e.g. 374 NH₃ and HCN) as a result of flaming; when MCE approaches 0.8, more smoldering 375 occurs such that N-reduction becomes dominant (Ferek et al., 1998; Goode et al., 1999; 376 McMeeking et al., 2009; Yokelson et al., 1996, 2008). Accordingly, the smoldering 377 combustion condition of fire no. 12 leads to lower concentration of oxidized nitrogen 378 species than the rest of the fires in this study. Although fires no. 15 and no. 17 have 379 relatively low MCE (~0.89), the pulse of HONO in first 5-10 minutes suggest an active 380 flaming phase followed by longer smoldering phase. This indicates both fires had 381 combustion conditions that consisted of a mixture of flaming and smoldering, and thus 382 significant HONO was still produced. In addition, HONO/NO_x ratio ranged from 0.13 to 383 0.53 with a mean of 0.29 ± 0.12 (1 σ), comparable with previous results of laboratory 384 experiments (0.11 ± 0.04) and field experiments (0.23 ± 0.09) (Akagi et al., 2013; Burling 385 et al., 2010, 2011)

386

387 **3.2 Verification of ADS collected HONO concentration**

388 The HONO collected with the ADS represents a mean value over the course of each 389 entire burn. We first compare HONO concentration recovered from the ADS, denoted as 390 [HONO]_{ADS}, with that measured with the collocated MC/IC when both measurements 391 were available (Fig. 2). The comparison demonstrates good consistency across all fires, 392 with the [HONO]_{ADS} of all available fires falling within the first and third quartile of 393 MC/IC HONO data. Additionally, we made intercomparisons between [HONO]_{ADS} with 394 mean values of various high resolution methods including MC/IC, OP-FTIR, ACES and 395 PTR-ToF that are also available from the NOAA data archive (Fig. 3; FIREX, 2016). The 396 mean values used for the comparison are shown in Table S2. The linear regression results 397 for all four comparisons are:

398 399 $[HONO]_{ADS} = (1.07 \pm 0.24) [HONO]_{MCIC} - 0.72$ Eq. (1) $(R^2 = 0.63; p_{slope} < 0.001, p_{intercept} = 0.95);$ 400 401
$$\label{eq:constraint} \begin{split} & [HONO]_{ADS} = (1.07 \pm 0.08) \; [HONO]_{ACES} - 4.63 \\ & (R^2 = 0.95; \; p_{slope} < 1 \times 10^{-6}, \; p_{intercept} = 0.32); \end{split}$$
402 Eq. (2) 403 404 $[HONO]_{ADS} = (1.07 \pm 0.22) [HONO]_{FTIR} + 5.48$ 405 Eq. (3) $(R^2 = 0.75; p_{slope} < 0.005, p_{intercept} = 0.48);$ 406 407 408 $[HONO]_{ADS} = (1.08 \pm 0.19) [HONO]_{PTR-ToF} - 8.81$ Eq. (4) 409 $(R^2 = 0.87; p_{slope} < 0.005, p_{intercept} = 0.28).$ 410

411 We found significant linear correlation between each of the [HONO] techniques and 412 [HONO] $_{ADS}$ with a slope of ~1. Note that the y-intercepts of Eq. (1)—(4) are much 413 smaller than the overall range of measured [HONO] (up to 121 ppbv). In addition, p-414 values of the intercepts for all 4 fittings are much greater than 0.05, suggesting the 415 intercepts are not significantly different from zero. All data except one fall within 95% 416 prediction interval bounds of the overall fitting (Fig. 3). Therefore, we conclude that the 417 ADS method has high capture efficiency of HONO in the biomass combustion 418 environment, which assures the accuracy of the isotopic composition analysis and 419 applicability of this method for field-based biomass combustion research.

- 420
- 421 422

3.3 Isotopic composition of HONO and NO_x from burning different biomass

423 δ^{15} N of NO_x and HONO emitted from burning various biomass types in this study ranged 424 from -4.3 ‰ to +7.0‰ and -5.3 to +5.8‰, respectively (Table 1). There is no direct dependence of δ^{15} N on concentration of either HONO or NO_x (Figure S2). In Fig. 4, δ^{15} N 425 426 values of NO_x and HONO are shown for each biomass type. Each value represents a 427 concentration-weighted mean (if multiple samples were collected for a biomass type) 428 with error bars representing propagation of replicate variation and method precision. For 429 biomass types burned in replicate (ponderosa pine, lodgepole pine, Engelmann spruce, and Douglas-fir), the δ^{15} N-NO_x and δ^{15} N-HONO variation within a given biomass type is 430 smaller than the full range across all fuel types. Additionally, we note that the variations 431 of δ^{15} N-NO_x and δ^{15} N-HONO for ponderosa pine and δ^{15} N-HONO for Engelmann spruce 432 are larger than the method analytical precision of δ^{15} N-NO_x (1.5‰) and δ^{15} N-HONO 433 434 (0.5%), respectively, which represents fire-by-fire variation likely due to different 435 combustion conditions and/or different fuel compositions. For example, fuel moisture 436 content derived from the original biomass weight and dry biomass weight reveal that the 437 ponderosa pine burned in fire no.3 had more moisture content (48.1%) than fire no.2 438 (32.1%), which could affect combustion temperature and thus product formation. Fig. 4 439 also illustrates burning different biomass parts from specific vegetation can result in 440 fairly diverse δ^{15} N-HONO and δ^{15} N-NO_x, e.g. among ponderosa pine mixture, canopy 441 and litter, as well as between Engelmann spruce mixture and duff.

442

443 Our δ^{15} N-NO_x range falls well within the range (-7‰ to +12‰) found in the FLAME-4 444 experiment (Fibiger and Hastings, 2016). The FLAME-4 study investigated NO_x 445 emissions from burning a relatively large range of vegetation biomass from all over the 446 world, and found a linear relationship (Eq. (5)), indicating that 83% of the variation of δ^{15} N-NO_x is explained by δ^{15} N-biomass. The biomass types burned in this work focused 447 on vegetation in the western U.S., and differ greatly from that in FLAME-4, with 448 449 Ponderosa pine being the only common biomass between the two studies. Specifically, the δ^{15} N-biomass range (-4.2% to +0.9%) for this work is much narrower than that of the 450 FLAME-4 experiment (-8‰ to +8‰). 451

452

453
$$\delta^{15}$$
N-NO_x = 0.41 δ^{15} N-biomass + 1.0 (r²=0.83, p<0.001) Eq. (5)
454

To compare with the relationship found in Fibiger and Hastings (2016) we mass weighted the contributions from different components of the same biomass type. For the same type

of biomass, δ^{15} N-biomass varies amongst different parts of the vegetation with 457 458 differences as great as 4.1‰, 2.4‰, 4.6‰ and 2.6‰ for ponderosa pine, lodgepole pine, 459 Douglas-fir and Engelmann spruce, respectively (Table S1). In the FIREX experiments, 460 many of the burns were conducted for mixtures of various vegetation parts. For instance, one ponderosa pine fire contains canopy ($\sim 30\%$), litter ($\sim 28\%$), and other parts ($\sim 42\%$) 461 including duff and shrub, and the compositions vary slightly amongst each burn. 462 Therefore, the $\delta^{15}N$ of a particular biomass mixture is mass weighted according to its 463 composition contribution from each part (Table S1). Similarly, the δ^{15} N-NO_x and δ^{15} N-464 HONO from fires of different biomass parts are weighted by concentrations for each 465 466 biomass type, i.e. ponderosa pine (including mixture, canopy and litter) and Engelmann 467 spruce (including mixture and duff), to produce a signature associated with combustion 468 of that biomass type.

469

For purpose of comparison among different biomass types, we average δ^{15} N-NO_x (δ^{15} N-470 HONO) weighted by concentrations for each biomass type, i.e. ponderosa pine (including 471 472 mixture, canopy and litter) and Engelmann spruce (including mixture and duff) (all data are listed in Table S3). Linear regressions between δ^{15} N-HONO and δ^{15} N-biomass, as 473 well as that between δ^{15} N-NO_x and δ^{15} N-biomass, show that both δ^{15} N-HONO and δ^{15} N-474 NO_x increase with $\delta^{15}N$ -biomass in general (Fig. S3). However, the linear regressions 475 476 performed here are limited by small datasets (4 data points each) and unsurprisingly yield insignificant linear correlations for δ^{15} N-HONO (or δ^{15} N-NO_x) versus δ^{15} N-biomass (p 477 478 values are 0.1 and 0.5, respectively). Still, combining our results of δ^{15} N-NO_x versus 479 δ^{15} N-biomass from this work with those from the FLAME-4 study (Fibiger and Hastings, 480 2016) results in a significant linear correlation (Eq. (6)) and is shown in Fig. 5. Despite differences in burned biomass types between the two studies, our δ^{15} N-NO_x reasonably 481 482 overlap with the FLAME-4 results within our δ^{15} N-biomass range. The relationship between δ^{15} N-NO_x and δ^{15} N-biomass (Eq. (6)) for the combined data highly reproduces 483 484 that obtained solely from FLAME-4 study (Eq. (5)) and confirms the dependence of δ^{15} N-NO_x on δ^{15} N-biomass. 485

486

$$\delta^{15}$$
N-NO_x = (0.42±0.17) δ^{15} N-biomass + 1.3 (r²=0.71, p<0.001) Eq. (6)

The mean values weighted by concentration plotted in Fig. 4 show ¹⁵N of HONO is 489 consistently slightly more depleted than that of NO_x (δ^{15} N-HONO < δ^{15} N-NO_x) across all 490 491 the biomass types, except for ponderosa pine (litter) that results in an opposite relationship between δ^{15} N-HONO and δ^{15} N-NO_x. Furthermore, δ^{15} N-HONO is linearly 492 correlated with δ^{15} N-NO_x following a relationship of Eq. (7) within the δ^{15} N-NO_x and 493 494 δ^{15} N-HONO range obtained in the current study (Fig. 6). This provides potential insights 495 into HONO-NO_x interactions and HONO formation pathways in fresh emissions from 496 biomass burning. Although a number of studies on wildfire biomass burning have 497 suggested that partitioning of N emissions between NO_x and NH₃ depends on combustion 498 conditions represented by MCE (Ferek et al., 1998; Goode et al., 1999; McMeeking et al., 499 2009; Yokelson et al., 1996, 2008), HONO formation pathways remain unclear 500 (Alvarado et al., 2009, 2015; Nie et al., 2015).

