

General Response: We thank the reviewer for your helpful comments. We have addressed all comments and provided point by point response below. The revised manuscript is presented in below.

Response to the referee#2' comments

Overall Comments: The authors use multiple measures to investigate the physical properties of non-sea salt aerosols collected in the Arctic. They qualitatively describe the mixing state of organic, soot, and sulfate aerosols. Although the core data set of this manuscript is informative, as is, the presentation and discussion are confusing and possibly misleading. No large scale descriptive statistics are provided, the methods and presentation are often unclear and/or redundant, the environmental context is lacking, and the optical conclusions are not well explained. This manuscript would benefit from additional analysis, clarification of methods, adding proper context to results, and careful grammar review.

Response: We carefully addressed all of the questions below. We attempted to re-write the manuscript. However, we are not able to address non-specific questions.

Unclear Sample Selection: In multiple sections of this paper the authors discuss a subset of their samples without explaining why they were chosen. This detracts from their discussion and leaves the reader questioning what is not being discussed. Anytime a subset of samples is chosen or used, an explanation needs to be provided explaining why these are the best samples for that specific analysis or investigation. A few example occurrences can be found in lines 275, 281, 118, 190.

Response: Thank you for good comments. We received some comments before and revised the manuscript several rounds. After that, we added Figure S1 showing what samples were analyzed by different instruments. We removed some number which confused for the readers and rephrased some sentences in the Methods section.

To SEM and NanoSIMS analysis after TEM observations, we need to select better samples. These samples should keep good conditions of carbon film because some samples have broken carbon film during the sampling period. In the NanoSIMS and SEM analysis, these samples need to be stucked on the sample stage, they could not be recycled again.

Please see response to reviewer 1, comments 2-4.

Environmental Context: The only environmental context explored in this text is the back trajectories provided for 4 of the sampling days. Whenever field data is discussed, external variables such as temperature, humidity, and back trajectories can help explain variability observed in the data. To improve this work, the context of the sampling should be explored as a driving factor between differences in particle characteristics.

Response: We added the RH, T, P in the table S1. These meteorological data were recorded during the sampling period. Here we did back trajectories of each sampling day as shown in Figure 1.

We added the following to the revised manuscript (line 144 to 146)

“The sample information such as local sampling date and time and meteorological conditions

(e.g., temperature (T), relative humidity (RH), pressure (P), wind direction (WD), wind speed (WS)) were listed in Table S1.”

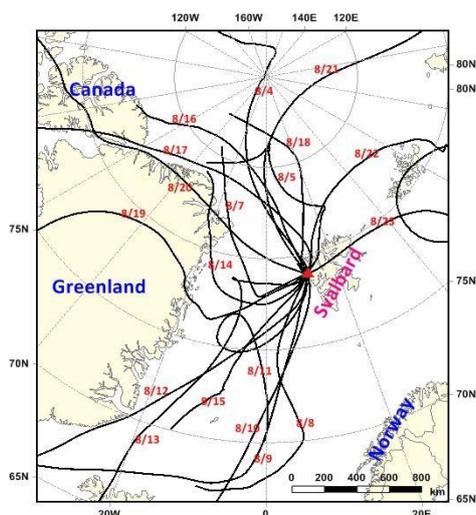


Figure 1 72 h back trajectories of air masses at 500m over Arctic Yellow River Station in Svalbard during 3–26 August 2012, and arriving time was set according to the sampling time

2.1 I am also not convinced that the daily averaged back trajectory calculations are meaningful for these samples. (I’m also not sure the trajectories were daily averaged as the methods section just said they were calculated for a given day, not the frequency of times or given time on the day.) Because the sample length was between 20 minutes and 2 hours, a back trajectory occurring during the actual sampling time for all samples would be more appropriate.

Response: We did choose the mid of the sampling time as the end point for the back trajectory analysis (see Figure 1)

3 Optical Property Calculations: The authors attempt to estimate the potential radiative forcing implications of their results using a core-shell model and Mie theory. Although this is a potentially meaningful result, the methodology needs to be discussed further and improved.

Response: We added more description in the Method section.

L263-271“BHCOAT Mie code by Bohren and Huffman (1983) was used to calculate the optical properties, including scattering cross section (SCS), absorption cross section (ACS), and single scattering albedo (SSA), assuming a core-shell structure. We firstly calculated these parameters assuming a sulfate core and OM shell structure only (ignoring some of the particles that contain soot core). Because the Mie code only can calculate the core-shell structure or homogeneous models, we assume sulfate as a core and OM as a shell in individual particle to build the core-shell model. Based on the core-shell standard mode (Li et al., 2016), we can calculate optical properties of individual internally mixed particles.”

3.1 When performing Mie theory calculations the input size distribution characteristics and wavelength dependence (angstrom exponent) have dramatic effects on the results. These data

need to be reported. The authors state that a single size distribution is used. How was this size distribution established? Could you run your analysis over all of the size distributions you observed to estimate the variability?

Response: Size distributions of particles are measured by microscopes These measurements can help us to evaluate sulfate and OM volume and then we made the Figure 8d. Once we obtained these basic data, we can input all the data to calculate optical properties of individual particles. Here the result is different from the bulk online optical properties. As the comments, we did correction between optical absorption cross section (ACS) and particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating. Interestingly, we found that ACS increase along with the particle diameter increase through one nonlinearity showed as below.

In this study, it is not necessary to consider wavelength dependence because lots of work how different size of aerosol particles interacted with the different wavelength. We just did pilot cases to confirm how BrC influence the optical absorption of sulfate particles based on our TEM data (Figure 8). The purpose of this study draws attention for BrC in the Arctic air.

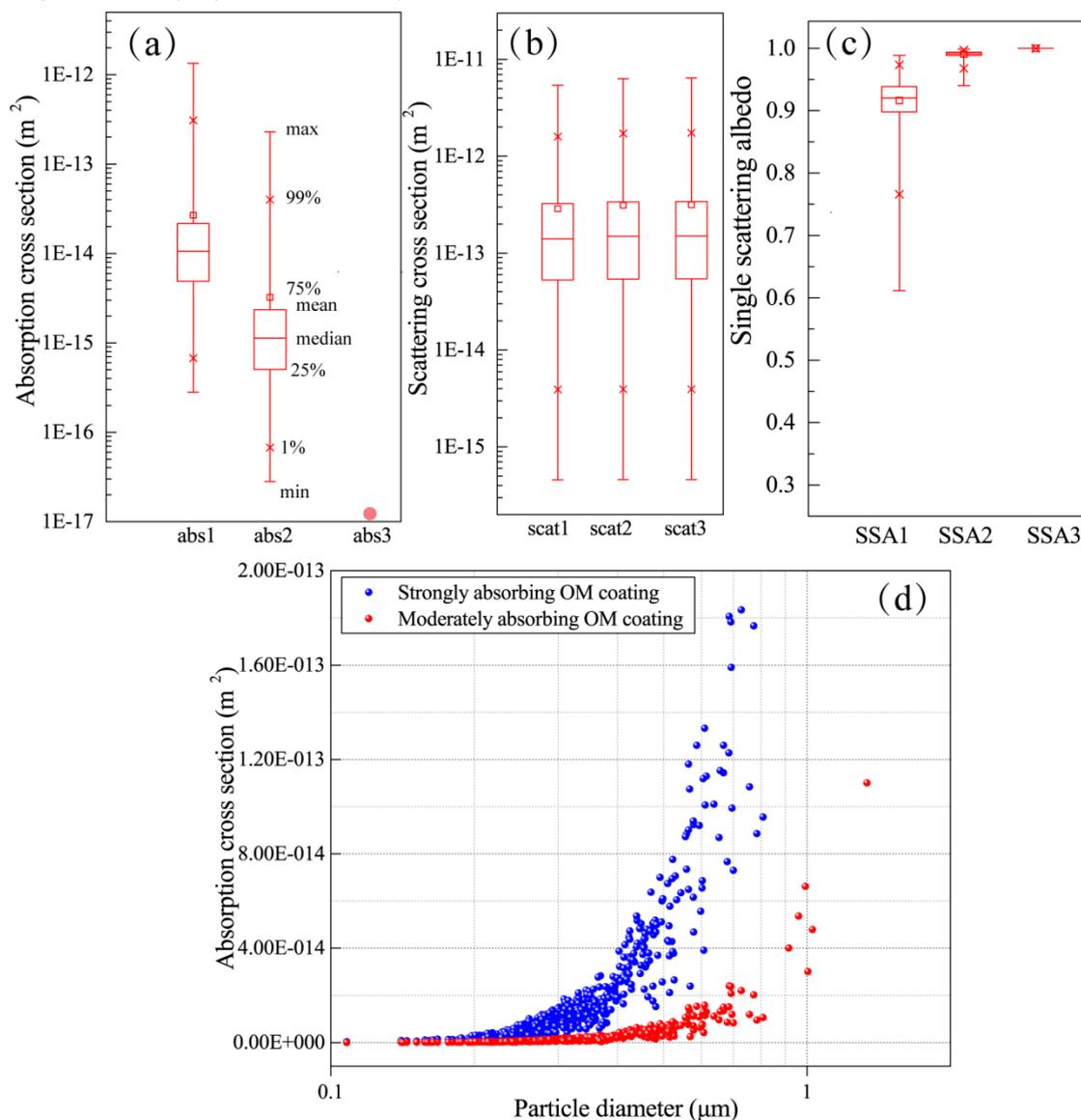


Figure 8 Optical properties of Box-and-whisker plots showing optical parameters of all

analysed particles assuming sulfate core and BrC shell (not considering soot cores in the particles). (a) Scattering cross section (b) Absorption cross section (c) Single scattering albedo. Top to bottom markers in the box-and-whisker represent max, 99%, 75%, mean, median, 25%, 1%, min values. (d) Absorption cross section along with particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating.

3.2 The way the authors include soot in their calculations is not appropriate, as soot is treated as a core with a mixed sulfate-OM shell. This is in direct conflict with their observations that soot is only observed to be associated with the OM shell and sulfate is always the core. It would be better not to include a soot calculation than include a misleading one.

Response: We deleted the soot part here. The updated data is shown in Figure 8

3.3 The relevancy of the refractive index (RI) choice needs to be discussed further. The authors use an RI from biomass burning brown carbon as their slightly absorbing case but frequently state that they believe the OM to be secondary. If that is the case, their estimations are likely too high for all cases but the non-absorbing case. The authors should include refractive indexes from secondary organic aerosol brown carbon to get a more realistic answer. These are generally much lower than the RI values that they used. Additionally, the scattering component chosen (1.65) is relatively high and not explained. Please explain why this value was chosen as it is especially important in core-shell cases.

Response: There are many different RIs from the laboratory experiments. For example, Jiang et al., (2019) suggested the measured RIs on different organic species. The RI is dependence on the wavelength. Fortunately, we just want to test how OM coating influence the sulfate particles. Based on the relationship between RIs and wavelength (Jiang et al., 2019; Feng et al., 2013), the RI doesn't influence our conclusions in this study.

Reference added: Jiang, H., A. L. Frie, A. Lavi, J. Y. Chen, H. Zhang, R. Bahreini, and Y.-H. Lin (2019), Brown Carbon Formation from Nighttime Chemistry of Unsaturated Heterocyclic Volatile Organic Compounds, Environmental Science & Technology Letters, DOI: 10.1021/acs.estlett.1029b00017.

Here we chose the RI reported in Feng et al., 2013. Through the comparisons between strongly absorbing and moderately absorbing, non-absorbing BrC, we generally knew how BrC coating could influence optical properties of sulfate particles. Although our current study could not get a more realistic answer, but the result suggests that BrC has a potentially role to play in light absorption properties of aerosols Arctic. The conclusion warrants further study in the Arctic area. Our results from the TEM can provide solid evidence about the mixing state of sulfate and other aerosol species. As suggested by the referee, we added a sentence to explain why we selected the 550 nm:

In the context:

"Although the refractive index has dependence on the wavelength between 350-870 nm, we tried to select the 550 nm as a case to test how OM coating influence sulfate particles in Arctic air."

3.4 Stemming from the above comment, the relevancy of each refractive index case needs to be discussed in the context of the Arctic. Which case seems the most likely?