501
$$\delta^{15}$$
N-HONO = 1.01 δ^{15} N-NO_x - 1.52 (R² = 0.89, p<0.001) Eq. (7)

502 Previous mechanistic studies on combustion of biomass/biofuel model compounds in a 503 well controlled closed system have investigated detailed nitrogen chemistry in the gas 504 phase, suggesting NO_x and HONO are formed from chain reactions involving oxidation 505 of precursors NH₃ and HCN, which are produced via devolatilization and pyrolysis of 506 amines and proteins in biomass/biofuel (Houshfar et al., 2012; Lucassen et al., 2011). 507 When the combustion conditions favor the oxidation of NH₃ and HCN, NO is first 508 formed and the chain reactions control the cycling of reactive nitrogen species (NO, NO₂ 509 and HONO). Detailed and mechanistic nitrogen chemistry for the chemical relationship 510 between NO_x and HONO in the combustion environment have been discussed in earlier 511 works (Chai and Goldsmith, 2017; Shrestha et al., 2018; Skreiberg et al., 2004). In 512 addition, Houshfar et al. (2012) performed biomass combustion kinetic modeling with 513 reduced mechanism via sensitivity analysis. From these works, we extract major 514 pathways (R1-R11) that are likely responsible for fast gas-phase inter-conversion 515 between NO_x and HONO within the combustion system. They found that whether HONO 516 is preferably converted from NO or NO₂ in series during nitrogen transformation 517 (referred to as nitrogen flow) critically depends on temperature. Specifically, within 1 518 second of residence time, at moderate temperatures (e.g. 700 °C), preferable nitrogen 519 flow following NO formation in biomass combustion is $NO \rightarrow NO_2 \rightarrow HONO \rightarrow NO$, and 520 major reactions involving NO_x-HONO conversion are listed in R1-R6; at high 521 temperatures (e.g. 850 °C and above), the nitrogen flow cycle NO \rightarrow HONO \rightarrow NO₂ \rightarrow NO 522 becomes preferable, and major reactions involving NO_x-HONO are R7-R11.

523
$$NO_2 + HNOH \rightarrow HONO + HNO$$
 R1

524
$$NO_2 + HNO \rightarrow HONO + NO$$
 R2

- 525 $NO_2 + HO_2 \rightarrow HONO + O_2$ R3
- 526 $NO_2 + H_2 \rightarrow HONO + H$ R4
- 527 NO₂ + C_xH_y (hydrocarbon) \rightarrow HONO + C_xH_{y-1} R5
- 528 HONO \rightarrow OH + NO R6
- 529 $OH + NO \rightarrow HONO$ R7
- 530 HONO + $NH_2 \rightarrow NO_2 + NH_3$ R8
- 531 HONO + NH \rightarrow NO₂ + NH₂ R9
- 532 HONO + O \rightarrow NO₂ + OH R10
- 533 HONO + OH \rightarrow NO₂ + H₂O R11

Although our studied fuels are more complicated in composition than a model system involving no more than a few starting species, results from the above studies provide fundamental underpinnings for biomass combustion. Also note that heterogeneous chemistry after these species were emitted was not considered here as the residence time 538 of the fresh plume in our study was \sim 5 seconds, which is of the same magnitude as that predicted in the nitrogen flow analysis (Houshfar et al., 2012). Kinetic isotope effects 539 540 (KIE) of these reactions have not been characterized; so only a semi-quantitative 541 prediction is presented here. At low temperatures, R1-R5 are all H-abstraction reactions 542 involving loose transition states that have significant activation energy; a primary KIE is expected for such conditions and leads to ¹⁵N depletion in the product (HONO) (Chai et 543 544 al., 2014; Matsson and Westaway, 1999, and references therein). Additionally, R6 is a 545 unimolecular dissociation reaction with no reaction barrier, and hence R6 could be expected to have a small kinetic isotope effect enriching ¹⁵N in HONO, somewhat 546 547 offsetting the depletion that arose from R1-R5. Consequently, the overall isotope effect of R1-R6 would lead to δ^{15} N-HONO < δ^{15} N-NO_x by a small difference, consistent with our 548 549 results (Fig. 4). On the other hand, the KIE for the reactions R7-R11 at higher 550 temperatures (> 850 °C) is expected to enrich ¹⁵N in HONO relative to NO_x (Chai and 551 Dibble, 2014), leading to an opposite isotope effect to that predicted at lower 552 temperatures.

Temperatures of the biomass combustion process span a large range involving different processes including preheating, drying, distillation, pyrolysis, gasification (aka "glowing combustion") and oxidation in turbulent diffusion flames at a range of temperatures associated with changing flame dynamics (Yokelson et al., 1996). Despite this complexity, our measured slight ¹⁵N enrichment in NO_x compared to HONO (Table 1, Fig. 4) suggests that the reactions R1-R6 played a more important role than R7-R11 in HONO formation during the FIREX Fire Lab experiments.

560

561 **3.4 Isotopic composition of nitrates collected on particle filters**

562 All Nylasorb filter extract solutions showed no detectable NO₃⁻ and NO₂⁻ concentrations, 563 indicating no significant amount of HNO₃ was collected on these filters, which is 564 consistent with the very low concentrations measured by MC/IC (note that low 565 concentration and limited sample volume also preclude further isotopic analysis of HNO₃ 566 collected by MC/IC). By contrast, we found 5 out of 20 particulate filter extract solutions 567 had detectable NO_3^- concentration that were sufficient (10 nmol N) for isotopic composition analysis (Table 1). $\delta^{15}N$ and $\delta^{18}O$ reported here are considered to represent 568 $NO_3(p)$. $\delta^{15}N-NO_3(p)$ of the five samples (burns) range from -10.6 to -7.4 %, all of 569 which are more ^{15}N depleted than that of HONO and NO_x. In addition, the smaller range 570 of δ^{15} N-NO₃⁻ than that of δ^{15} N-HONO and δ^{15} N-NO_x rules out possible transformation of 571 572 NO_x and HONO to nitrate on the filters, which could distort the isotopic composition of 573 NO_x and HONO.

574

575 In the FLAME-4 experiments, only one particulate filter had captured $NO_3^{-}(p)$ above the 576 concentration detection limit, whereas HNO₃ collected on Nylasorb filters from 7 577 experiments were above the concentration detection limit and therefore only δ^{15} N-HNO₃ 578 (-0.3‰ to 11.2‰) were reported (Fibiger and Hastings, 2016). The contrast with our 579 filter results are likely attributed to different formation mechanisms under different 580 conditions, in addition to variation of fuel types. Of the 7 detectable HNO₃ collections 581 from FLAME-4, 5 represented room burns for which samples were collected from smoke 582 aged for 1-2 hours in the lab, and the sampled HNO₃ was likely a secondary product. By contrast all our observed NO₃⁻(p) were in fresh emissions and may have been derived from plant nitrate (Cárdenas-Navarro et al., 1999) and/or combustion reactions. There have been no other studies on δ^{15} N of NO₃⁻(p) and HNO₃ directly emitted from fresh plumes to the best of our knowledge, so more investigation using both laboratory work (isotope effect) and kinetic modeling will be needed in order to understand formation mechanisms of HNO₃ and NO₃⁻(p) in the biomass combustion process and their respective isotope effects.

590

In addition to $\delta^{15}N$, we report $\delta^{18}O$ of HONO and NO₃(p) directly emitted from biomass 591 burning plumes with ranges of 5.2% to 15.2% and 11.5% to 14.8%, respectively. These 592 are the first observations reported for δ^{18} O of reactive nitrogen species directly emitted 593 from biomass burning and low values are expected for the δ^{18} O, which, in this case, is 594 595 mainly extracted from that of molecular oxygen ($\delta^{18}O = -23.5\%$) (Kroopnick and Craig, 1972), biomass/cellulose ($\delta^{18}O = 15\%-35\%$), and/or biomass contained water ($\delta^{18}O =$ 596 597 ~ 0 %– 16%) (Keel et al., 2016). In field studies where photochemistry and O₃ are 598 inevitably involved in the reactive nitrogen cycle in various stages of aged plumes, we expect to see much more elevated δ^{18} O values of HONO and NO₃(p) due to the 599 extremely high value of δ^{18} O-O₃ (~110‰) (Vicars and Savarino, 2014). Therefore, the 600 601 δ^{18} O found in the lab is helpful in understanding conditions where photochemistry would 602 not apply (e.g. nighttime fresh smoke) and should be distinguishable from the expected 603 higher δ^{18} O that would be found in aged smoke and/or daytime fresh smoke.