Response: We can't make this judgement because there is no data to show this. We showed that this is potentially important to consider, which warrants further study.

3.5 The authors use $m=1.55$ for sulfate. RI values have wavelength dependence, was this included in any way?

Response: Thanks. It does. We corrected the writing here.

Line 256 to 257: "The refractive index used for the non-light-absorbing sulfate component was set to $m=1.55$ at 550 nm (Seinfeld and Pandis, 2006)."

3.6 The authors average the absorption cross section on a per particle basis. This isn't meaningful since observed absorption cross sections will depend on the whole aerosol population (and the size of the particle). The authors should calculate an absorption cross section for the ambient aerosol concentration during their sampling and compare it to other absorption observations.

Response: Our methods are not able to give the ambient aerosol concentration. We tested how the absorbing OM influenced optical properties of sulfate in the Arctic based our individual particle measurements. These results showed that this is potentially important to consider, which warrants further study.

4 Redundancy and Clarity: The writing of this paper needs improvement. There are multiple times when the authors restate the same point twice or fail to introduce a topic before discussing their results. An example of redundancy can be observed when Copper TEM grids are introduced in lines 120, 140, and 156. An example of an improper discussion occurs in the discussion of satellite particles, which are introduced in line 280 with no context or explanation. This forces the reader to infer what the authors mean by satellite particles, possibly leading them to wrong conclusions. To improve the manuscript, I recommend careful reconsideration of the presentation of the data, with special consideration to avoiding redundancy and ensuring the appropriate context is present.

Response: We carefully revised them. Please see our replies as below associated with your specific comments.

4.1 There are multiple points in the methods, a few of which I've included in the specific comments, that are unclear or confusing. The reader needs to be able to understand exactly which analysis was performed on which filters for how many particles if they are to believe your result.

Response: We made one major revision in this section. We added the Table S1 to make specific information.

5 Sample Information: The only summary of the total data set is provided in figure S6 and this figure states that only 3 samples (of 46 collected and 21 analyzed). No explanation is given as to why is summary is so limited given the authors have EDS data (which they used for classification) of 20-30 particles over 21 samples (at least 400 data points). All the of following discussions only make sense if they are provided in the context of overall sample composition.

Response: The Referee is probably referring to Figure S4 rather than Figure S6. It is a spelling

mistake, which we have corrected. At the revised ms, the Figure was moved to the main ms as Figure 1.

Specific Comments:

1 There are many grammatical errors and redundancies that I have not addressed below.

Response: We carefully improved English writing in the manuscript.

2 The phrase internal mixing is used throughout the manuscript without an explicit definition. An explanation of what exactly you mean when you say something is “internally mixed” would improve the manuscript

Response: We added the following to the text

Add an explanation the first time you write it, e.g., line 88

“Internal mixing means that a single particle simultaneously contains two or more types of aerosol components (Li et al., 2016).”

3 Line 55: Change “Arctic” to “the Arctic”

Response: Corrected

4 Line 60: Change “nature” to “natural”

Response: Corrected

5 Line 64-66: Restructure this sentence for clarity. Treat the percentages in a consistent way as to not confuse the reader. For example, this sentence could be changed to:” For example, Winger et al.(2017) showed most Arctic BC is sourced from domestic activities (35%) and transportation (38%), with only minor contributions from gas flaring (6%), power plants (9%), and open fires (12%)”

Response: Corrected

6 Line 72: I’m not sure exactly how this sentence fits in with the brown carbon theme of this paragraph. Are these compounds commonly found in brown carbon or organic aerosols in general? Please add some context.

Response: Thank you very much. We move the positions of this sentence and add more description here.

L73-78: “Accumulation of secondary organic aerosol, a significant fraction of the new particles grow to sizes that are active in cloud droplet formation in the Arctic (Abbatt et al., 2019). More than 100 organic species were detected in the Arctic aerosols and polyacids were found to be the most abundant compound class, followed by phthalates, aromatic acids, fatty acids, fatty alcohols, sugars/sugar alcohols, and n-alkanes (Fu et al., 2008).”

7 Line 84: replace “were” with “have been”

Response: Corrected

8 Line 96-68: This sentence is confusing. Please rewrite it more concisely and clearly.

Response: Revised as the following (line 106 to 109)

“The poor understanding on mixing state of BC and BrC in individual particles will prevent the further simulation of atmospheric climate and aerosol-cloud interaction in the Arctic through the

current atmospheric models (Browse et al., 2013; Samset et al., 2014; Zanatta et al., 2018)."

9 Line 99-100: Change "collected on 7 to 23 August, 2012 in the Arctic." to "collected in the Arctic between August 7th and 23rd, 2012."

Response: Corrected.

10 Line 104: replace "on substrate" with "on a substrate"

Response: deleted the word here

11 Line 118: Change "samples between 7 and 23 August, 2012." to "samples collected between August 7th and 23rd, 2012."

Response: Corrected

12 Line 119: Replace "analyzed for TEM analysis" with "analyzed with TEM"

Response: Revised

13 114-131: Restructure your sampling section. As written the reader may think that you sampled with 2 separate samplers an individual particle sampler and cascade impactor. After reading the paper, there is only one sampler. This confusion can be remedied by introducing the cascade impactor earlier.

Response: Revised the part as below (line 133 to 136)

"A sampler containing a single-stage impactor with a 0.5-mm-diameter jet nozzle (Genstar Electronic Technology, China) was used to collect individual particles by the air flow rate at 1.5 l min⁻¹. Aerosol particles were collected onto copper TEM grids coated with carbon film."

14 122-123: Add the top size cutoff for this sampler.

Response: Added

15 138: replace "within a" with "for a"

Response: Corrected

16 148: 2002 particles examined over all the samples, or in a specific filter?

Response: Added the Table S1 in the sentence. Table S1 can show how many particles we analyzed in each sample.

17 151: Clarify this sentence. Do you mean to say you "we only checked elemental compositions of 20-30 particles" in each sample?

Response: We added more detailed description (line 165 to 180)

"In the clean Arctic air, there are simply particle types including sea salt, sulfate, soot, and OM. Because soot particles have chain-like aggregation, it is not necessary to check their elemental composition. Sea salt particles display spherical or square shapes and are stable under the electron beam in TEM but sulfate particles are spherical but flats on the substrate and produce unstable bubble under the electron beam (Buseck and Posfai, 1999; Chi et al., 2015). TEM observations also can clearly identify sulfate particles or sulfate with OM coating. Therefore, we

can easily identify Arctic particle types based on their morphology. Because of the time-consuming in the experiment, it is not necessary to frequently check elemental composition of the same particle type. For the data statistic in this study, we randomly checked elemental composition of 20-30 particles in each sample (Table S1). EDS spectra of 575 particles were manually selected and saved in the computer for elemental composition analysis. Particles examined by TEM were dry at the time of observation in the vacuum of the electron microscope. In our study, the effects of water and other semi-volatile organics were not considered as they evaporate in the vacuum.”

18 155-156: The statement about Cu is redundant. This has already been stated in line 120 and 140.

Response: Corrected:

19 156-157: What is the difference between what is stated here and what is stated in lines 147-149?

Response: In lines 147-149 describe how to analyze particles on the substrate using TEM/EDS. In 156-157, we used scanning TEM. The method is one other function in the TEM which can give element profile along with the one line or elemental mapping in the targeted individual particle. We added more description here and tell the reader why we did the STEM here.

The following paragraph was revised to (line 181 to 188)

“ Elemental mapping and line profile of individual aerosol particles were obtained from the EDX scanning operation mode of TEM (STEM). The STEM information can clearly display elemental distribution in the targeted individual particles which cannot be provided by the above EDS examination. Based on preliminary individual analysis, we further chose the typical samples containing abundant sulfate with OM coating for the STEM analysis. The high-resolution details of elemental distribution in individual particles can further prove the details of the mixing structure of sulfate and OM in individual particles.”

20 162-163: Replace “is the image analysis platform...” with “is an image analysis platform”.

Response: Corrected

21 177: Replace “Organic Matters” with “Organic Matter”

Response: Corrected

22 Line 184: Replace “TEM grids was” with either “The TEM grid was” or “TEM grids were”

Response: Corrected

23 Line 192-194: Incomplete sentence

Response: Corrected

“AFM with a digital nanoscope IIIa instrument operating in the tapping mode was used to observe surface morphology of individual aerosol particles and measure particle thickness.”

24 Line 194: I stopped making basic grammar and structure critiques at this point.

Response: Thank you very much. We carefully checked them and improved the English writing. Please see the red markers in the revised ms.

25 Line 222: Does treating this as a core-shell system with BC in the middle and sulfate and OM mixed on the outside have any basis? You've indicated that you have soot inclusions on the outside of predominantly sulfate particles, so why would soot be on the inside?

Response: We've deleted the part with soot – see response in comment above.

26 Line 224: This sentence reads as if you've calculated the refractive index of the particles. Did you measure the optical properties of these particles?

Response: We didn't measure their optical properties. We deleted this sentence and revised the part.

27 Line 228: "In this study". It is unclear whether this is referring to your previous work or this manuscript.

Response: corrected

28 Line 253: Is it possible that coagulation of primary organic particles and S-rich particles could have led to the formation of organic coatings? Are you sure assumption that all organic coatings are secondary valid?

Response: The possibility could happen in the polluted air due to high particle number. As our previous study, Chi et al., (2017) determined sea salt particles collected in the Arctic air. We didn't observe the association of sulfate and sea salts particles, although we found many aged sea salt particles. The observations again suggest the particle coagulation is unlikely to be important in the Arctic air.

29 Line 272: Can you say a percentage of NSS particles that are S-rich with an OM coating? Or a percentage of S-rich particles that have a coating? This would strengthen the paper if an actual number was given.

Response: No, we cannot say NSS particles are S-rich with OM coating particles. Figure S4 shows 39% by number of all the analyzed particles were NSS-particles but 29% particles contain sulfate. Here we found 73% of the analyzed NSS-particles are S-rich with OM coating. We've changed this

line 310 to 312: "Here we focused on S-rich, soot, and OM particles as the major non-sea salt particle (NSS-particle, 39±5%) in the analyzed samples, which are approximately 29±7% of 2002 particles (Figure 3). "

L323-324: "A majority of 781 analyzed NSS-particles (74% by particle number) have a sulfate core and OM coating (Figures 4 and 5)."

30 Line 274-276: Are these specific samples special or is there something that you think may have caused the low frequency of soot inclusions? If so, please explain why.

Response: We only found some fresh soot particles in three samples out of the 21 samples. We

found many sulfate with soot inclusions. Previous works showed that BC is very low in the Arctic. We've changed this (line 326 to 333).

"The mixing structure is different from our previous findings in polluted air that soot is normally mixed with sulfate instead of OM coating (Li et al., 2016). Moreover, we noticed that a few chain-like soot aggregates (1.3% in all analyzed particles) (Figure S5) only occurred in three samples during the sampling period (Table S1). Considering the remoteness of the sampling site, such fresh soot particles are likely to be of local origin, including shipping and flaring (Gilgen et al., 2018; Peters et al., 2011). Indeed, we found a few of ships moving in Arctic Ocean during these days from the Ny-Ålesund town."

31 Line 277: This statement needs to be better supported. Just because a site is remote does not mean particles are local in origin. If this is supported by your trajectory calculations, mention them here. Also, don't the soot inclusions also imply that perhaps the OM is not secondary in nature? Soot is 100% primary and often co-emitted with primary OM, so if there is soot associated with OM coatings the soot itself is primary and so possibly some of the OM is primary as well.

Response: We found fresh soot particles in Figure 6. Indeed, TEM image shows that the soot particles have very thin OM coating.

It is true that soot is often co-emitted with primary OM. However, there are plenty of evidence to show that the primary soot aggregate particles are not heavily mixed with OM and in particularly not coated with thick OM (Wang et al., 2017). Furthermore, It is well known that SOA is dominant in the remote air (Jimenez et al., 2009).

32 Line 280: You need to define what satellite particles are; you have not discussed or defined them previously. Are they simply splatter of liquid portions of the particle when the particle is collected?

Response: Thanks. We define it.

L334-337 "TEM observations showed that some sulfate particles had unique morphology that a sulfate particle was surrounded by some smaller particles (Figure 5a). They are often called "satellite" particles as they were distributed from the central particles when impacted on the substrate during sample collection."

33 Line 280-281: is there a reason why satellite particles would have been observed on these days but not other days?

Response: The samples collected during 9-15, August contained abundant sulfate particles, OM, and soot particles (Table S1). Also, Figure 2 shows that air masses from North American during these days were mainly dominant. The reasons should depend on their air masses. Although the mechanism of this point is interesting, we could not give specific evidence under the current data.

34 Line 281: This is misleading and implies you performed the NanoSIMS analysis on 11 samples. In the methods section, it says only two samples were analyzed with NanoSIMS.

Response: Corrected

L339-342: "NanoSIMS analysis further provided more information that the satellite particles

selected from the samples (Table S1) have strong 32S- (Figure 6a, c) and 16O- signals (Figure 6d) as well as weak 12C14N- signals (Figure 6a, b).”

35 Line 287-289: This is misleading. It reads as if you have done molecular characterization of the organic matter.

Response: Corrected

L346-347 “Indeed, Fu et al. (2008) found that polyacids are the most abundant organic compounds, followed by phthalates, aromatic acids, and fatty acids in Arctic aerosol particles.”

36 Line 314: You back trajectories are only for specific days, be transparent about this in the Discussion

Response: Corrected. Add one Figure S2

37 Line 317-318: This conflicts with your earlier comment that most BC should be local.

Response: The result does not conflict with earlier comments. We only pointed out the fresh soot particles (only 1.3%) may come from local shipping or combustion activities. It doesn't mean most BC were from local sources. In this study, we did not determine where the BC comes from.

38 Line 326-342: This discussion reads like a list of facts, but why they are all relevant is not always stated. Explain why each observation is important and how it adds context to your results.

Response: We carefully revised the part as below

L338-408: “The sulfate core-OM shell structure observed in the Arctic summer atmosphere is similar to those in the background or rural air in other places (Li et al., 2016; Moffet et al., 2013). Based on the images from electron microscopies, we can infer that OM coating thickness in the arctic atmosphere was comparable with them in rural places but higher than them in urban places. During the transports, organic coatings on sulfates were considered as the secondary organic aerosols and their masses increase following particle aging and growth (Li et al., 2016; Moffet et al., 2013; Sierau et al., 2014). Figures 1 and 2 show that most of particles in the air masses transported long distance from North American. The result indicates that these long-range transportation of secondary sulfate particles have enough time to experience the possible atmospheric heterogeneous reactions on particle surfaces or cloud processes in the Arctic air. Similarly, Moffet et al. (2013) found that soot inclusions occurred in OM coating when OM coating on sulfates built up through photochemical activity and pollution buildup the Sacramento urban plume aged. On the other hand, the sulfate/OM particles with soot inclusions are probably formed in a similar way as those found elsewhere (Li et al., 2016) – e.g., soot particles may have acted as nuclei for secondary sulfate or organic uptake during their transports (Riemer et al., 2009). Similarly, besides the OM coating in the Arctic particles, Leck and Svensson (2015) found some biogenic aerosols like gel-aggregate containing bacterium in ultrafine particles. However, we didn't find any gel-like particles in the samples because our sampler had very low efficiency for ultrafine particles.”

39 Line 343: Why is dry included here? Are there also wet particles that you have not discussed?

Response: Deleted the “Dry” here

40 Line 344-347: Explain LLPS in simpler terms and why it's important.

Response: Revised the part.

L409-428: “TEM images show that most of the internally mixed sulfate particles display sulfate core and OM coating on the substrate (Figures 4a and 5b, c). The sulfate and OM separation in individual particles were defined by You et al. (2012) as liquid-liquid phase separation (LLPS). Concerning the knowledges of the LLPS can better understand particle hygroscopicity, heterogeneous reactions between reactive gases on particle surface, and organic aging (You et al., 2012). They also reported that the LLPS can reflect the O:C ratio in the OM, which is roughly ≤ 0.5 . In this study, we did observe the LLPS in almost all the fine sulfate particles, which indicates that the secondary OM in the coating might be not highly aged. Therefore, we speculate that the thick OM coatings were consistently built up during the long range transport of sulfate particles and part of secondary OM in the coating likely formed in Arctic area. Indeed, some studies reported that there are various sources of organic precursors during the Arctic area, such as biogenic VOCs from ice melting and open water (Dall’Osto et al., 2017) and anthropogenic VOCs from shipping emissions in summertime (Gilgen et al., 2018). The dependence of OM volume on particle size (Figure S6) suggests that the suspended sulfate particles are initially important surface for secondary OM formation. Moreover, the common OM coating on sulfate particles indicates that secondary OM as the surfaces of fine particles might govern the possible heterogeneous reactions between reactive gases and sulfate particles in the Arctic air.”

41 Line 346: It’s unclear what 90% is referring to in this statement.

Response: We deleted this sentence here

42 Line 348: There’s likely a complex relationship between phase state, oxidation state, and humidity. This needs to be investigated and explained further if statements about aerosol age are going to be made. Additionally, shouldn’t you see a variety of ages of aerosol in your samples? Showing contrast between aged and unaged particles would be interesting and convincing.

Response: We re-wrote this part. We pointed out that the implications including LLSP in the Arctic air. Almost all particles we observed the aged. We do not have a parameter to define the age of the aerosol population and we are unable to answer the question clearly.

43 Line 376-378: This is circular reasoning because $^{12}\text{C}^{14}\text{N}^-$ was what you used to identify OM so of course it was observed in the OM coatings. I don’t think this data set is appropriate to make conclusions about the N content of OM coatings. That said, if you were able to calculate the mass concentration of N in the coatings with NanoSIMS that might give you a better indication of the BrC potential of the OM.

Response: It is impossible to calculate the N concentration in the coatings based on the NanoSIMS data here. The literature as the below proved that CN^- can indicate N-containing organic matter.

Herrmann, A. M., K. Ritz, N. Nunan, P. L. Clode, J. Pett-Ridge, M. R. Kilburn, D. V. Murphy, A. G.

O'Donnell, and E. A. Stockdale (2007), Nano-scale secondary ion mass spectrometry — A new analytical tool in biogeochemistry and soil ecology: A review article, *Soil Biology and Biochemistry*, 39(8), 1835-1850.

44 Line 390-393: The average absorption cross-section is reported on a particle basis. This would be much more meaningful if it was extrapolated to environmental conditions. Because you've sampled from the atmosphere, you should be able to approximate particle concentrations, correct? If so, you could back calculate this to an actual atmospheric absorption contribution and compare it to expected absorption from other species and measurements. This would be significantly more meaningful.

Response: Yes, we agree it would be useful to extrapolate to environmental conditions but as stated by the reviewers, we do not want to over-interpret our results. We cannot extrapolate because we do not have ambient particle concentration data. One single paper will not be able to answer all questions and we do suggest for future work to answer these questions.

45 Line 413, 260: 29% number should include a standard deviation.

Response: added

46 Line 410: The last section shouldn't simply repeat what was stated in the above sections, but instead present the data in additional context and discuss the implications.

Response: We made major revision in the section. Please see the revised red words.

References:

- Abbatt, J.P.D., Leaitch, W.R., Aliabadi, A.A., Bertram, A.K., Blanchet, J.P., Boivin-Rioux, A., Bozem, H., Burkart, J., Chang, R.Y.W., Charette, J., Chaubey, J.P., Christensen, R.J., Cirisan, A., Collins, D.B., Croft, B., Dionne, J., Evans, G.J., Fletcher, C.G., Galí, M., Ghahremaninezhad, R., Girard, E., Gong, W., Gosselin, M., Gourdal, M., Hanna, S.J., Hayashida, H., Herber, A.B., Hesaraki, S., Hoor, P., Huang, L., Hussherr, R., Irish, V.E., Keita, S.A., Kodros, J.K., Köllner, F., Kolonjari, F., Kunkel, D., Ladino, L.A., Law, K., Lévassieur, M., Libois, Q., Liggio, J., Lizotte, M., Macdonald, K.M., Mahmood, R., Martin, R.V., Mason, R.H., Miller, L.A., Moravek, A., Mortenson, E., Mungall, E.L., Murphy, J.G., Namazi, M., Norman, A.L., O'Neill, N.T., Pierce, J.R., Russell, L.M., Schneider, J., Schulz, H., Sharma, S., Si, M., Staebler, R.M., Steiner, N.S., Thomas, J.L., von Salzen, K., Wentzell, J.J.B., Willis, M.D., Wentworth, G.R., Xu, J.W., Yakobi-Hancock, J.D.: Overview paper: New insights into aerosol and climate in the Arctic, *Atmos. Chem. Phys.*, 19 (4), 2527-2560, 2019.
- Bohren, C.F., Huffman, D.R., 1983. Absorption and scattering of light by small particles. John Wiley & Sons, Inc., New York, USA.
- Browse, J., Carslaw, K.S., Schmidt, A., Corbett, J.J.: Impact of future Arctic shipping on high-latitude black carbon deposition, *Geophys. Res. Lett.*, 40 (16), 4459-4463, 2013.
- Buseck, P.R., Posfai, M.: Airborne minerals and related aerosol particles: Effects on climate and the environment, *P. Natl. Acad. Sci. USA*, 96 (7), 3372-3379, 1999.
- Chi, J.W., Li, W.J., Zhang, D.Z., Zhang, J.C., Lin, Y.T., Shen, X.J., Sun, J.Y., Chen, J.M., Zhang, X.Y., Zhang, Y.M., Wang, W.X.: Sea salt aerosols as a reactive surface for inorganic and organic acidic gases in the Arctic troposphere, *Atmos. Chem. Phys.*, 15 (19), 11341-11353, 2015.
- Dall'Osto, M., Beddows, D.C.S., Tunved, P., Krejci, R., Ström, J., Hansson, H.C., Yoon, Y.J., Park, K.-T., Becagli, S., Udisti, R., Onasch, T., O'Dowd, C.D., Simó, R., Harrison, R.M.: Arctic sea ice melt leads to atmospheric new particle formation, *Sci. Rep.*, 7 (1), 3318, 2017.
- Fu, P., Kawamura, K., Barrie, L.A.: Photochemical and Other Sources of Organic Compounds in the Canadian High Arctic Aerosol Pollution during Winter-Spring, *Environ. Sci. Technol.*, 43 (2), 286-292, 2008.
- Gilgen, A., Huang, W.T.K., Ickes, L., Neubauer, D., Lohmann, U.: How important are future marine and shipping aerosol emissions in a warming Arctic summer and autumn?, *Atmos. Chem. Phys.*, 18 (14), 10521-10555, 2018.
- Jimenez, J.L., Canagaratna, M.R., Donahue, N.M., Prevot, A.S.H., Zhang, Q., Kroll, J.H., DeCarlo, P.F., Allan, J.D., Coe, H., Ng, N.L., Aiken, A.C., Docherty, K.S., Ulbrich, I.M., Grieshop, A.P., Robinson, A.L., Duplissy, J., Smith, J.D., Wilson, K.R., Lanz, V.A., Hueglin, C., Sun, Y.L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J.M., Collins, D.R., Cubison, M.J., E., Dunlea, J., Huffman, J.A., Onasch, T.B., Alfarra, M.R., Williams, P.I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J.Y., Zhang, Y.M., Dzepina, K., Kimmel, J.R., Sueper, D., Jayne, J.T., Herndon, S.C., Trimborn, A.M., Williams, L.R., Wood, E.C., Middlebrook, A.M., Kolb, C.E., Baltensperger, U., Worsnop, D.R.: Evolution of Organic Aerosols in the Atmosphere, *Science*, 326 (5959), 1525-1529, 2009.
- Leck, C., Svensson, E.: Importance of aerosol composition and mixing state for cloud droplet activation over the Arctic pack ice in summer, *Atmos. Chem. Phys.*, 15 (5), 2545-2568, 2015.