604

605 4 Conclusions

606

In this study we applied new methods for characterizing the isotopic composition of 607 reactive nitrogen species including NO_x (δ^{15} N), HONO (δ^{15} N and δ^{18} O), and NO₃(p) 608 $(\delta^{15}N \text{ and } \delta^{18}O)$ emitted directly from biomass burning. We measured fresh (stack) 609 610 emissions from 20 laboratory fires of different fuels during the 2016 FIREX Fire Lab 611 experiments. NO_x, HONO and HNO₃ emitted in fresh smoke reached their peak in most 612 of our fires within five minutes of ignition of biomass (i.e. when flaming combustion 613 peaked). The HONO mixing ratio was typically ~ 2 orders of magnitude larger than 614 HNO₃, and HONO/NO_x ratio ranged from 0.13 to 0.53.

615

616 Our HONO collection method (ADS) for isotopic analysis was applied to biomass 617 burning (BB) for the first time. The good agreement for concentration comparison 618 between our method and 4 high time-resolution HONO concentration methods suggests 619 high collection efficiency of HONO from BB emissions, which ensures accurate isotopic 620 compositional analysis. Comparison with concurrent observations and a previous study 621 show that the combination of our HONO and NO_x collection methods are compatible, 622 allowing for simultaneous determination of the isotopic composition of both HONO and 623 NO_x. This provides important potential for investigating the photochemical and non-624 photochemical relationships between HONO and NO_x in a variety of environments, and 625 especially in BB plumes.

626

 δ^{15} N-NO_x emitted from burning various Western U.S. biomass types in this study ranged from -4.3 ‰ to +7.0 ‰, falling well within the range found by Fibiger and Hastings 629 (2016), although the vegetation types were much broader in the earlier study. We report the first δ^{15} N-HONO emitted directly from burning, ranging from -5.3% to +5.8%. δ^{15} N-630 NO_x and δ^{15} N-HONO range derived from BB can be further compared with that from 631 632 other sources using the same methods presented here, and provide insights into source 633 signatures for both NO_x and HONO. This study also showed the important capability of determining δ^{18} O-HONO and δ^{18} O-NO₃ (p) from BB plumes, and we expect δ^{18} O of both 634 HONO and $NO_3(p)$ produced under photochemical conditions will be much higher than 635 636 the lab results due to the important role of O₃ in reactive nitrogen oxidation.

637

Interestingly, the linear correlation between δ^{15} N-HONO and δ^{15} N-NO_x for the biomass 638 639 we studied suggests systematic co-production of NO_x and HONO occurs during biomass 640 combustion and both of them are released as primary pollutants in fresh smoke. The relationship between δ^{15} N-HONO and δ^{15} N-NO_x likely reflects that HONO was produced 641 to a larger extent at moderate combustion temperatures (< ~800 °C) than higher 642 643 temperatures on the basis of a simplified mechanism for flow of reactive nitrogen species. 644 However, we note that this relationship is derived from all measured δ^{15} N-HONO and 645 δ^{15} N-NO_x in fires ranging from smoldering to flaming, so is not necessarily 646 representative of a particular combustion condition. Still, it is likely that a compilation 647 over a range of conditions is more useful for potentially distinguishing HONO sources 648 and formation pathways in the environment since it will always be a challenge to assess 649 exact combustion temperatures. Determining these relationships in real wildfire smoke 650 will be essential for better constraint on NO_x and HONO budgets, and eventually may 651 improve ozone and secondary aerosol predictions for regional air quality.

652

653 Data availability. The data from the laboratory tests are available on request from the

654 corresponding authors. Data from the 2016 Missoula Fire lab are available here:

- 655 https://esrl.noaa.gov/csd/groups/csd7/measurements/2016firex/FireLab/DataDownload/
- 656 Supplement.

657 Author contribution. JC, MH and JD designed this work. JC and DJM conducted the

sample collections at the Fire Lab, with additional support from MH, JD and ES. JC

659 carried out the isotopic composition measurements; DJM supported the isotopic research

and interpretation. ES helped analyze the MC/IC data. VS and RY provided the OP-FTIR

- data. KJZ and SSB provided the ACES data. ARK and CW provided the PTR-ToF data.
- G62 JC wrote the manuscript, and all authors provided edits and feedback.
- 663 *Competing interests.* The authors declare that they have no conflicts of interest.
- 664 Acknowledgement. This work was supported by funding from the National Oceanic and
- Atmospheric Administration (AC4 Award NA16OAR4310098 to MH) and
- the National Science Foundation (AGS-1351932 to MH). The FIREX Fire Lab study was

supported in part by the NOAA Climate Office's Atmospheric Chemistry, Carbon Cycle,

and Climate program. We are grateful for Ruby Ho for laboratory support and Marshall

669 Otter for the biomass δ^{15} N analysis. We also thank James Roberts and Matthew Coggon

670 for helpful discussions. We are thankful for the two anonymous reviewers for their

- 671 helpful comments.
- 672
- 673

674 References675

- 676 Akagi, S. K., Craven, J. S., Taylor, J. W., McMeeking, G. R., Yokelson, R. J., Burling, I.
- R., Urbanski, S. P., Wold, C. E., Seinfeld, J. H., Coe, H., Alvarado, M. J. and Weise, D.
- 678 R.: Evolution of trace gases and particles emitted by a chaparral fire in California, Atmos
- 679 Chem Phys, 12(3), 1397–1421, doi:10.5194/acp-12-1397-2012, 2012.
- 680 Akagi, S. K., Yokelson, R. J., Burling, I. R., Meinardi, S., Simpson, I., Blake, D. R.,
- 681 McMeeking, G. R., Sullivan, A., Lee, T., Kreidenweis, S., Urbanski, S., Reardon, J.,
- 682 Griffith, D. W. T., Johnson, T. J. and Weise, D. R.: Measurements of reactive trace gases 683 and variable O3 formation rates in some South Carolina biomass burning plumes, Atmos
- 684 Chem Phys, 13(3), 1141–1165, doi:10.5194/acp-13-1141-2013, 2013.
- Alvarado, M. J. and Prinn, R. G.: Formation of ozone and growth of aerosols in young
 smoke plumes from biomass burning: 1. Lagrangian parcel studies, J. Geophys. Res.
 Atmospheres 114(D0). doi:10.1020/2008/D0111144_2000
- 687 Atmospheres, 114(D9), doi:10.1029/2008JD011144, 2009.
- Alvarado, M. J., Wang, C. and Prinn, R. G.: Formation of ozone and growth of aerosols
 in young smoke plumes from biomass burning: 2. Three-dimensional Eulerian studies, J.
- 690 Geophys. Res. Atmospheres, 114(D9), doi:10.1029/2008JD011186, 2009.
- Alvarado, M. J., Lonsdale, C. R., Yokelson, R. J., Akagi, S. K., Coe, H., Craven, J. S.,
- Fischer, E. V., McMeeking, G. R., Seinfeld, J. H., Soni, T., Taylor, J. W., Weise, D. R.
- and Wold, C. E.: Investigating the links between ozone and organic aerosol chemistry in
- a biomass burning plume from a prescribed fire in California chaparral, Atmos Chem
- 695 Phys, 15(12), 6667–6688, doi:10.5194/acp-15-6667-2015, 2015.
- Böhlke, J. K., Mroczkowski, S. J. and Coplen, T. B.: Oxygen isotopes in nitrate: new
- reference materials for 180:170:160 measurements and observations on nitrate-water
- equilibration, Rapid Commun. Mass Spectrom., 17(16), 1835–1846,
- 699 doi:10.1002/rcm.1123, 2003.
- 700 Burling, I. R., Yokelson, R. J., Griffith, D. W. T., Johnson, T. J., Veres, P., Roberts, J.
- 701 M., Warneke, C., Urbanski, S. P., Reardon, J., Weise, D. R., Hao, W. M. and de Gouw,
- J.: Laboratory measurements of trace gas emissions from biomass burning of fuel types
- from the southeastern and southwestern United States, Atmos Chem Phys, 10(22),
- 704 11115–11130, doi:10.5194/acp-10-11115-2010, 2010.
- 705 Burling, I. R., Yokelson, R. J., Akagi, S. K., Urbanski, S. P., Wold, C. E., Griffith, D. W.
- T., Johnson, T. J., Reardon, J. and Weise, D. R.: Airborne and ground-based
- measurements of the trace gases and particles emitted by prescribed fires in the United
- 708 States, Atmos Chem Phys, 11(23), 12197–12216, doi:10.5194/acp-11-12197-2011, 2011.