- Li, W., Sun, J., Xu, L., Shi, Z., Riemer, N., Sun, Y., Fu, P., Zhang, J., Lin, Y., Wang, X., Shao, L., Chen, J., Zhang, X., Wang, Z., Wang, W.: A conceptual framework for mixing structures in individual aerosol particles, *Journal of Geophysical Research: Atmospheres*, 121 (22), 13,784-713,798, 2016.
- Moffet, R.C., Rödel, T.C., Kelly, S.T., Yu, X.Y., Carroll, G.T., Fast, J., Zaveri, R.A., Laskin, A., Gilles, M.K.: Spectro-microscopic measurements of carbonaceous aerosol aging in Central California, *Atmos. Chem. Phys.*, 13 (20), 10445-10459, 2013.
- Peters, G.P., Nilssen, T.B., Lindholt, L., Eide, M.S., Glomsrød, S., Eide, L.I., Fuglestad, J.S.: Future emissions from shipping and petroleum activities in the Arctic, *Atmos. Chem. Phys.*, 11 (11), 5305-5320, 2011.
- Riemer, N., West, M., Zaveri, R.A., Easter, R.C.: Simulating the evolution of soot mixing state with a particle resolved aerosol model, *J. Geophys. Res.*, 114, doi:10.1029/2008JD011073, 2009.
- Samset, B.H., Myhre, G., Herber, A., Kondo, Y., Li, S.M., Moteki, N., Koike, M., Oshima, N., Schwarz, J.P., Balkanski, Y., Bauer, S.E., Bellouin, N., Bernsten, T.K., Bian, H., Chin, M., Diehl, T., Easter, R.C., Ghan, S.J., Iversen, T., Kirkevåg, A., Lamarque, J.F., Lin, G., Liu, X., Penner, J.E., Schulz, M., Seland, Ø., Skeie, R.B., Stier, P., Takemura, T., Tsigaridis, K., Zhang, K.: Modelled black carbon radiative forcing and atmospheric lifetime in AeroCom Phase II constrained by aircraft observations, *Atmos. Chem. Phys.*, 14 (22), 12465-12477, 2014.
- Seinfeld, J., Pandis, S., 2006. *Atmospheric Chemistry and Physics: From air pollution to climate change* (2nd ed.). 1-1203 pp., John Wiley & Son, Inc., Hoboken, New Jersey.
- Sierau, B., Chang, R.Y.W., Leck, C., Paatero, J., Lohmann, U.: Single-particle characterization of the high-Arctic summertime aerosol, *Atmos. Chem. Phys.*, 14 (14), 7409-7430, 2014.
- Wang, Y., Liu, F., He, C., Bi, L., Cheng, T., Wang, Z., Zhang, H., Zhang, X., Shi, Z., Li, W.: Fractal Dimensions and Mixing Structures of Soot Particles during Atmospheric Processing, *Environ. Sci. Tech. Lett.*, 4 (11), 487-493, 2017.
- You, Y., Renbaum-Wolff, L., Carreras-Sospedra, M., Hanna, S.J., Hiranuma, N., Kamal, S., Smith, M.L., Zhang, X., Weber, R.J., Shilling, J.E., Dabdub, D., Martin, S.T., Bertram, A.K.: Images reveal that atmospheric particles can undergo liquid-liquid phase separations, *Proceedings of the National Academy of Sciences of the United States of America* 109 (33), 13188-13193, 2012.
- Zanatta, M., Laj, P., Gysel, M., Baltensperger, U., Vratolis, S., Eleftheriadis, K., Kondo, Y., Dubuisson, P., Winiarek, V., Kazadzis, S., Tunved, P., Jacobi, H.W.: Effects of mixing state on optical and radiative properties of black carbon in the European Arctic, *Atmos. Chem. Phys.*, 18 (19), 14037-14057, 2018.
- Jiang, H., A. L. Frie, A. Lavi, J. Y. Chen, H. Zhang, R. Bahreini, and Y.-H. Lin (2019), Brown Carbon Formation from Nighttime Chemistry of Unsaturated Heterocyclic Volatile Organic Compounds, *Environmental Science & Technology Letters*, DOI: 10.1021/acs.estlett.1029b00017.
- Yang, Y., Wang, H., Smith, S.J., Easter, R.C., Rasch, P.J.: Sulfate Aerosol in the Arctic: Source Attribution and Radiative Forcing, *Journal of Geophysical Research: Atmospheres*, 123 (3), 1899-1918, 2018.

24 **Abstract**

25 Interaction of anthropogenic particles with radiation and clouds plays an important
26 role on Arctic climate change. Mixing state of aerosols is a key parameter to influence
27 aerosol-cloud and aerosol-radiation interaction. However, little is known on this
28 parameter in the Arctic, preventing an accurate representation of this information in
29 global models. Here we used transmission electron microscopy with
30 energy-dispersive X-ray spectrometry (TEM/EDS), scanning TEM, scanning electron
31 microscopy (SEM), nanoscale secondary ion mass spectrometry (NanoSIMS), and
32 atomic forces microscopy (AFM) to determine the size and mixing properties of
33 individual particles at 100 nm – 10 μ m, with a particular focus on sulfate and
34 carbonaceous particles. We found that non-sea salt sulfate particles with size range at
35 100-2000 nm were commonly coated with organic matter (OM) in summer. 20% of
36 sulfate particles also had soot inclusions which only appeared in the OM coating. The
37 OM coating is estimated to contribute to 63% of the particle volume on average. **To**
38 **understand how OM coating influences optical properties of sulfate particles, the Mie**
39 **theory of the core-shell model was applied to calculate optical properties of individual**
40 **sulfate particles. The result shows** that absorption cross section (ACS) of individual
41 OM-coated particles significantly increased **when assuming the OM coating as**
42 **light-absorbing brown carbon (BrC) and the ACS also increased following the**
43 **increasing particle size.** The microscopic observations suggest that OM modulates the
44 mixing structure of fine Arctic sulfate particles, which may determine their
45 hygroscopicity and optical properties.

46

47 **1. Introduction**

48 Surface temperatures are rising faster in the Arctic than the rest of globe (IPCC,
49 2013). Although increased human-induced emissions of long-lived greenhouse gases
50 are certainly one of the driving factors, air pollutants, such as aerosols and ozone, are
51 also important contributors to climate change in the Arctic (Law and Stohl, 2007;
52 Shindell, 2007). It is well known that aerosols from northern mid-altitude continents
53 affect the sea ice albedo by altering the heat balance of the atmosphere and surface
54 (Hansen and Nazarenko, 2004; Jacob et al., 2010; Shindell, 2007). These aerosols in
55 Arctic atmosphere include sea salt, sulfate, particulate organic matter (OM), and to a
56 lesser extent, ammonium, nitrate, black carbon (BC) (Hara et al., 2003; Quinn et al.,
57 2007) and mineral dust particles (Dagsson-Waldhauserova et al., 2013). Studies show
58 BC in the Arctic absorbs solar radiation in the atmosphere and when deposited on
59 snow (Iziomon et al., 2006; Koch and Hansen, 2005; Sand et al., 2013; Shindell,
60 2007). Moreover, Maahn et al. (2017) used aircraft in situ observation of clouds and
61 aerosols and found that concentration of BC are enhanced below the clouds in the
62 Arctic and further influence the mean effective radii of cloud droplets which lead to
63 the suppressed drizzle production and precipitation.

64 BC, commonly called “soot” is derived from the combustion sources such as
65 diesel engines, residential solid fuel, and open burning (Bond et al., 2013). Some
66 studies investigated the possible sources of these BC particles, including natural gas
67 flaring (Qi et al., 2017) and ship emissions in the Arctic (Browse et al., 2013;
68 Weinbruch et al., 2012) and emissions of biomass burning and fossil fuels in the
69 northern hemisphere (Winiger et al., 2016; Xu et al., 2017). For example, Winiger et
70 al.(2017) showed that most Arctic BC is sourced from domestic activities (35%) and
71 transportation (38%), with only minor contributions from gas flaring (6%), power
72 plants (9%), and open fires (12%).

73 Accumulation of secondary organic aerosols, a significant fraction of the new
74 particles grows to sizes that are active in cloud droplet formation in the Arctic (Abbatt
75 et al., 2019). More than 100 organic species have been detected in the Arctic aerosols
76 and polyacids are the most abundant compound class, followed by phthalates,

77 aromatic acids, fatty acids, fatty alcohols, sugars/sugar alcohols, and n-alkanes (Fu et
78 al., 2008). Recently, certain organic aerosols, referred to as brown carbon (BrC), have
79 been recognized as an important light-absorbing carbonaceous aerosol after BC in the
80 troposphere (Alexander et al., 2008; Andreae and Gelencser, 2006; Feng et al., 2013;
81 Lack et al., 2012). BrC can be directly emitted from combustion sources or formed in
82 the atmosphere via photo-chemical aging (Jiang et al., 2019; Saleh et al., 2013;
83 Updyke et al., 2012). **Moreover**, aging of secondary organic aerosols can significantly
84 contribute to BrC during atmospheric transports (Laskin et al., 2015). Feng et al.(2013)
85 estimated that on average, BrC **accounts** for 66% of total OM mass globally and its
86 light absorption is about 26% of BC.

87 BC and BrC are often internally mixed with other non-absorbing aerosols, such as
88 sulfate (Lack et al., 2012; Laskin et al., 2015). **Internal mixing means that a single**
89 **particle simultaneously contains two or more types of aerosol components (Li et al.,**
90 **2016)**. This internal mixing can enhance BC absorption by a factor of up to two (Bond
91 et al., 2013) and change the activity of cloud condensation nuclei (CCN) in the Arctic
92 atmosphere (Leck and Svensson, 2015; Martin et al., 2011). Spatial and temporal
93 variations of aerosol composition, size distribution, and sources of Arctic aerosols
94 **have been** studied extensively in numerous ground-based, ship, airborne observations,
95 **and various atmospheric models** (Brock et al., 2011; Burkart et al., 2017; Chang et al.,
96 2011; Dall'Osto et al., 2017; Fu et al., 2008; Hara et al., 2003; Hegg et al., 2010;
97 Iziomon et al., 2006; Karl et al., 2013; Latham et al., 2013; Leck and Bigg, 2008;
98 Leck and Svensson, 2015; Moore et al., 2011; Raatikainen et al., 2015;
99 Wörnschimmel et al., 2013; Winiger et al., 2017; Yang et al., 2018; Zangrando et al.,
100 2013). A few previous studies also looked at the mixing states of coarse aerosol
101 particles in Arctic troposphere (Behrenfeldt et al., 2008; Chi et al., 2015; Geng et al.,
102 2010; Hara et al., 2003; Leck and Svensson, 2015; Moroni et al., 2017; Raatikainen et
103 al., 2015; Sierau et al., 2014), but those of fine non-sea salt particles, including the
104 most important short-lived climate forcers – BC and BrC (Feng et al., 2013; Fu et al.,
105 2008; Kirpes et al., 2018; Laskin et al., 2015; Leck and Svensson, 2015), are poorly
106 characterized. **The poor understanding on mixing state of BC and BrC in individual**

107 particles will prevent the further simulation of atmospheric climate and aerosol-cloud
108 interaction in the Arctic through the current atmospheric models (Browse et al., 2013;
109 Samset et al., 2014; Zanatta et al., 2018).

110 In this study, individual aerosol particles were collected in the Arctic during 7-23
111 August, 2012. We combined the data from various microscopic instruments to
112 systematically determine the size, composition, and mixing properties of individual
113 particles, with a particular focus on sulfate and carbonaceous particles. Mie theory
114 was used to test how OM coating influences optical properties of sulfate particles in
115 the Arctic when OM was assumed as BrC. The results are discussed in the context of
116 aerosol-radiation and cloud interaction.

117

118 **2. Experimental section**

119 **2.1 Field campaign**

120 The Svalbard archipelago includes all landmasses between 74 and 81 degrees
121 North and 10 and 35 degrees East (Figure 1). The islands cover 63000 km².
122 Ny-Ålesund town is situated on the west coast of the largest island, Spitsbergen.
123 Ny-Ålesund town is situated only 1200 km from the North Pole and represents a
124 central platform for Arctic research. The sampling place represents remote Arctic
125 conditions.

126 An individual particle sampler at Chinese Arctic Yellow River Station (78°55'N,
127 11°56'E) collected individual particles (Chi et al., 2015; Geng et al., 2010). The
128 sampling site is about 2 km far away from the Zeppelin observatory station (78.9N
129 11.88E) running by the Ny-Ålesund Science Managers Committee
130 (<https://www.esrl.noaa.gov/psd/iasoa/stations/nyalesund>). Two to three samples were
131 regularly collected at 9:00, 16:00, 21:00 (local time) of each day, with a total of 46
132 samples during 7-23 August, 2012.

133 A sampler containing a single-stage impactor with a 0.5-mm-diameter jet nozzle
134 (Genstar Electronic Technology, China) was used to collect individual particles by the
135 air flow rate at 1.5 l min⁻¹. Aerosol particles were collected onto copper TEM grids
136 coated with carbon film. This sampler has a collection efficiency of 31% at 100 nm

137 aerodynamic diameter and 50% at 200 nm if the density of the particles is 2 g cm^{-3} .
138 The sampler can collect particles with $< 10 \text{ }\mu\text{m}$ aerodynamic diameter on TEM grids.
139 Sampling times varied from twenty minutes to two hours in clean remote Arctic area.
140 After collection, each sample was placed in a sealed dry plastic tube and stored in a
141 desiccator at $20 \pm 3\%$ RH for analysis. Ambient laboratory conditions (17–23% RH
142 and 19–21 °C) is effective at preserving individual hygroscopic aerosol particles and
143 reducing changes that would alter samples and subsequent data interpretation
144 (Laskina et al., 2015). The sample information such as local sampling date and time
145 and meteorological conditions (e.g., temperature (T), relative humidity (RH), pressure
146 (P), wind direction (WD), wind speed (WS)) are listed in Table S1.

147

148 2.2 TEM measurement

149 Individual particle samples were examined by a JEOL JEM-2100 transmission
150 electron microscopy operated at 200 kV with an energy-dispersive X-ray
151 spectrometry (TEM/EDS). TEM can observe the mixing structure of different aerosol
152 components within an individual particle on the substrate because electron beam
153 transmit through the specimen to form an image. EDS spectra were acquired for a
154 maximum time of 30 s to minimize potential beam damage and collect particle X-ray
155 spectra with sufficient intensity. TEM grids are made of copper (Cu) and covered by a
156 carbon-reinforced substrate, so Cu is excluded from the quantitative analyses of the
157 particles. Because of the substrate contribution, C content in TEM grid coated by
158 carbon film might be overestimated in EDS spectra of individual particles.

159 The distribution of aerosol particles on TEM grids was not uniform, with coarser
160 particles occurring near the center and finer particles on the periphery. Therefore, to
161 ensure that the analyzed particles are representative, five areas were chosen from the
162 center and periphery of the sampling spot on each grid. Through a labor-intensive
163 operation, 2002 aerosol particles with diameter $< 10 \text{ }\mu\text{m}$ in 21 samples were analyzed
164 by TEM/EDS (Table S1). To check elemental composition of individual particles,
165 EDX was manually used to obtain EDS spectra of individual particles. In the clean
166 Arctic air, there are simply particle types including sea salt, sulfate, soot, and OM.

167 Because soot particles have chain-like aggregation, it is not necessary to check their
168 elemental composition. Sea salt particles display spherical or square shapes and are
169 stable under the electron beam in TEM but sulfate particles are spherical but flats on
170 the substrate and produce unstable bubble under the electron beam (Buseck and Posfai,
171 1999; Chi et al., 2015). TEM observations also can clearly identify sulfate particles or
172 sulfate with OM coating. Therefore, we can easily identify Arctic particle types based
173 on their morphology. Because of the time-consuming in the experiment, it is not
174 necessary to frequently check elemental composition of the same particle type. For
175 the data statistic in this study, we randomly checked elemental composition of 20-30
176 particles in each sample (Table S1). EDS spectra of 575 particles were manually
177 selected and saved in the computer for elemental composition analysis. Particles
178 examined by TEM were dry at the time of observation in the vacuum of the electron
179 microscope. In our study, the effects of water and other semi-volatile organics were
180 not considered as they evaporate in the vacuum.

181 Elemental mapping and line profile of individual aerosol particles were obtained
182 from the EDX scanning operation mode of TEM (STEM). The STEM information can
183 clearly display elemental distribution in the targeted individual particles which cannot
184 be provided by the above EDS examination. Based on preliminary individual analysis,
185 we further chose the typical samples containing abundant sulfate with OM coating for
186 the STEM analysis. The high-resolution details of elemental distribution in individual
187 particles can further prove the details of the mixing structure of sulfate and OM in
188 individual particles.

189 The iTEM software (Olympus soft imaging solutions GmbH, Germany) is an
190 image analysis platform for electron microscopy. In this study, it was used to
191 manually or automatically obtain area, perimeter, and equivalent circle diameter
192 (ECD) of individual particles through identifying boundary of every particle in TEM
193 images. In these analyzed samples, we found there were abundant fine sulfate
194 particles in 11 samples collected during 9-15 August, 2012. In other samples, there
195 were only a few sulfate particles and more sea salt particles. Based on the TEM
196 observations, we selected the samples containing more sulfate particles to further do

197 other microscopic analyses as below.

198

199 **2.3 NanoSIMS measurement**

200 Because the sulfate particles collected in the Arctic had good consistent property
201 (e.g., elemental composition and mixing state) from TEM observations, we just
202 selected three samples containing abundant fine sulfate particles (Table S1) for further
203 studies. These three samples listed in Table S1 were analyzed using a nanoscale
204 secondary ion mass spectrometry (NanoSIMS) 50L (CAMECA Instruments,
205 Geneviers, France) instrument. A micro-cesium source was used to generate Cs^+
206 primary ions, with an impact energy of 16 kV for sample interrogation. The primary
207 beam was stepped across the sample to produce element specific, quantitative digital
208 images. The Cs^+ primary ion beam was used to obtain $^{16}\text{O}^-$, $^{12}\text{C}^{14}\text{N}^-$, $^{14}\text{N}^{16}\text{O}^-$, $^{32}\text{S}^-$,
209 $^{35}\text{Cl}^-$, and $^{16}\text{O}^{23}\text{Na}^-$ ions in this study. The NanoSIMS analysis can obtain ion mapping
210 of particles with nanometer spatial resolution over a broad range of particle sizes
211 (Figure S1). Because the substrate of TEM grid is carbon, CN^- is adopted to represent
212 OM in individual particles (Chi et al., 2015; Ghosal et al., 2014). S^- is used to infer the
213 presence of sulfates in individual particles (Li et al., 2017). Finally, the NanoSIMS
214 obtained ion mapping of 32 sulfate particles.

215

216 **2.4 SEM and AFM measurement**

217 Because TEM could not vertically observe OM coating and sulfate core, we
218 conducted one special experiment using a Zeiss ultra 55 scanning electron microscopy
219 (SEM) with EDS. The TEM grids were mounted onto an aluminum SEM stub and
220 directly observed in secondary electron image mode. SEM analysis was operated at
221 10 kv of extra high tension (EHT) and 9.7 mm of work distance (WD). Processes such
222 as sample moving, analysis region selection and imaging were controlled by computer.
223 The specimen stage in SEM was tilted at the range of 0-75°, and then we vertically
224 observed thickness of OM coating and sulfate core on the substrate. To verify vertical
225 property of individual S-rich particles impacting on the substrate, we observed two
226 typical samples containing abundant sulfate particles using the SEM (Table S1).

227 AFM with a digital nanoscope IIIa instrument operating in the tapping mode was
228 used to observe surface morphology of individual aerosol particles and measure
229 particle thickness. The tapping AFM has a cantilever and conical tip of 10 nm radius.
230 By using AFM, a general image of the particles is taken at 10 μm full scan size, which
231 generally includes 1-2 particles depending on the exact location. In this study, we are
232 only interested in the sulfate-containing particles. AFM provides surface information
233 and morphology of 17 particles but no composition. Samples were firstly quickly
234 examined by the TEM under low magnification mode. In case, the operation roughly
235 identified S-containing particles and didn't damage the secondary sulfate particles
236 under the electron. Because TEM grids have coordinates letters, we can exactly find
237 the same particles on the substrate in AFM examined in TEM experiments. The
238 procedures can exclude sea salt particles in the AFM image. As a result, the same
239 samples observed by TEM were then examined in AFM to obtain 3-D image of
240 secondary sulfate particles and their volume. Because individual particles collected in
241 Arctic air were scattered on the substrate, we only obtained 17 effective data. After we
242 obtained AFM images of sulfate particles, the NanoScope analysis software can
243 automatically obtain bearing area (A) and bearing volume (V) of each analyzed
244 particle according to the following formula.

$$245 \quad A = \frac{4}{3} \pi r^2 = \frac{\pi d^2}{3} \rightarrow d = \sqrt{\frac{3A}{\pi}} \quad (1)$$

$$246 \quad V = \frac{4}{3} \pi r^3 = \frac{4}{3} \times \frac{\pi D^3}{8} \rightarrow D = \sqrt[3]{\frac{6V}{\pi}} \quad (2)$$

247 Where x is the equivalent circle diameter (ECD) and y is the equivalent spherical
248 diameter (ESD).

249 ECD of individual aerosol particles measured from the iTEM software can be
250 further converted into ESD. Based on these data, we estimate one good linear
251 correlation ($y=0.38x$) between ESD and ECD of sulfate particles impacting on the
252 substrate. The value was further used to correct all the analyzed particles in TEM
253 images (Chi et al., 2015).

254

255 **2.5 Calculation of BrC optical properties**

256 The refractive index used for the non-light-absorbing sulfate component was set to
257 $m=1.55$ at 550 nm (Seinfeld and Pandis, 2006). The refractive index of OM (as BrC)
258 is not known so we considered three scenarios: strongly absorbing ($1.65-0.03i$ at 550
259 nm), moderately absorbing ($1.65-0.003i$ at 550 nm), and non-absorbing OM (1.65 at
260 550 nm) (Feng et al., 2013). Although the refractive index has dependence on the
261 wavelength between 350-870 nm, we tried to select the 550 nm as a case to test how
262 OM coating influence sulfate particles in Arctic air.

263 BHCOAT Mie code by Bohren and Huffman (1983) was used to calculate the
264 optical properties, including scattering cross section (SCS), absorption cross section
265 (ACS), and single scattering albedo (SSA), assuming a core-shell structure. We firstly
266 calculated these parameters assuming a sulfate core and OM shell structure only
267 (ignoring some of the particles that contain soot core). Because the Mie code only can
268 calculate the core-shell structure or homogeneous models, we assume sulfate as a core
269 and OM as a shell in individual particle to build the core-shell model. Based on the
270 core-shell standard mode (Li et al., 2016), we can calculate optical properties of
271 individual internally mixed particles.

272 **2.6 Back trajectories of air masses and Lagrangian particle dispersion model**

273 Three-day (72 h) back trajectories of air masses were generated using a Hybrid
274 Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model at the Chinese
275 Arctic Yellow River Station during August 2012, at an altitude of 500m above sea
276 level (Figure 1). Most air masses originate in the Arctic Ocean, and are restricted to
277 this vast marine region during the sampling periods. Based on the TEM observations,
278 air masses from North America and Greenland brought abundant sulfate particles into
279 the sampling area in summertime.

280 In order to determine the particle origins, the lagrangian particle dispersion model
281 FLEXPART-WRF 3.1 (Brioude et al., 2013) was used. The FLEXPART-WRF model
282 is using meteorological parameters from WRF dynamical simulation. The domain
283 resolution is 50×50 km with 64 vertical levels. The FLEXPART-WRF simulations
284 were launched in a backward mode over 10 days, with the Chinese Arctic Yellow

285 River Station as an origin. For each simulation (one per sample), 20000
286 pseudo-particles were released in a small volume around the station position. Each
287 single particle position evolution backward in time was determined by Lagrangian
288 dispersion calculation. Based on the TEM experiments **and back trajectory of air**
289 **masses (Figure 1)**, we found that there were more S-rich with OM coating particles in
290 the samples collected on August 11, 12, 14 and 15, 2012. Therefore, we **further** did the
291 FLEXPART-WRF simulation of these four days **(Figure 2)**. The emission intensity in
292 the Arctic area has been also shown in Figure **S2**.