- Cárdenas-Navarro, R., Adamowicz, S. and Robin, P.: Nitrate accumulation in plants: a role for water, J. Exp. Bot., 50(334), 613–624, doi:10.1093/jxb/50.334.613, 1999.
- 711 Casciotti, K. L., Sigman, D. M., Hastings, M. G., Böhlke, J. K. and Hilkert, A.:
- 712 Measurement of the Oxygen Isotopic Composition of Nitrate in Seawater and Freshwater
- 713 Using the Denitrifier Method, Anal. Chem., 74(19), 4905–4912, doi:10.1021/ac020113w,
- 714 2002.
- 715 Casciotti, K. L., Böhlke, J. K., McIlvin, M. R., Mroczkowski, S. J. and Hannon, J. E.:
- 716 Oxygen Isotopes in Nitrite: Analysis, Calibration, and Equilibration, Anal. Chem., 79(6),
- 717 2427–2436, doi:10.1021/ac061598h, 2007.
- 718 Chai, J. and Dibble, T. S.: Pressure Dependence and Kinetic Isotope Effects in the
- Absolute Rate Constant for Methoxy Radical Reacting with NO2, Int. J. Chem. Kinet.,
 46(9), 501–511, doi:10.1002/kin.20865, 2014.
- 721 Chai, J. and Goldsmith, C. F.: Rate coefficients for fuel + NO2: Predictive kinetics for
- HONO and HNO2 formation, Proc. Combust. Inst., 36(1), 617–626,
- doi:10.1016/j.proci.2016.06.133, 2017.
- Chai, J. and Hastings, M.: Collection Method for Isotopic Analysis of Gaseous Nitrous
 Acid, Anal. Chem., 90(1), 830–838, doi:10.1021/acs.analchem.7b03561, 2018.
- 726 Chai, J., Hu, H., Dibble, T. S., Tyndall, G. S. and Orlando, J. J.: Rate Constants and
- 727 Kinetic Isotope Effects for Methoxy Radical Reacting with NO2 and O2, J. Phys. Chem.
- 728 A, 118(20), 3552–3563 [online] Available from:
- 729 http://pubs.acs.org/doi/abs/10.1021/jp501205d (Accessed 24 June 2015), 2014.
- 730 Christian, T. J., Kleiss, B., Yokelson, R. J., Holzinger, R., Crutzen, P. J., Hao, W. M.,
- 731 Shirai, T. and Blake, D. R.: Comprehensive laboratory measurements of biomass-burning
- emissions: 2. First intercomparison of open-path FTIR, PTR-MS, and GC-MS/FID/ECD,
- 733 J. Geophys. Res. Atmospheres, 109(D2), doi:10.1029/2003JD003874, 2004.
- 734 Cook, P. A., Savage, N. H., Turquety, S., Carver, G. D., O'Connor, F. M., Heckel, A.,
- 735 Stewart, D., Whalley, L. K., Parker, A. E., Schlager, H., Singh, H. B., Avery, M. A.,
- 736 Sachse, G. W., Brune, W., Richter, A., Burrows, J. P., Purvis, R., Lewis, A. C., Reeves,
- 737 C. E., Monks, P. S., Levine, J. G. and Pyle, J. A.: Forest fire plumes over the North
- 738 Atlantic: p-TOMCAT model simulations with aircraft and satellite measurements from
- the ITOP/ICARTT campaign, J. Geophys. Res. Atmospheres, 112, D10S43,
- 740 doi:10.1029/2006JD007563, 2007.
- 741 Crutzen, P. J. and Andreae, M. O.: Biomass burning in the tropics: impact on atmospheric
- chemistry and biogeochemical cycles, Science, 250(4988), 1669–1678,
- 743 doi:10.1126/science.250.4988.1669, 1990.
- 744 Dibb, J. E., Arsenault, M., Peterson, M. C. and Honrath, R. E.: Fast nitrogen oxide
- photochemistry in Summit, Greenland snow, Atmos. Environ., 36(15–16), 2501–2511,
- 746 doi:10.1016/S1352-2310(02)00130-9, 2002.

- 747 Ferek, R., Reid, J., Hobbs, P., Blake, D. and Liousse, C.: Emission factors of
- hydrocarbons, halocarbons, trace gases and particles from biomass burning in Brazil, J.
- 749 Geophys. Res.-ATMOSPHERES, 103, 32107–32118, doi:10.1029/98JD00692, 1998.
- Fibiger, D. L. and Hastings, M. G.: First Measurements of the Nitrogen Isotopic
- 751 Composition of NOx from Biomass Burning, Environ. Sci. Technol., 50(21), 11569–
- 752 11574, doi:10.1021/acs.est.6b03510, 2016.
- Fibiger, D. L., Hastings, M. G., Lew, A. F. and Peltier, R. E.: Collection of NO and NO2
- for Isotopic Analysis of NOx Emissions, Anal. Chem., 86(24), 12115–12121,
- 755 doi:10.1021/ac502968e, 2014.
- FIREX: FIREX 2016 Fire Lab Data Archive, [online] Available from:
- 757 https://esrl.noaa.gov/csd/groups/csd7/measurements/2016firex/FireLab/DataDownload/in
- dex.php?page=/csd/groups/csd7/measurements/2016firex/FireLab/DataDownload/
- 759 (Accessed 25 March 2019), 2016.
- Goode, J. G., Yokelson, R. J., Susott, R. A. and Ward, D. E.: Trace gas emissions from
- laboratory biomass fires measured by open-path Fourier transform infrared spectroscopy:
 Fires in grass and surface fuels, J. Geophys. Res., 104, 21, doi:10.1029/1999JD900360,
- 763 1999.
- Hastings, M. G., Casciotti, K. L. and Elliott, E. M.: Stable Isotopes as Tracers of
- Anthropogenic Nitrogen Sources, Deposition, and Impacts, Elements, 9(5), 339–344,
 doi:10.2113/gselements.9.5.339, 2013.
- Houshfar, E., Skreiberg, Ø., Glarborg, P. and Løvås, T.: Reduced chemical kinetic mechanisms for NO_x emission prediction in biomass combustion, Int. J. Chem. Kinet., 4(44), 219–231, doi:10.1002/kin.20716, 2012.
- Jaeglé, L., Steinberger, L., Martin, R. V. and Chance, K.: Global partitioning of NOx
 sources using satellite observations: relative roles of fossil fuel combustion, biomass
 burning and soil emissions, Faraday Discuss., 130, 407–423; discussion 491-517, 519–
 524, 2005.
- Jaffe, D. and Briggs, N.: Ozone production from wildfires: A critical review, Atmos.
 Environ., 51, 1–10, doi:10.1016/j.atmosenv.2011.11.063, 2012.
- Keel, S. G., Joos, F., Spahni, R., Saurer, M., Weigt, R. B. and Klesse, S.: Simulating
- 777 oxygen isotope ratios in tree ring cellulose using a dynamic global vegetation model,
- 778 Biogeosciences, 13(13), 3869–3886, doi:https://doi.org/10.5194/bg-13-3869-2016, 2016.
- 779 Keene, W. C., Lobert, J. M., Crutzen, P. J., Maben, J. R., Scharffe, D. H., Landmann, T.,
- 780 Hély, C. and Brain, C.: Emissions of major gaseous and particulate species during
- 781 experimental burns of southern African biomass, J. Geophys. Res. Atmospheres,
- 782 111(D4), doi:10.1029/2005JD006319, 2006.

- 783 Koss, A. R., Sekimoto, K., Gilman, J. B., Selimovic, V., Coggon, M. M., Zarzana, K. J.,
- Yuan, B., Lerner, B. M., Brown, S. S., Jimenez, J. L., Krechmer, J., Roberts, J. M.,
- 785 Warneke, C., Yokelson, R. J. and Gouw, J. de: Non-methane organic gas emissions from
- 786 biomass burning: identification, quantification, and emission factors from PTR-ToF
- during the FIREX 2016 laboratory experiment, Atmospheric Chem. Phys., 18(5), 3299-
- 788 3319, doi:https://doi.org/10.5194/acp-18-3299-2018, 2018.
- 789 Kraus, S.: DOASIS: a framework design for DOAS., 2006.
- Kroopnick, P. and Craig, H.: Atmospheric oxygen: isotopic composition and solubility
 fractionation, Science, 175(4017), 54–55, doi:10.1126/science.175.4017.54, 1972.
- Lapina, K., Honrath R. E., Owen R. C., Val Martín M., Hyer E. J. and Fialho P.: Late
 summer changes in burning conditions in the boreal regions and their implications for
 NOx and CO emissions from boreal fires, J. Geophys. Res. Atmospheres, 113(D11),
- 795 doi:10.1029/2007JD009421, 2008.
- Liu, X., Zhang, Y., Huey, L. G., Yokelson, R. J., Wang, Y., Jimenez, J. L., Campuzano-
- Jost, P., Beyersdorf, A. J., Blake, D. R., Choi, Y., Clair, J. M. S., Crounse, J. D., Day, D.
- A., Diskin, G. S., Fried, A., Hall, S. R., Hanisco, T. F., King, L. E., Meinardi, S.,
- 799 Mikoviny, T., Palm, B. B., Peischl, J., Perring, A. E., Pollack, I. B., Ryerson, T. B.,
- 800 Sachse, G., Schwarz, J. P., Simpson, I. J., Tanner, D. J., Thornhill, K. L., Ullmann, K.,
- 801 Weber, R. J., Wennberg, P. O., Wisthaler, A., Wolfe, G. M. and Ziemba, L. D.:
- 802 Agricultural fires in the southeastern U.S. during SEAC4RS: Emissions of trace gases
- and particles and evolution of ozone, reactive nitrogen, and organic aerosol, J. Geophys.
- 804 Res. Atmospheres, 121(12), 7383–7414, doi:10.1002/2016JD025040, 2016.
- Lucassen, A., Labbe, N., Westmoreland, P. R. and Kohse-Höinghaus, K.: Combustion
- 806 chemistry and fuel-nitrogen conversion in a laminar premixed flame of morpholine as a
- 807 model biofuel, Combust. Flame, 158(9), 1647–1666,
- 808 doi:10.1016/j.combustflame.2011.02.010, 2011.
- 809 Matsson, O. and Westaway, K. C.: Secondary Deuterium Kinetic Isotope Effects and
- 810 Transition State Structure, in Advances in Physical Organic Chemistry, vol. 31, edited by
- 811 D. Bethell, pp. 143–248, Academic Press., 1999.
- 812 McMeeking, G. R., Kreidenweis, S. M., Baker, S., Carrico, C. M., Chow, J. C., Collett, J.
- 813 L., Hao, W. M., Holden, A. S., Kirchstetter, T. W., Malm, W. C., Moosmüller, H.,
- 814 Sullivan, A. P. and Wold, C. E.: Emissions of trace gases and aerosols during the open
- combustion of biomass in the laboratory, J. Geophys. Res. Atmospheres, 114(D19),
- 816 D19210, doi:10.1029/2009JD011836, 2009.
- 817 Miller, D. J., Wojtal, P. K., Clark, S. C. and Hastings, M. G.: Vehicle NOx emission
- 818 plume isotopic signatures: Spatial variability across the eastern United States, J. Geophys.
- 819 Res. Atmospheres, 122(8), 4698–4717, doi:10.1002/2016JD025877, 2017.