293 **3. Results**

294 **3.1 Composition and sources of aerosol particles**

295 We summarized average elemental weight and frequency of individual **Arctic**
296 particles derived from the TEM/EDX. The result shows that O, Na, S, and Cl in
297 individual particles are dominant elements (Figure S3). On basis of the composition
298 and morphology of individual particles, we classified the particles into four major
299 groups: Na-rich (i.e. NaCl, Na₂SO₄, and NaNO₃), S-rich (i.e. ammonium sulfate and
300 sulfuric acid), and carbonaceous (soot and OM). The classification criteria of different
301 particle types and their sources have been described in a separate study (Li et al.,
302 2016). S-rich particles representing secondary inorganic particles (e.g., SO₄²⁻, NO₃⁻,
303 and NH₄⁺) are transformed from gaseous SO₂, NO_x, and NH₃. OM can be divided into
304 primary organic matter (POM) and secondary organic matter (SOM). SOM is
305 produced from the chemical oxidation of volatile organic compounds (VOCs) and
306 **often** exhibits OM coating on S-rich particles. Na-rich particles in the marine air are
307 from sea spray and have typical near cubic shape. Soot particles, which contain C
308 with minor O, appear as a chain-like aggregate of carbon-bearing spheres. Our
309 previous study well characterized aging mechanism of sea salt particles in summer
310 Arctic air (Chi et al., 2015). **Here** we focused on S-rich, soot, and OM particles as the
311 major non-sea salt particle (NSS-particle, **39 ±5%**) in the analyzed samples, which are
312 approximately **29 ±7% of 2002** particles (Figure 3).

313

314 **3.2 OM coating on sulfate particles**

315 TEM observations revealed a common core-shell mixing structure in fine sulfate
316 particles (Figure 4a). Elemental mapping of such internally mixed sulfate particles
317 shows C signals in the coating (C map, Figure 4b) and S and O signals in the center (S
318 and O map, Figure 4c, d). The elemental line profile of a sulfate particle also shows
319 sulfate core and C coating (Figure S4). Furthermore, ion maps of individual particles
320 from the NanoSIMS further exhibit $^{12}\text{C}^{14}\text{N}^-$ signals in the coating (red color in Figure
321 4e, f) and $^{32}\text{S}^-$ signals in the core (green color in Figure 4e, g). These results provide
322 strong evidence that the coating is OM and the core is sulfate.

323 A majority of 781 analyzed NSS-particles (74% by particle number) have a sulfate
324 core and OM coating (Figures 4 and 5). ~20% of them also contain small soot
325 inclusions but they only appeared in organic coating, rather than as the core mixed in
326 sulfate (Figure 5b). The mixing structure is different from our previous findings in
327 polluted air that soot is normally mixed with sulfate instead of OM coating (Li et al.,
328 2016). Moreover, we noticed that a few chain-like soot aggregates (1.3% in all
329 analyzed particles) (Figure S5) only occurred in three samples during the sampling
330 period (Table S1). Considering the remoteness of the sampling site, such fresh soot
331 particles are likely to be of local origin, including shipping and flaring (Gilgen et al.,
332 2018; Peters et al., 2011). Indeed, we found a few of ships moving in Arctic Ocean
333 during these days from the Ny-Ålesund town.

334 TEM observations showed that some sulfate particles had unique morphology that
335 a sulfate particle was surrounded by some smaller particles (Figure 5a). They are
336 often called “satellite” particles as they were distributed from the central particles
337 when impacted on the substrate during sample collection. 16% of the analyzed sulfate
338 particles with satellite particles as shown in Figure 5a were detected in the samples
339 (Table S1) collected during 9-15 August. NanoSIMS analysis further provided more
340 information that the satellite particles selected from the samples (Table S1) have
341 strong $^{32}\text{S}^-$ (Figure 6a, c) and $^{16}\text{O}^-$ signals (Figure 6d) as well as weak $^{12}\text{C}^{14}\text{N}^-$ signals
342 (Figure 6a, b). The CN^- signal normally can represent organic aerosols (Chi et al.,
343 2015; Ghosal et al., 2014). Previous studies showed that the similar satellite particles
344 are normally considered as acidic sulfate (Buseck and Posfai, 1999; Iwasaka et al.,

1983). Therefore, we can conclude that these acidic satellites not only contain sulfuric acid but also some OM or organic acids. Indeed, Fu et al. (2008) found that polyacids are the most abundant organic compounds, followed by phthalates, aromatic acids, and fatty acids in Arctic aerosol particles. As a result, these Arctic sulfate particles with satellites contain certain amounts of sulfuric or organic acids with liquid phase. Back trajectories of air masses and FLEXPART both shows abundant sulfate particles and some containing satellite particles were transported from Greenland and North American (Figures 1 and 2).

AFM was used to obtain 3D image of individual secondary particles impacting on the substrate. Figure 7a shows that the secondary particles normally have smooth surface which is different from uneven surface of the Arctic fresh and aged NaCl particles (Chi et al., 2015). Furthermore, we observed particle thickness through tilting the specimen stage up to 75° in SEM. Figure 7a-b both shows that the secondary particles look like thin pancake sticking on the substrate. Furthermore, the sections of two secondary particles in the AFM images shows that the highest heights of particles are only 0.15 (green line) and 0.26 (red line) of the corresponding horizontal size (Figure 7a). Here we can conclude that shape of individual particles was modified when they impacted on the substrate following the airflow. Therefore, the measured ECDs of individual particles in TEM images are much larger than the real particle diameter. To calibrate the particle diameter, we obtained volume of dry particles on the substrate and then calculated their equivalent sphere diameter (ESD) in the AFM images (Figure 7c). ESD distribution of the secondary Arctic particles displayed a peak at 340 nm, ranging from 100 nm to 2000 nm (Figure 7d). The core particles, as sulfate or soot, had a peak at 240 nm and 120 nm, respectively (Figure 7d). In the core-shell particles, we knew size in all the analyzed particles and further calculated volume of sulfate, OM, and/or soot within individual particles. We can estimate that OM on average accounted for 63 ± 23% of the dry sulfate particle volume. Our result shows that the OM volume increases following the particle size increase (Figure S6).

374

375 4. Discussion

376 4.1 Mixing mechanism of organic, soot, and sulfate

377 Lagrangian particle dispersion modeling using the FLEXPART-WRF 3.1 showed
378 that air masses arriving at the sampling site during our field measurement periods
379 were likely originated from the Greenland and North America (Figure 2). Previous
380 studies reported that air masses from North America or Greenland during the summer
381 contain higher concentration of black carbon, OM, and sulfate (Burkart et al., 2017;
382 Chang et al., 2011; Fu et al., 2008; Moore et al., 2011; Park et al., 2013). Indeed, there
383 is strong emission intensity of OC and SO₂ around the Arctic area from emission
384 simulation as shown in Figure S2. However, Weinbruch et al. (2012) observed soot
385 particles when cruise ships were present in the area around Ny-Ålesund town. It is
386 possible that **minor** soot particles are sourced from the ship emissions **and most of**
387 **them are** transported from out of Arctic area in the free troposphere (Figure S2).

388 The sulfate core-OM shell structure observed in the Arctic summer atmosphere is
389 similar to those in the background or rural air in other places (Li et al., 2016; Moffet
390 et al., 2013). Based on the images from electron microscopies, we can infer that OM
391 coating thickness in the **Arctic air** was comparable with them in rural places **but**
392 **higher than them in urban places**. During the transports, organic coatings on sulfates
393 were considered as the secondary organic aerosols and their masses increase
394 following particle aging and growth (Li et al., 2016; Moffet et al., 2013; Sierau et al.,
395 2014). **Figures 1 and 2 show that most of particles in the air masses transported long**
396 **distance from North American. The result indicates that these long-range**
397 **transportation of secondary sulfate particles have enough time to experience the**
398 **possible atmospheric heterogeneous reactions on particle surfaces or cloud processes**
399 **in the Arctic air. Similarly**, Moffet et al. (2013) found that soot inclusions occurred in
400 OM coating when OM coating on sulfates built up through photochemical activity and
401 pollution buildup the Sacramento urban plume aged. On the other hand, the
402 sulfate/OM particles with soot inclusions are probably formed in a similar way as
403 those found elsewhere (Li et al., 2016) – e.g., soot particles may have acted as nuclei
404 for secondary sulfate or organic uptake during their transports (Riemer et al., 2009).

405 Similarly, besides the OM coating in the Arctic particles, Leck and Svensson (2015)
406 found some biogenic aerosols like gel-aggregate containing bacterium in ultrafine
407 particles. However, we didn't find any gel-like particles in the samples because our
408 sampler had very low efficiency for ultrafine particles.

409 TEM images show that most of the internally mixed sulfate particles display
410 sulfate core and OM coating on the substrate (Figures 4a and 5b, c). The sulfate and
411 OM separation in individual particles were defined by You et al. (2012) as
412 liquid-liquid phase separation (LLPS). Concerning the knowledges of the LLPS can
413 better understand particle hygroscopicity, heterogeneous reactions of reactive gases on
414 particle surface, and organic aging (You et al., 2012). They also reported that the
415 LLPS can reflect the O:C ratio in the OM, which is roughly ≤ 0.5 . In this study, we
416 did observe the LLPS in almost all the fine sulfate particles, which indicates that the
417 secondary OM in the coating might be not highly aged. Therefore, we speculate that
418 the thick OM coatings were consistently built up during the long-range transport of
419 sulfate particles and part of secondary OM in the coating likely formed in Arctic area.
420 Indeed, some studies reported that there are various sources of organic precursors
421 during the Arctic area, such as biogenic VOCs from ice melting and open water
422 (Dall'Osto et al., 2017) and anthropogenic VOCs from shipping emissions in
423 summertime (Gilgen et al., 2018). The dependence of OM volume on particle size
424 (Figure S6) suggests that the suspended sulfate particles are initially important surface
425 for secondary OM formation. Moreover, the common OM coating on sulfate particles
426 indicates that secondary OM as the surfaces of fine particles might govern the
427 possible heterogeneous reactions between reactive gases and sulfate particles in the
428 Arctic air.

429 It should be noted that most of secondary OM not only occurred on the surfaces
430 of sulfate particles but also its mass (mean mass at $63 \pm 23\%$) dominated in individual
431 particles (Figure 7d). The OM dominating in individual particles can influence the IN
432 and CCN activities of secondary sulfate particles (Latham et al., 2013; Martin et al.,
433 2011). For example, some studies found that an increase in organic mass fraction in
434 particles of a certain size would lead to a suppression of the Arctic CCN activity

435 (Leck and Svensson, 2015; Martin et al., 2011). Moreover, OM as particle surfaces
436 can significantly influence **hygroscopicity and IN activity** of sulfate particles (Wang et
437 al., 2012).

438

439 **4.2 Potential impact of OM on optical properties of sulfate-containing particles**

440 The internal mixing of soot, sulfate, and OM can change optical properties of
441 individual particles in the atmosphere. Recent studies showed that BrC has been
442 detected in the OM in the polluted and clean air and even in upper troposphere
443 (Laskin et al., 2015; Wang et al., 2018). Feng et al. (2013) further calculated the
444 contribution up to 19% of the optical absorption of the strongly absorbing BrC in
445 global simulations which is after the absorption BC aerosols. Although we didn't
446 directly measure the optical absorption and BrC in the Arctic atmosphere, various
447 colored OM (e.g. nitrated/polycyclic aromatics and phenols), referred as BrC, were
448 detected in the Arctic atmosphere in different seasons (Fu et al., 2008;
449 Wöhrensimmel et al., 2013; Zangrando et al., 2013) and in surface ice or snowpack
450 (Browse et al., 2013; Doherty et al., 2013; Hegg et al., 2010). We also noticed that the
451 $^{12}\text{C}^{14}\text{N}^-$ signal generally occurred in all analyzed OM coating in sulfate particles
452 (Figure 4e-f). Herrmann et al. (2007) **considered that $^{12}\text{C}^{14}\text{N}^-$ from NanoSIMS**
453 **represents nitrogen-containing organic in the detected materials. In this study,**
454 **although we could not determine that all the organic materials in the OM coating were**
455 **nitrogen-containing OM, the NanoSIMS data as shown in Figure 4 indicated that the**
456 **OM coating more or less homogenously contained nitrogen-containing OM. As a**
457 **result,** the nitrogen-containing OM indicates that the OM coating could contain
458 certain amounts of secondary BrC (Jiang et al., 2019; Laskin et al., 2015).