- 820 Miller, D. J., Chai, J., Guo, F., Dell, C. J., Karsten, H. and Hastings, M. G.: Isotopic
- 821 Composition of In Situ Soil NOx Emissions in Manure-Fertilized Cropland, Geophys.
- 822 Res. Lett., 45(21), 12,058-12,066, doi:10.1029/2018GL079619, 2018.
- 823 Min, K.-E., Washenfelder, R. A., Dubé, W. P., Langford, A. O., Edwards, P. M.,
- 824 Zarzana, K. J., Stutz, J., Lu, K., Rohrer, F., Zhang, Y. and Brown, S. S.: A broadband
- 825 cavity enhanced absorption spectrometer for aircraft measurements of glyoxal,
- 826 methylglyoxal, nitrous acid, nitrogen dioxide, and water vapor, Atmospheric Meas.
- 827 Tech., 9(2), 423–440, doi:https://doi.org/10.5194/amt-9-423-2016, 2016.
- 828 Neuman, J. A., Trainer, M., Brown, S. S., Min, K.-E., Nowak, J. B., Parrish, D. D.,
- 829 Peischl, J., Pollack, I. B., Roberts, J. M., Ryerson, T. B. and Veres, P. R.: HONO
- 830 emission and production determined from airborne measurements over the Southeast
- 831 U.S., J. Geophys. Res. Atmospheres, 121(15), 9237–9250, doi:10.1002/2016JD025197, 832 2016.
- 833 Nie, W., Ding, A. J., Xie, Y. N., Xu, Z., Mao, H., Kerminen, V.-M., Zheng, L. F., Qi, X.
- 834 M., Huang, X., Yang, X.-Q., Sun, J. N., Herrmann, E., Petäjä, T., Kulmala, M. and Fu, C.
- 835 B.: Influence of biomass burning plumes on HONO chemistry in eastern China, Atmos
- 836 Chem Phys, 15(3), 1147–1159, doi:10.5194/acp-15-1147-2015, 2015.
- 837 Perrino, C., De Santis, F. and Febo, A.: Criteria for the choice of a denuder sampling
- 838 technique devoted to the measurement of atmospheric nirous and nitric acids,
- 839 Atmospheric Environ. Part Gen. Top., 24(3), 617-626, doi:10.1016/0960-
- 840 1686(90)90017-H, 1990.
- 841 Pinto, J. P., Dibb, J., Lee, B. H., Rappenglück, B., Wood, E. C., Levy, M., Zhang, R.-Y.,
- Lefer, B., Ren, X.-R., Stutz, J., Tsai, C., Ackermann, L., Golovko, J., Herndon, S. C., 842
- 843 Oakes, M., Meng, Q.-Y., Munger, J. W., Zahniser, M. and Zheng, J.: Intercomparison of
- 844 field measurements of nitrous acid (HONO) during the SHARP campaign, J. Geophys.
- 845 Res. Atmospheres, 119(9), 5583-5601, doi:10.1002/2013JD020287, 2014.
- 846 Reinhardt, E. D., Keane, R. E. and Brown, J. K.: First Order Fire Effects Model: FOFEM
- 847 4.0, user's guide, Gen Tech Rep INT-GTR-344 Ogden UT US Dep. Agric. For. Serv. 848 Intermt. Res. Stn. 65 P, 344, doi:10.2737/INT-GTR-344, 1997.
- 849 Roberts, J. M., Veres, P., Warneke, C., Neuman, J. A., Washenfelder, R. A., Brown, S.
- 850 S., Baasandorj, M., Burkholder, J. B., Burling, I. R., Johnson, T. J., Yokelson, R. J. and
- 851 de Gouw, J.: Measurement of HONO, HNCO, and other inorganic acids by negative-ion
- 852 proton-transfer chemical-ionization mass spectrometry (NI-PT-CIMS): application to
- 853 biomass burning emissions, Atmos Meas Tech, 3(4), 981-990, doi:10.5194/amt-3-981-
- 854 2010, 2010.
- 855 Scheuer, E., Talbot, R. W., Dibb, J. E., Seid, G. K., DeBell, L. and Lefer, B.: Seasonal
- 856 distributions of fine aerosol sulfate in the North American Arctic basin during TOPSE, J. 857 Geophys. Res. Atmospheres, 108(D4), doi:10.1029/2001JD001364, 2003.

- 858 Selimovic, V., Yokelson, R. J., Warneke, C., Roberts, J. M., de Gouw, J., Reardon, J. and
- 859 Griffith, D. W. T.: Aerosol optical properties and trace gas emissions by PAX and OP-
- 860 FTIR for laboratory-simulated western US wildfires during FIREX, Atmos Chem Phys,
- 861 18(4), 2929–2948, doi:10.5194/acp-18-2929-2018, 2018.
- Shrestha, K. P., Seidel, L., Zeuch, T. and Mauss, F.: Detailed Kinetic Mechanism for the
 Oxidation of Ammonia Including the Formation and Reduction of Nitrogen Oxides,
- 864 Energy Fuels, doi:10.1021/acs.energyfuels.8b01056, 2018.
- 865 Sigman, D. M., Casciotti, K. L., Andreani, M., Barford, C., Galanter, M. and Böhlke, J.
- 866 K.: A Bacterial Method for the Nitrogen Isotopic Analysis of Nitrate in Seawater and
- 867 Freshwater, Anal. Chem., 73(17), 4145–4153, doi:10.1021/ac010088e, 2001.
- 868 Skreiberg, Ø., Kilpinen, P. and Glarborg, P.: Ammonia chemistry below 1400 K under 869 fuel-rich conditions in a flow reactor, Combust. Flame, 136(4), 501–518,
- doi:10.1016/j.combustflame.2003.12.008, 2004.
- 871 Stockwell, C. E., Yokelson, R. J., Kreidenweis, S. M., Robinson, A. L., DeMott, P. J.,
- 872 Sullivan, R. C., Reardon, J., Ryan, K. C., Griffith, D. W. T. and Stevens, L.: Trace gas
- 873 emissions from combustion of peat, crop residue, domestic biofuels, grasses, and other
- fuels: configuration and Fourier transform infrared (FTIR) component of the fourth Fire
 Lab at Missoula Experiment (FLAME-4), Atmos Chem Phys, 14(18), 9727–9754,
- 875 Lab at Missoura Experiment (FLAME-4), Atmos Chem Phys, 14(18 976 doi:10.5104/acm.14.0727.2014.2014
- doi:10.5194/acp-14-9727-2014, 2014.
- Stutz, J., Oh, H.-J., Whitlow, S. I., Anderson, C., Dibb, J. E., Flynn, J. H., Rappenglück,
 B. and Lefer, B.: Simultaneous DOAS and mist-chamber IC measurements of HONO in
 Houston, TX, Atmos. Environ., 44(33), 4090–4098, doi:10.1016/j.atmosenv.2009.02.003,
- 880 2010.
- 881 Tkacik, D. S., Robinson, E. S., Ahern, A., Saleh, R., Stockwell, C., Veres, P., Simpson, I.
- J., Meinardi, S., Blake, D. R., Yokelson, R. J., Presto, A. A., Sullivan, R. C., Donahue, N.
- 883 M. and Robinson, A. L.: A dual-chamber method for quantifying the effects of
- atmospheric perturbations on secondary organic aerosol formation from biomass burning
- 885 emissions, J. Geophys. Res. Atmospheres, 122(11), 6043–6058,
- doi:10.1002/2016JD025784, 2017.
- 887 Travis, K. R., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Zhu, L., Yu, K.,
- 888 Miller, C. C., Yantosca, R. M., Sulprizio, M. P., Thompson, A. M., Wennberg, P. O.,
- 889 Crounse, J. D., St. Clair, J. M., Cohen, R. C., Laughner, J. L., Dibb, J. E., Hall, S. R.,
- Ullmann, K., Wolfe, G. M., Pollack, I. B., Peischl, J., Neuman, J. A. and Zhou, X.: Why
- 891 do models overestimate surface ozone in the Southeast United States?, Atmos Chem
- 892 Phys, 16(21), 13561–13577, doi:10.5194/acp-16-13561-2016, 2016.
- 893 Trentmann, J., Yokelson, R. J., Hobbs, P. V., Winterrath, T., Christian, T. J., Andreae, M.
- 0. and Mason, S. A.: An analysis of the chemical processes in the smoke plume from a
- 895 savanna fire, J. Geophys. Res. Atmospheres, 110(D12), doi:10.1029/2004JD005628, 2005
- 896 2005.