459 To understand **how OM coating influence optical properties of sulfate particles,**
460 **we assume three scenarios of OM coating as BrC:** strongly absorbing (case 1),
461 moderately absorbing (case 2) or non-absorbing OM (case 3) with a refractive index
462 of 1.65-0.03i, 1.65-0.003i, and 1.65 at 550 nm according to *Feng et al.* (2013). Based
463 on the size measurements shown in Figure 7d, we can calculate volume of sulfate and
464 OM within each particle. We input volume of each component and the corresponding

465 refractive index into the Mie code and then calculated optical properties of individual
466 sulfate particles in the samples. Based on optical data statistic of 575 particles, Figure
467 8a show that the OM coating is strongly absorbing BrC (referred to case Abs1), as by
468 Feng et al.(2013), the average absorption cross section (ACS) of individual particles is
469 estimated to be $2.67 \times 10^{-14} \text{ m}^2$. This value is 8.30 times higher than the aerosol ACS
470 ($3.22 \times 10^{-15} \text{ m}^2$) when assuming that the BrC is moderately absorbing (referred to case
471 Abs2, Figure 8a). However, the scattering cross section (SCS) of individual particles
472 only shows a small change (Figure 8b). Figure 8c also shows that the single scattering
473 albedos (SSAs) of individual particles are 0.92, 0.99, and 1 when assuming the BrC as
474 strongly, moderately and non-absorbing (cases SSA1 to SSA3). These results suggest
475 whether we consider organic coating as BrC may have a significant influence on the
476 absorption properties of individual sulfate particles.

477 In this study, we explored the relationship between ACS of individual particles and
478 particle diameters. Interestingly, Figure 8d shows that ACS of individual fine
479 OM-coating sulfate particles increased following the increasing particle size. The
480 result shows that the ACS can be enhanced following particle size growing and
481 particle aging. In other word, OM-coating sulfate particles transported more longer
482 distances and they might have stronger optical absorption in the Arctic air.

483 Current climate models estimated the radiative force of Arctic BC (Sand et al.,
484 2013; Shindell, 2007; Winiger et al., 2017; Zanatta et al., 2018), but none specifically
485 considered optical properties of Arctic BrC. Our study well revealed OM coating on
486 sulfate particles and this detail microphysical complexity of aerosol particles will be
487 useful to construct the atmospheric radiation and CCN/IN simulation in Arctic
488 atmospheric models in the future.

489

490 5 Summary

491 Different individual particle techniques, such as TEM/EDS, STEM, SEM,
492 NanoSIMS, and AFM, were applied to study S-rich, soot, and OM particles in the
493 Arctic air in summer. Sulfate particles accounted for approximately $29 \pm 7\%$ by
494 number of all analyzed particles in Arctic air. TEM and NanoSIMS commonly

495 observed OM coating and sulfate core individual sulfate particles, defined as the
496 LLSP. The common OM coating on sulfate particles indicates that secondary OM as
497 the surfaces of fine particles might govern the possible heterogeneous reactions
498 between reactive gases and sulfate particles in the Arctic air. Moreover, 20% of them
499 also contain small soot inclusions but they only appeared in organic coating, rather
500 than as the core mixed in sulfate. The mixing structure is totally different from the
501 previous findings that soot is internally mixed with sulfate instead of OM coating in
502 urban polluted air.

503 Size distribution of the secondary Arctic particles displayed a peak at 340 nm,
504 ranging from 100 nm to 2000 nm. The core particles, as sulfate or soot, had a peak at
505 240 nm and 120 nm, respectively. Furthermore, we can estimate that OM on average
506 accounted for $63 \pm 23\%$ of the dry NSS-particle volume. Based on microscopic
507 measurements of individual particles, we not only built up one core-shell model but
508 also quantify volume of OM and sulfate in individual particles. The Mie code was
509 used to calculate optical properties of internally mixed sulfate/OM particles when we
510 considered OM as non-absorbing, moderately absorbing BrC, and strongly absorbing
511 BrC. We found that the aerosol ACS is 8.30 times higher than the BrC as moderately
512 absorbing. We concluded that whether we consider organic coating as BrC may have a
513 significant influence on the absorption properties of individual particles in the Arctic
514 air. Moreover, individual fine OM-coating sulfate particles increased following the
515 increasing particle size. Therefore, we proposed that further studies should focus on
516 the BrC in Arctic aerosols: What mass concentrations of BrC are in fine particles?
517 What kinds of BrC are in fine particles? The optical mass absorption of BrC in fine
518 particles should be investigated? These results can be used to evaluate how BrC
519 aerosols influence the Arctic climate.

520

521 **Author Contributions:** WL and ZS designed the study. YZ and XS collected aerosol
522 particles. WL, HY, and JZ contributed laboratory experiments and data analysis. HY
523 and WL performed optical calculation and wrote part of first draft. PT and MD
524 provided the online measurement data of new particle formation and growth. JS and
525 XZ coordinated the field campaign. All authors commented and edited the paper.

526

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528

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References

- Abbatt, J.P.D., Leaitch, W.R., Aliabadi, A.A., Bertram, A.K., Blanchet, J.P., Boivin-Rioux, A., Bozem, H., Burkart, J., Chang, R.Y.W., Charette, J., Chaubey, J.P., Christensen, R.J., Cirisan, A., Collins, D.B., Croft, B., Dionne, J., Evans, G.J., Fletcher, C.G., Gal í M., Ghahremaninezhad, R., Girard, E., Gong, W., Gosselin, M., Gourdal, M., Hanna, S.J., Hayashida, H., Herber, A.B., Hesaraki, S., Hoor, P., Huang, L., Hussherr, R., Irish, V.E., Keita, S.A., Kodros, J.K., Köllner, F., Kolonjari, F., Kunkel, D., Ladino, L.A., Law, K., Levasseur, M., Libois, Q., Liggio, J., Lizotte, M., Macdonald, K.M., Mahmood, R., Martin, R.V., Mason, R.H., Miller, L.A., Moravek, A., Mortenson, E., Mungall, E.L., Murphy, J.G., Namazi, M., Norman, A.L., O'Neill, N.T., Pierce, J.R., Russell, L.M., Schneider, J., Schulz, H., Sharma, S., Si, M., Staebler, R.M., Steiner, N.S., Thomas, J.L., von Salzen, K., Wentzell, J.J.B., Willis, M.D., Wentworth, G.R., Xu, J.W., Yakobi-Hancock, J.D.: Overview paper: New insights into aerosol and climate in the Arctic, *Atmos. Chem. Phys.*, 19 (4), 2527-2560, 2019.
- Alexander, D.T.L., Crozier, P.A., Anderson, J.R.: Brown Carbon Spheres in East Asian Outflow and Their Optical Properties, *Science*, 321 (5890), 833-836, 2008.
- Andreae, M.O., Gelencser, A.: Black carbon or brown carbon? The nature of light-absorbing carbonaceous aerosols, *Atmos. Chem. Phys.*, 6 (10), 3131-3148, 2006.
- Behrenfeldt, U., Krejci, R., Ström, J., Stohl, A.: Chemical properties of Arctic aerosol particles collected at the Zeppelin station during the aerosol transition period in May and June of 2004, *Tellus B*, 60 (3), 405-415, 2008.
- Bohren, C.F., Huffman, D.R., 1983. Absorption and scattering of light by small particles. John Wiley & Sons, Inc. , New York, USA.
- Bond, T.C., Doherty, S.J., Fahey, D.W., Forster, P.M., Berntsen, T., DeAngelo, B.J., Flanner, M.G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P.K., Sarofim, M.C., Schultz, M.G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N., Guttikunda, S.K., Hopke, P.K., Jacobson, M.Z., Kaiser, J.W., Klimont, Z., Lohmann, U., Schwarz, J.P., Shindell, D., Storelvmo, T., Warren, S.G., Zender, C.S.: Bounding the role of black carbon in the climate system: A scientific assessment, *J. Geophys. Res.*, 118 (11), 5380-5552, 2013.
- Brioude, J., Arnold, D., Stohl, A., Cassiani, M., Morton, D., Seibert, P., Angevine, W., Evan, S., Dingwell, A., Fast, J.D., Easter, R.C., Pisso, I., Burkhart, J., Wotawa, G.: The Lagrangian particle dispersion model FLEXPART-WRF version 3.1, *Geosci. Model Dev.*, 6 (6), 1889-1904, 2013.
- Brock, C.A., Cozic, J., Bahreini, R., Froyd, K.D., Middlebrook, A.M., McComiskey, A., Brioude, J., Cooper, O.R., Stohl, A., Aikin, K.C., de Gouw, J.A., Fahey, D.W., Ferrare, R.A., Gao, R.S., Gore, W., Holloway, J.S., Hübler, G., Jefferson, A., Lack, D.A., Lance, S., Moore, R.H., Murphy, D.M., Nenes, A., Novelli, P.C., Nowak, J.B., Ogren, J.A., Peischl, J., Pierce, R.B., Pilewskie, P., Quinn, P.K., Ryerson, T.B., Schmidt, K.S., Schwarz, J.P., Sodemann, H., Spackman, J.R., Stark, H., Thomson, D.S., Thornberry, T., Veres, P., Watts, L.A., Warneke, C., Wollny, A.G.: Characteristics, sources, and transport of aerosols measured in spring 2008 during the aerosol, radiation, and cloud processes affecting Arctic Climate (ARCPAC) Project, *Atmos. Chem. Phys.*, 11 (6), 2423-2453, 2011.
- Browse, J., Carslaw, K.S., Schmidt, A., Corbett, J.J.: Impact of future Arctic shipping on high-latitude black carbon deposition, *Geophys. Res. Lett.*, 40 (16), 4459-4463, 2013.
- Burkart, J., Willis, M.D., Bozem, H., Thomas, J.L., Law, K., Hoor, P., Aliabadi, A.A., Köllner, F., Schneider, J., Herber, A., Abbatt, J.P.D., Leaitch, W.R.: Summertime observations of elevated levels of ultrafine particles in the high Arctic marine boundary layer, *Atmos. Chem. Phys.*, 17 (8),