- 897 Val Martín, M., Honrath, R. E., Owen, R. C., Pfister, G., Fialho, P. and Barata, F.:
- 898 Significant enhancements of nitrogen oxides, black carbon, and ozone in the North
- 899 Atlantic lower free troposphere resulting from North American boreal wildfires, J.
- 900 Geophys. Res. Atmospheres, 111(D23), doi:10.1029/2006JD007530, 2006.
- 901 Veres, P., Roberts, J. M., Burling, I. R., Warneke, C., Gouw, J. de and Yokelson, R. J.:
- 902 Measurements of gas-phase inorganic and organic acids from biomass fires by negative-
- 903 ion proton-transfer chemical-ionization mass spectrometry, J. Geophys. Res.
- 904 Atmospheres, 115(D23), doi:10.1029/2010JD014033, 2010.
- 905 Vicars, W. C. and Savarino, J.: Quantitative constraints on the 17O-excess (Δ17O)
- signature of surface ozone: Ambient measurements from 50°N to 50°S using the nitrite-
- 907 coated filter technique, Geochim. Cosmochim. Acta, 135(Supplement C), 270–287,
 908 doi:10.1016/j.gca.2014.03.023, 2014.
- 909 Wojtal, P. K., Miller, D. J., O'Conner, M., Clark, S. C. and Hastings, M. G.: Automated,
- High-resolution Mobile Collection System for the Nitrogen Isotopic Analysis of NOx, J.
 Vis. Exp. JoVE, (118), doi:10.3791/54962, 2016.
- Ye, C., Zhou, X., Pu, D., Stutz, J., Festa, J., Spolaor, M., Tsai, C., Cantrell, C., Mauldin,
 R. L., Campos, T., Weinheimer, A., Hornbrook, R. S., Apel, E. C., Guenther, A., Kaser,
 L., Yuan, B., Karl, T., Haggerty, J., Hall, S., Ullmann, K., Smith, J. N., Ortega, J. and
- 915 Knote, C.: Rapid cycling of reactive nitrogen in the marine boundary layer, Nature,
- 916 532(7600), 489–491, doi:10.1038/nature17195, 2016.
- 917 Yokelson, R., W. T. Griffith, D. and Ward, D.: Open-path Fourier transform infrared
 918 studies of large-scale laboratory biomass fires, J. Geophys. Res., 101, 21067,
- 919 doi:10.1029/96JD01800, 1996.
- 920 Yokelson, R. J., Karl, T., Artaxo, P., Blake, D. R., Christian, T. J., Griffith, D. W. T.,
 921 Guenther, A. and Hao, W. M.: The Tropical Forest and Fire Emissions Experiment:
- 921 Ouenther, A. and Hao, W. M.. The Hopical Forest and File Emissions Experiment. 922 overview and airborne fire emission factor measurements, Atmos Chem Phys, 7(19),
- 923 5175–5196, doi:10.5194/acp-7-5175-2007, 2007.
- Yokelson, R. J., Christian, T. J., Karl, T. G. and Guenther, A.: The tropical forest and fire
 emissions experiment: laboratory fire measurements and synthesis of campaign data,
 Atmos Chem Phys, 8(13), 3509–3527, doi:10.5194/acp-8-3509-2008, 2008.
- 927 Yokelson, R. J., Crounse, J. D., DeCarlo, P. F., Karl, T., Urbanski, S., Atlas, E., Campos,
- 928 T., Shinozuka, Y., Kapustin, V., Clarke, A. D., Weinheimer, A., Knapp, D. J., Montzka,
- 929 D. D., Holloway, J., Weibring, P., Flocke, F., Zheng, W., Toohey, D., Wennberg, P. O.,
- 930 Wiedinmyer, C., Mauldin, L., Fried, A., Richter, D., Walega, J., Jimenez, J. L., Adachi,
- 931 K., Buseck, P. R., Hall, S. R. and Shetter, R.: Emissions from biomass burning in the
- 932 Yucatan, Atmos Chem Phys, 9(15), 5785–5812, doi:10.5194/acp-9-5785-2009, 2009.
- 933 Yuan, B., Koss, A., Warneke, C., Gilman, J. B., Lerner, B. M., Stark, H. and Gouw, J. A.
- de: A high-resolution time-of-flight chemical ionization mass spectrometer utilizing
- hydronium ions (H_3O^+ ToF-CIMS) for measurements of volatile organic compounds in

- the atmosphere, Atmospheric Meas. Tech., 9(6), 2735–2752,
- 937 doi:https://doi.org/10.5194/amt-9-2735-2016, 2016.

938 Zarzana, K. J., Selimovic, V., Koss, A. R., Sekimoto, K., Coggon, M. M., Yuan, B.,

- 939 Dubé, W. P., Yokelson, R. J., Warneke, C., Gouw, J. A. de, Roberts, J. M. and Brown, S.
- 940 S.: Primary emissions of glyoxal and methylglyoxal from laboratory measurements of
- 941 open biomass burning, Atmospheric Chem. Phys., 18(20), 15451–15470,
- 942 doi:https://doi.org/10.5194/acp-18-15451-2018, 2018.
- 943 Zhou, S., Young, C. J., VandenBoer, T. C., Kowal, S. F. and Kahan, T. F.: Time-
- 944 Resolved Measurements of Nitric Oxide, Nitrogen Dioxide, and Nitrous Acid in an
- 945 Occupied New York Home, Environ. Sci. Technol., 52(15), 8355–8364,
- 946 doi:10.1021/acs.est.8b01792, 2018.

947	,	
948	}	
949)	
950		
951		
952		
953		
954		
955		
956		
957		
958		
959		
960		
961		
962		
963		
964		
965		
966		
967 968		
968 969		
969 970		
970 971		
971 972		
973		
974		
975		
976		
977		
978		
979		

0.13 0.93 0.29 0.94 0.21 0.93		_		0.44 0.93	0.19 0.94	0.28 0.92	0.29 0.93	0.18 0.94	0.20 0.95	0.47 0.87	0.36 0.93	0.53 0.89	0.16 0.95	0.89	0.93	0.32 0.93	0.94	0.93	0.95	0.95
NO3	14.3			14.8				14.5		11.5		12.7								
δ ¹⁵ N-p- NO ₃ ⁻	-7.5			-7.4				-10.6		<u> </u>		-8.9								
δ ¹⁵ N- biomass	0.3	0.3	0.3	-3.4	-3.4	-2.4	-2.8	-1.4	-2.0	-1.4	-1.9	-2.6	0.9	-3.5	-1.4	-0.1	-3.1	-4.2	-2.3	-2.6
δ ¹⁵ N- NO _x	-1.1	2.3	-3.6	-1.1	1.4	0.1	-1.3	1.9	3.3	4.3	2.1	3.4	5.2			7.0				
NO _x (ppbv)	147.9	124.7	716.8	170.8	94.7	91.7	73.6	229.7	571.8	36.2	70.0	95.5	443.3			73.3				
δ ¹⁸ Ο- HONO	12.6	11.6	10.6	8.8	8.4	14.6	9.5	5.2	15.2	8.5	14.9	9.9	7.5	14.8	14.0	14.8	14.9	15.2	10.2	12.2
δ ¹⁵ N- HONO	-5.3	1.7	-3.1	-2.3	-1.9	-1.7	-4.8	-0.5	-0.4	-4.6	0.1	2.1	5.8	6.1	2.5	5.3	3.0	0.3	1.9	0.5
(vdqd) ONOH	19.9	35.8	152.9	74.8	17.6	25.7	21.3	42.2	112.3	17.1	25.3	51.0	70.0	47.1	45.3	23.8	52.5	9.9	40.0	40.8
Fire no.	7	3	4	5	9	8	6	10	11	12	14	15	16	17	18	19	20	21	22	23
Biomass	OdId	OdId	OdId	PICO	PICO	PIEN	PIEN	PSME	PSME	PIEN-d	PSME	ABLA-c	PIPO-I	PIEN-c	PSME-c	PIPO-c	PICO-c	PICO-I	PSME-1	ABLA-c



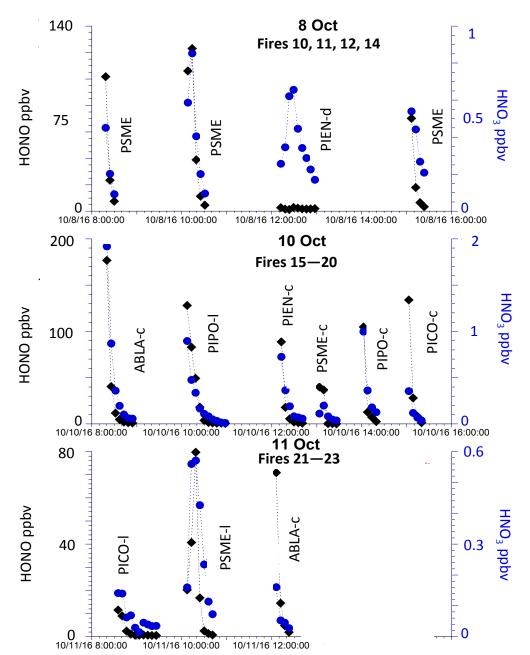
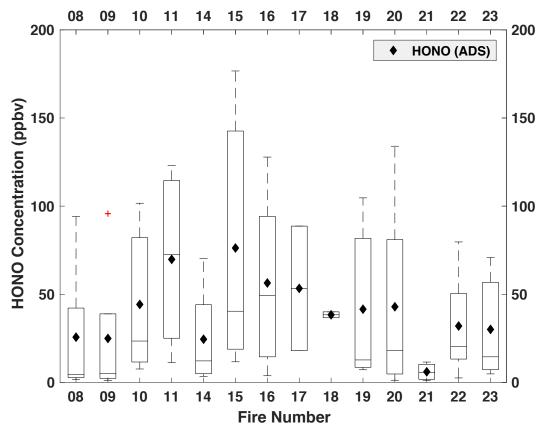
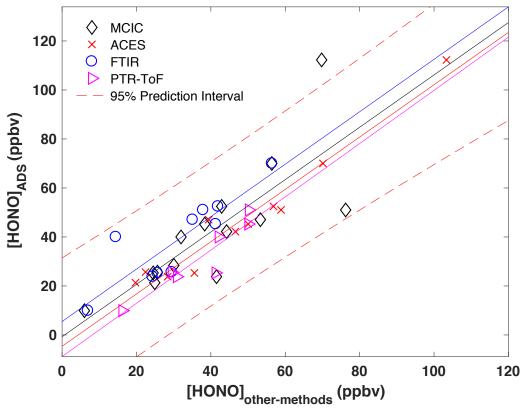


Figure 1. Temporal profile of HONO (black diamond) and HNO₃ (blue circle)
concentration measured using MC/IC method for various stack fires (fire numbers are
referred to Table 1).



992 993 Figure 2. Box plot of MC/IC HONO measurement with 5 minutes resolution over the 994 course of each fire. Each box whisker represents 5th, 25th, 50th, 75th, 95th percentile of 995 HONO concentration during each collection period. Black diamond is the mean HONO 996 concentration recovered from ADS collection. The red cross symbolize outliers. Note no 997 isotopic analysis was performed for fire no. 12 (shown in Figure 1) due to insufficient 998 amount of collected nitrite. 999



 1001
 1002 Figure 3. Comparison of ADS measured HONO concentration with mean values of 1003 various high resolution methods including MC/IC, FTIR, ACES and PTR-ToF for 1004 available fires. Solid lines are linear regression of each dataset with the same symbol

1005 color.

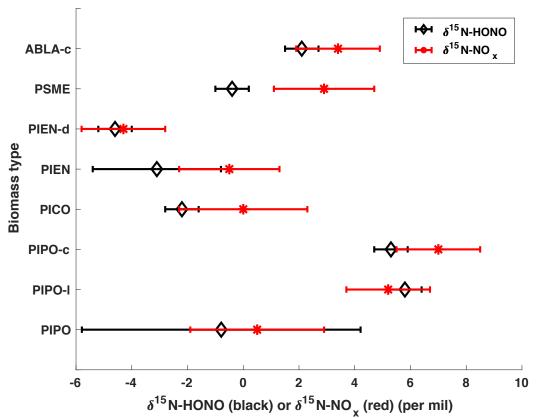
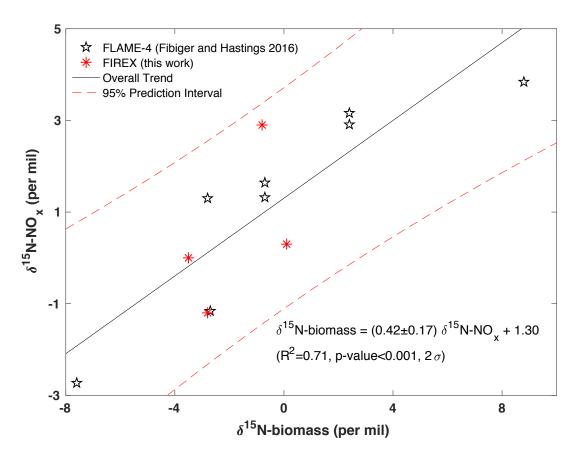




Figure 4. Concentration weighted mean δ^{15} N- of HONO and NO_x versus biomass type. The error bars are propagation of replicate ±1 σ uncertainty (when n>1) and method uncertainty; otherwise, the error bars stand for method uncertainty. PIPO is ponderosa pine, PICO is lodgepole pine, PIEN is Engelmann spruce, PSME is Douglas-fir, ABLA is

1011 subalpine (from Fish Lake, canopy). l indicates litter, c indicates canopy, d indicates duff.





1021 Figure 5. Dependence of δ^{15} N-NO _x on δ^{15} N-biomass. Star data points represent	sent results
---------------------------------------------------------------------------------------------------------------------	--------------

from FLAME-4 study (Fibiger and Hastings, 2016); Asterisk data points represent results from this work; solid line is linear regression between δ^{15} N-NO_x and δ^{15} N-biomass for

the combined dataset; dashed lines are 95% prediction interval (2σ) .

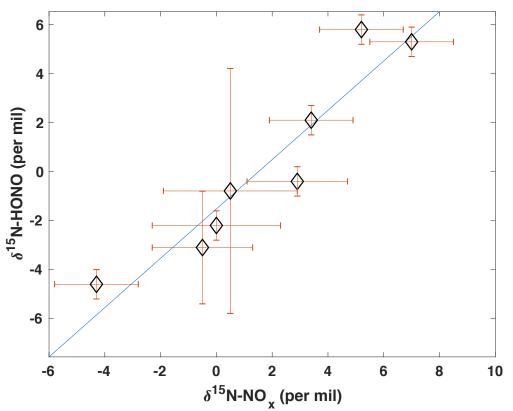




Figure 6. Scatter plot between δ^{15} N-HONO and δ^{15} N-NO_x. All error bars are propagation of replicate uncertainty (±1 σ) and method uncertainty. Linear regression follows δ^{15} N-HONO = 1.01 δ^{15} N-NO_x - 1.52 (R² = 0.89, p<0.001).

Supplement

Isotopic characterization of nitrogen oxides (NO_x), nitrous acid (HONO), and nitrate (NO₃ (p)) from laboratory biomass burning during FIREX

Jiajue Chai¹, David J. Miller^{1,a}, Eric Scheuer², Jack Dibb², Vanessa Selimovic³, Robert Yokelson³, Kyle J. Zarzana^{4,5,b}, Steven S. Brown^{4,6}, Abigail R. Koss^{4,5,6,c}, Carsten Warneke^{5,6}, Meredith Hastings¹

- 1. Department of Earth, Environmental and Planetary Sciences, and Institute at Brown for Environment and Society, Brown University, Providence, RI, USA
- 2. Institute for the Study of Earth, Ocean and Space, University of New Hampshire, Durham, NH, USA
- 3. Department of Chemistry, University of Montana, Missoula, USA
- 4. Chemical Sciences Division, NOAA Earth System Research Laboratory, Boulder, CO, USA
- 5. Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, CO, USA
- 6. Department of Chemistry, University of Colorado, Boulder, CO, USA
- a. Now at: Environmental Defense Fund, Boston, MA, USA
- b. Now at: Department of Chemistry, University of Colorado, Boulder, CO, USA
- c. Department of Civil and Environmental Engineering, Massachusetts Institute of Technology, Cambridge, MA, USA

Correspondence: Jiajue Chai (jiajue_chai@brown.edu)

Experimental details

NO_x online concentration measurement

NO and NO_x concentrations were measured with a Thermo Scientific Model 42i chemiluminescence NO/NO_x analyzer, with ±0.4 ppbv precision and 0.2 ppbv zero noise at 1 minute time resolution. In the NO channel, O₃ generated by an ozonator titrates NO to excited state NO₂ which subsequently produces luminescence that is proportional to NO concentration. In the NO_x channel, the sample gas stream first flows through a heated molybdenum catalyst (325 °C) that converts NO₂ to NO before entering the NO+O₃ reaction chamber. The auto cycle mode (NO/NO_x) switches the mode solenoid valve automatically on a 10 second cycle so that NO, NO₂, and NO_x concentrations are determined. It is known that some NO_y species including HONO, HNO₃, organic nitrate and PAN can be partially converted to NO in the hot molybdenum catalyst, causing positive artifacts in measured NO_x (Reed et al., 2016). In this study, only the HONO interference was corrected for. This was done by subtracting the ADS measured HONO concentration (mean value across each whole fire) from Thermo analyzer measured NO_x

concentration averaged across the whole fire; this provided the approximate lower limit of the NO_x concentration by assuming HONO is 100% converted to NO on the Molybdenum catalyst (e.g. (Dunlea et al., 2007; Febo et al., 1995). Contributions from HNO₃, PAN and gaseous organic nitrate are not of major concern because no photooxidation is involved in indoor fires (Koss et al., 2018; Selimovic et al., 2018; Stockwell et al., 2014). In addition, we do not expect that other reactive nitrogen species such as NH₃ and hydrogen cyanide (HCN) interfere with NO₂ measurement. A particulate matter filter (Millipore, 1µm PTFE) was always placed before the inlet of the NO_x analyzer. The NO channel was calibrated before and after the entire Fire Lab experiments with standard NO (10 ppmv in N₂) diluted with zero air (Thermo Fisher Scientific, Model 111) via a gas dilution calibrator (Thermo Fisher Scientific, Model 146i) and NO₂ response of the NO_x channel using O₃ titration is within ±5% accuracy. The NO_x measurement verified the concentration of the NO_x collected for isotopic analysis, and the original NO_x data is available in the NOAA FIREX archive (FIREX, 2016).

Table S1 Information of fuels measured for δ^{15} N-biomass. Acronyms: ponderosa pine (PIPO), lodgepole pine (PICO), Engelmann spruce (PIEN), Douglas-fir (PSME) and subalpine fir (ABLA). Each fuel is a mixture of one or multiple compositions (different parts from the vegetation) including duff, litter, canopy, rotten, shrub. δ^{15} N of each composition (5th column) was measured in replicates using the method described in section 2.2.3. Composition mass weighted δ^{15} N in 6th column are calculated by mass weighting δ^{15} N of each composition (5th column) with nitrogen content (=sample weight×%N). Mixture mass weighted δ^{15} N (8th column) is calculated by mass weighting δ^{15} N (6th column) with fraction in mixture (7th column).

Sample	Fuel Compo.	Sample Weight (mg)	%N	δ ¹⁵ N (‰)	Compo. Mass Weighted $\delta^{15}N \%$	Fraction in Mixture	Mixture Mass Weighted $\delta^{15}N$ (‰)
PIPO	Duff	4.87	1.1 1	0.32	0.39	0.16	0.1
PIPO	Duff	5.00	1.1 1	0.31			
PIPO	Duff	5.36	1.2 0	0.51			
PIPO	Litter	4.75	0.5 7	1.27	0.94	0.29	
PIPO	Litter	7.60	0.5 4	0.59			
PIPO	Canopy	4.76	0.9	-0.11	-0.10	0.31	

			7				
PIPO	Canopy	5.16	0.9 7	-0.10			
PIPO	Rotten	7.06	0.1 9	1.15	-1.33	0.18	
PIPO	Rotten	10.14	0.1 7	-2.29			
PIPO	Rotten	10.30	0.1 6	-1.55			
PIPO	Rotten	10.37	0.1 8	-2.82			
PICO	Duff	4.69	0.5 1	-2.95	-2.53	0.20	-3.5
PICO	Duff	16.31	0.4 2	-1.83			
PICO	Duff	10.58	0.6 8	-2.63			
PICO	Litter	4.45	0.8 4	-2.73	-3.09	0.11	
PICO	Litter	4.75	0.9 1	-3.38			
PICO	Litter	7.06	0.8 5	-3.15			
PICO	Canopy	4.45	0.9 3	-4.17	-4.16	0.40	
PICO	Canopy	5.24	0.8 8	-4.14			
PICO	Shrub	4.48	0.9 0	-3.51	-3.36	0.09	
PICO	Shrub	6.60	0.8 8	-3.21			
PSME	Duff	4.90	0.7 4	-0.08	0.39	0.15	-0.8
PSME	Duff	9.95	0.8 7	0.79			
PSME	Litter	4.53	0.7 2	-2.41	-2.30	0.11	
PSME	Litter	6.69	0.7 2	-2.19			
PSME	Canopy	4.66	0.8 7	-2.59	-2.33	0.46	
PSME	Canopy	5.99	0.8 6	-2.08			
PSME	Rotten	7.76	0.3 1	2.02	1.67	0.28	
PSME	Rotten	7.08	0.3	1.74			

			1				
PSME	Rotten	10.09	0.3 0	1.23			
Chamise	Canopy	5.27	1.1 4	-3.03	-2.84		
Chamise	Canopy	5.42	1.1 4	-2.66			
PIEN	Duff	4.68	1.3 2	-1.38	-1.41	0.17	-2.8
PIEN	Duff	5.03	1.4 0	-1.43			
PIEN	Canopy	4.59	0.9 5	-3.95	-3.50	0.31	
PIEN	Canopy	6.19	0.9 5	-2.70			
PIEN	Canopy	5.48	0.9 8	-3.84			
ABLA	Duff	5.25	1.1 7	-1.57	-1.40		
ABLA	Duff	6.47	1.2 0	-1.25			
ABLA	Litter	4.38	1.0 0	-4.02	-3.85		
ABLA	Litter	6.24	0.9 1	-3.66			

Table S2 Comparisons between [HONO]_{ADS} with mean values of various high resolution methods including MC/IC, FTIR, CES and PTR-ToF. Missing data points are results of instrumental issues.

Fire no.	ADS(ppb)	MCIC(ppb)	CES(ppb)	FTIR(ppb)	PTR-ToF (ppb)
8	25.7	25.7	22.4	29.5	29.5
9	21.3	24.9	19.7		
10	42.2	44.2	46.6		
11	112.3	69.8	103.3		
14	25.3	24.5	35.6	25.7	41.3
15	51.0	76.2	58.9	37.9	50.2
16	70.0	56.4	70.1	56.4	
17	47.1	53.3	39.4	35.1	
18	45.3	38.3	50.0	41.3	50.0
19	23.8	41.5	28.4	24.3	30.9
20	52.5	42.9	56.8	41.9	
21	9.9	6.0		7.0	16.2
22	40.0	32.0		14.5	42.1

Fire #	NO _x _analyzer	NO _x _collected
THC #	(ppb)	(ppb)
2	113.5	147.9
3	151.3	124.7
5	182.7	123.8
6	60.2	94.7
7	313.0	398.3
8	100.5	91.7
9	80.5	73.6
10	156.2	229.7
11	498.9	571.8
12	33.9	36.2
14	39.5	70.0
15	38.9	43.5
16	338.3	443.3
19	84.3	73.3

 Table S3 Data for NOx concentration measured by NOx analyzer and NOx collection system.

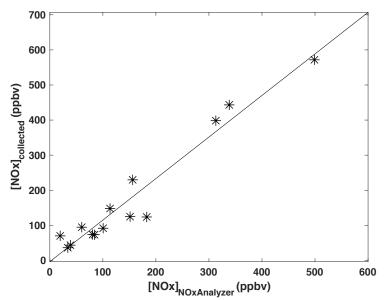


Figure S1 NO_x concentration comparison between NO_x analyzer measurement (mean value over the entire fire) and NO_x collected by the collection system for isotopic

analysis. Solid line is linear regression of the dataset: $y = (1.18\pm0.08)x + (-3.5\pm17.2)$, with $R^2 = 0.94$, $p_{slope} < 0.001$, $p_{intercept}=0.84$, and uncertainty=1 σ .

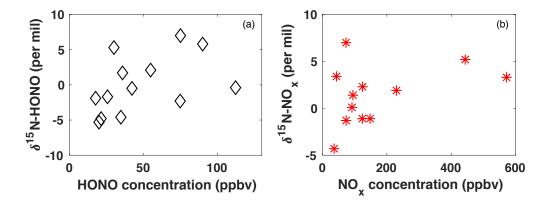


Figure S2. Relationship between δ^{15} N value versus concentration for HONO (a) and NO_x (b). p-values for linear correlation are 0.12 (a) and 0.93 (b) respectively.

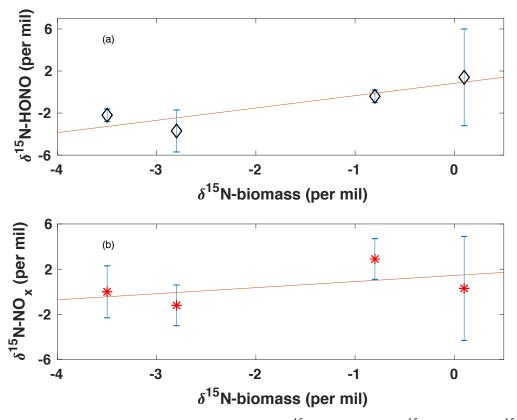


Figure S3. Linear regressions between (a) δ^{15} N-HONO and δ^{15} N-biomass, δ^{15} N-HONO = 1.2 δ^{15} N-biomass + 0.80 (r²=0.83, p_{slope}=0.1), and (b) δ^{15} N-NO_x and δ^{15} N-biomass

 δ^{15} N-NO_x = 0.54 δ^{15} N-biomass + 1.4 (r²=0.28, p_{slope}=0.5). The error bars are propagation of replicate uncertainty (1 σ) and method uncertainty.

References

Dunlea, E. J., Herndon, S. C., Nelson, D. D., Volkamer, R. M., San Martini, F., Sheehy, P. M., Zahniser, M. S., Shorter, J. H., Wormhoudt, J. C., Lamb, B. K., Allwine, E. J., Gaffney, J. S., Marley, N. A., Grutter, M., Marquez, C., Blanco, S., Cardenas, B., Retama, A., Ramos Villegas, C. R., Kolb, C. E., Molina, L. T. and Molina, M. J.: Evaluation of nitrogen dioxide chemiluminescence monitors in a polluted urban environment, Atmospheric Chemistry and Physics, 7(10), 2691–2704, doi:https://doi.org/10.5194/acp-7-2691-2007, 2007.

Febo, Antonio., Perrino, Cinzia., Gherardi, Monica. and Sparapani, Roberto.: Evaluation of a High-Purity and High-Stability Continuous Generation System for Nitrous Acid, Environ. Sci. Technol., 29(9), 2390–2395, doi:10.1021/es00009a035, 1995.

FIREX: FIREX 2016 Fire Lab Data Archive, [online] Available from: https://esrl.noaa.gov/csd/groups/csd7/measurements/2016firex/FireLab/DataDo wnload/index.php?page=/csd/groups/csd7/measurements/2016firex/FireLab/Da taDownload/ (Accessed 25 March 2019), 2016.

Koss, A. R., Sekimoto, K., Gilman, J. B., Selimovic, V., Coggon, M. M., Zarzana, K. J., Yuan, B., Lerner, B. M., Brown, S. S., Jimenez, J. L., Krechmer, J., Roberts, J. M., Warneke, C., Yokelson, R. J. and Gouw, J. de: Non-methane organic gas emissions from biomass burning: identification, quantification, and emission factors from PTR- ToF during the FIREX 2016 laboratory experiment, Atmospheric Chemistry and Physics, 18(5), 3299–3319, doi:https://doi.org/10.5194/acp-18-3299-2018, 2018.

Reed, C., Evans, M. J., Carlo, P. D., Lee, J. D. and Carpenter, L. J.: Interferences in photolytic NO₂ measurements: explanation for an apparent missing oxidant?, Atmospheric Chemistry and Physics, 16(7), 4707–4724, doi:https://doi.org/10.5194/acp-16-4707-2016, 2016.

Selimovic, V., Yokelson, R. J., Warneke, C., Roberts, J. M., de Gouw, J., Reardon, J. and Griffith, D. W. T.: Aerosol optical properties and trace gas emissions by PAX and OP-FTIR for laboratory-simulated western US wildfires during FIREX, Atmos. Chem. Phys., 18(4), 2929–2948, doi:10.5194/acp-18-2929-2018, 2018.

Stockwell, C. E., Yokelson, R. J., Kreidenweis, S. M., Robinson, A. L., DeMott, P. J., Sullivan, R. C., Reardon, J., Ryan, K. C., Griffith, D. W. T. and Stevens, L.: Trace gas emissions from combustion of peat, crop residue, domestic biofuels, grasses, and other fuels: configuration and Fourier transform infrared (FTIR) component of the fourth Fire Lab at Missoula Experiment (FLAME-4), Atmos. Chem. Phys., 14(18), 9727–9754, doi:10.5194/acp-14-9727-2014, 2014.