- 5515-5535, 2017.
- Buseck, P.R., Posfai, M.: Airborne minerals and related aerosol particles: Effects on climate and the environment, *P. Natl. Acad. Sci. USA*, 96 (7), 3372-3379, 1999.
- Chang, R.Y.W., Leck, C., Graus, M., Müller, M., Paatero, J., Burkhardt, J.F., Stohl, A., Orr, L.H., Hayden, K., Li, S.M., Hansel, A., Tjernström, M., Leaitch, W.R., Abbatt, J.P.D.: Aerosol composition and sources in the central Arctic Ocean during ASCOS, *Atmos. Chem. Phys.*, 11 (20), 10619-10636, 2011.
- Chi, J.W., Li, W.J., Zhang, D.Z., Zhang, J.C., Lin, Y.T., Shen, X.J., Sun, J.Y., Chen, J.M., Zhang, X.Y., Zhang, Y.M., Wang, W.X.: Sea salt aerosols as a reactive surface for inorganic and organic acidic gases in the Arctic troposphere, *Atmos. Chem. Phys.*, 15 (19), 11341-11353, 2015.
- Dagsson-Waldhauserova, P., Arnalds, O., Olafsson, H.: Long-term frequency and characteristics of dust storm events in Northeast Iceland (1949–2011), *Atmos. Environ.*, 77 (0), 117-127, 2013.
- Dall'Osto, M., Beddows, D.C.S., Tunved, P., Krejci, R., Ström, J., Hansson, H.C., Yoon, Y.J., Park, K.-T., Becagli, S., Udisti, R., Onasch, T., O'Dowd, C.D., Simó, R., Harrison, R.M.: Arctic sea ice melt leads to atmospheric new particle formation, *Sci. Rep.*, 7 (1), 3318, 2017.
- Doherty, S.J., Grenfell, T.C., Forsström, S., Hegg, D.L., Brandt, R.E., Warren, S.G.: Observed vertical redistribution of black carbon and other insoluble light-absorbing particles in melting snow, *J. Geophys. Res.*, 118 (11), 5553-5569, 2013.
- Feng, Y., Ramanathan, V., Kotamarthi, V.R.: Brown carbon: a significant atmospheric absorber of solar radiation?, *Atmos. Chem. Phys.*, 13 (17), 8607-8621, 2013.
- Fu, P., Kawamura, K., Barrie, L.A.: Photochemical and Other Sources of Organic Compounds in the Canadian High Arctic Aerosol Pollution during Winter–Spring, *Environ. Sci. Technol.*, 43 (2), 286-292, 2008.
- Geng, H., Ryu, J., Jung, H.-J., Chung, H., Ahn, K.-H., Ro, C.-U.: Single-Particle Characterization of Summertime Arctic Aerosols Collected at Ny-Alesund, Svalbard, *Environ. Sci. Technol.*, 44 (7), 2348-2353, 2010.
- Ghosal, S., Weber, P.K., Laskin, A.: Spatially resolved chemical imaging of individual atmospheric particles using nanoscale imaging mass spectrometry: insight into particle origin and chemistry, *Analytical Methods*, 6 (8), 2444-2451, 2014.
- Gilgen, A., Huang, W.T.K., Ickes, L., Neubauer, D., Lohmann, U.: How important are future marine and shipping aerosol emissions in a warming Arctic summer and autumn?, *Atmos. Chem. Phys.*, 18 (14), 10521-10555, 2018.
- Hansen, J., Nazarenko, L.: Soot climate forcing via snow and ice albedos, *P. Natl. Acad. Sci. USA*, 101 (2), 423-428, 2004.
- Hara, K., Yamagata, S., Yamanouchi, T., Sato, K., Herber, A., Iwasaka, Y., Nagatani, M., Nakata, H.: Mixing states of individual aerosol particles in spring Arctic troposphere during ASTAR 2000 campaign, *J. Geophys. Res.*, 108 (D7), 2003.
- Hegg, D.A., Warren, S.G., Grenfell, T.C., Sarah, J.D., Clarke, A.D.: Sources of light-absorbing aerosol in arctic snow and their seasonal variation, *Atmos. Chem. Phys.*, 10 (22), 10923-10938, 2010.
- Herrmann, A.M., Ritz, K., Nunan, N., Clode, P.L., Pett-Ridge, J., Kilburn, M.R., Murphy, D.V., O'Donnell, A.G., Stockdale, E.A.: Nano-scale secondary ion mass spectrometry — A new analytical tool in biogeochemistry and soil ecology: A review article, *Soil Biol. Biochem.*, 39 (8), 1835-1850, 2007.
- IPCC (2013), Clouds and Aerosols, in *Climate Change 2013: The Physical Science Basis. Contribution*

- of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, 571-657 pp, Cambridge, U.K. and New York, NY.
- Iwasaka, Y., Minoura, H., Nagaya, K.: The transport and spatial scale of Asian dust-storm clouds: A case study of the dust-storm event of April 1979, *Tellus, Ser. B*, 35, 189-196, 1983.
- Iziomon, M.G., Lohmann, U., Quinn, P.K.: Summertime pollution events in the Arctic and potential implications, *J. Geophys. Res.*, 111 (D12), D12206, 2006.
- Jacob, D.J., Crawford, J.H., Maring, H., Clarke, A.D., Dibb, J.E., Emmons, L.K., Ferrare, R.A., Hostetler, C.A., Russell, P.B., Singh, H.B., Thompson, A.M., Shaw, G.E., McCauley, E., Pederson, J.R., Fisher, J.A.: The Arctic Research of the Composition of the Troposphere from Aircraft and Satellites (ARCTAS) mission: design, execution, and first results, *Atmos. Chem. Phys.*, 10 (11), 5191-5212, 2010.
- Jiang, H., Frie, A.L., Lavi, A., Chen, J.Y., Zhang, H., Bahreini, R., Lin, Y.-H.: Brown Carbon Formation from Nighttime Chemistry of Unsaturated Heterocyclic Volatile Organic Compounds, *Environ. Sci. Techn. Lett.*, DOI: 10.1021/acs.estlett.1029b00017, 2019.
- Karl, M., Leck, C., Coz, E., Heintzenberg, J.: Marine nanogels as a source of atmospheric nanoparticles in the high Arctic, *Geophys. Res. Lett.*, 40 (14), 3738-3743, 2013.
- Kirpes, R.M., Bondy, A.L., Bonanno, D., Moffet, R.C., Wang, B., Laskin, A., Ault, A.P., Pratt, K.A.: Secondary sulfate is internally mixed with sea spray aerosol and organic aerosol in the winter Arctic, *Atmos. Chem. Phys.*, 18 (6), 3937-3949, 2018.
- Koch, D., Hansen, J.: Distant origins of Arctic black carbon: A Goddard Institute for Space Studies ModelE experiment, *J. Geophys. Res.*, 110 (D4), D04204, 2005.
- Lack, D.A., Langridge, J.M., Bahreini, R., Cappa, C.D., Middlebrook, A.M., Schwarz, J.P.: Brown carbon and internal mixing in biomass burning particles, *P. Natl. Acad. Sci. USA*, 109 (37), 14802-14807, 2012.
- Laskin, A., Laskin, J., Nizkorodov, S.A.: Chemistry of Atmospheric Brown Carbon, *Chem. Rev.*, 115 (10), 4355-4382, 2015.
- Laskina, O., Morris, H.S., Grandquist, J.R., Estillore, A.D., Stone, E.A., Grassian, V.H., Tivanski, A.V.: Substrate-Deposited Sea Spray Aerosol Particles: Influence of Analytical Method, Substrate, and Storage Conditions on Particle Size, Phase, and Morphology, *Environ. Sci. Tech.*, 49 (22), 13447-13453, 2015.
- Latham, T.L., Beyersdorf, A.J., Thornhill, K.L., Winstead, E.L., Cubison, M.J., Hecobian, A., Jimenez, J.L., Weber, R.J., Anderson, B.E., Nenes, A.: Analysis of CCN activity of Arctic aerosol and Canadian biomass burning during summer 2008, *Atmos. Chem. Phys.*, 13 (5), 2735-2756, 2013.
- Law, K.S., Stohl, A.: Arctic Air Pollution: Origins and Impacts, *Science*, 315 (5818), 1537-1540, 2007.
- Leck, C., Bigg, E.K.: Comparison of sources and nature of the tropical aerosol with the summer high Arctic aerosol, *Tellus, Ser. B*, 60 (1), 118-126, 2008.
- Leck, C., Svensson, E.: Importance of aerosol composition and mixing state for cloud droplet activation over the Arctic pack ice in summer, *Atmos. Chem. Phys.*, 15 (5), 2545-2568, 2015.
- Li, W., Sun, J., Xu, L., Shi, Z., Riemer, N., Sun, Y., Fu, P., Zhang, J., Lin, Y., Wang, X., Shao, L., Chen, J., Zhang, X., Wang, Z., Wang, W.: A conceptual framework for mixing structures in individual aerosol particles, *J. Geophys. Res.*, 121 (22), 13784-713,798, 2016.
- Li, W., Xu, L., Liu, X., Zhang, J., Lin, Y., Yao, X., Gao, H., Zhang, D., Chen, J., Wang, W., Harrison, R.M., Zhang, X., Shao, L., Fu, P., Nenes, A., Shi, Z.: Air pollution-aerosol interactions produce more bioavailable iron for ocean ecosystems, *Sci. Adv.*, 3 (3), e1601749, 2017.

- Maahn, M., de Boer, G., Creamean, J.M., Feingold, G., McFarquhar, G.M., Wu, W., Mei, F.: The observed influence of local anthropogenic pollution on northern Alaskan cloud properties, *Atmos. Chem. Phys.*, 17 (23), 14709-14726, 2017.
- Martin, M., Chang, R.Y.W., Sierau, B., Sjogren, S., Swietlicki, E., Abbatt, J.P.D., Leck, C., Lohmann, U.: Cloud condensation nuclei closure study on summer arctic aerosol, *Atmos. Chem. Phys.*, 11 (22), 11335-11350, 2011.
- Moffet, R.C., Rödel, T.C., Kelly, S.T., Yu, X.Y., Carroll, G.T., Fast, J., Zaveri, R.A., Laskin, A., Gilles, M.K.: Spectro-microscopic measurements of carbonaceous aerosol aging in Central California, *Atmos. Chem. Phys.*, 13 (20), 10445-10459, 2013.
- Moore, R.H., Bahreini, R., Brock, C.A., Froyd, K.D., Cozic, J., Holloway, J.S., Middlebrook, A.M., Murphy, D.M., Nenes, A.: Hygroscopicity and composition of Alaskan Arctic CCN during April 2008, *Atmos. Chem. Phys.*, 11 (22), 11807-11825, 2011.
- Moroni, B., Cappelletti, D., Crocchianti, S., Becagli, S., Caiazzo, L., Traversi, R., Udusti, R., Mazzola, M., Markowicz, K., Ritter, C., Zielinski, T.: Morphochemical characteristics and mixing state of long range transported wildfire particles at Ny-Ålesund (Svalbard Islands), *Atmos. Environ.*, 156, 135-145, 2017.
- Park, K., Kim, G., Kim, J.-s., Yoon, Y.-J., Cho, H.-j., Ström, J.: Mixing State of Size-Selected Submicrometer Particles in the Arctic in May and September 2012, *Environ. Sci. Technol.*, 48 (2), 909-919, 2013.
- Peters, G.P., Nilssen, T.B., Lindholt, L., Eide, M.S., Glomsrød, S., Eide, L.I., Fuglestad, J.S.: Future emissions from shipping and petroleum activities in the Arctic, *Atmos. Chem. Phys.*, 11 (11), 5305-5320, 2011.
- Qi, L., Li, Q., Li, Y., He, C.: Factors controlling black carbon distribution in the Arctic, *Atmos. Chem. Phys.*, 17 (2), 1037-1059, 2017.
- Quinn, P.K., Shaw, G., Andrews, E., Dutton, E.G., Ruoho-Airola, T., Gong, S.L.: Arctic haze: current trends and knowledge gaps, *Tellus B*, 59 (1), 99-114, 2007.
- Raatikainen, T., Brus, D., Hyvärinen, A.P., Svensson, J., Asmi, E., Lihavainen, H.: Black carbon concentrations and mixing state in the Finnish Arctic, *Atmos. Chem. Phys.*, 15 (17), 10057-10070, 2015.
- Riemer, N., West, M., Zaveri, R.A., Easter, R.C.: Simulating the evolution of soot mixing state with a particle resolved aerosol model, *J. Geophys. Res.*, 114, doi:10.1029/2008JD011073, 2009.
- Saleh, R., Hennigan, C.J., McMeeking, G.R., Chuang, W.K., Robinson, E.S., Coe, H., Donahue, N.M., Robinson, A.L.: Absorptivity of brown carbon in fresh and photo-chemically aged biomass-burning emissions, *Atmos. Chem. Phys.*, 13 (15), 7683-7693, 2013.
- Samset, B.H., Myhre, G., Herber, A., Kondo, Y., Li, S.M., Moteki, N., Koike, M., Oshima, N., Schwarz, J.P., Balkanski, Y., Bauer, S.E., Bellouin, N., Bernsten, T.K., Bian, H., Chin, M., Diehl, T., Easter, R.C., Ghan, S.J., Iversen, T., Kirkevåg, A., Lamarque, J.F., Lin, G., Liu, X., Penner, J.E., Schulz, M., Seland, Ø., Skeie, R.B., Stier, P., Takemura, T., Tsigaridis, K., Zhang, K.: Modelled black carbon radiative forcing and atmospheric lifetime in AeroCom Phase II constrained by aircraft observations, *Atmos. Chem. Phys.*, 14 (22), 12465-12477, 2014.
- Sand, M., Bernsten, T.K., Kay, J.E., Lamarque, J.F., Seland, Ø., Kirkevåg, A.: The Arctic response to remote and local forcing of black carbon, *Atmos. Chem. Phys.*, 13 (1), 211-224, 2013.
- Seinfeld, J., Pandis, S., 2006. *Atmospheric Chemistry and Physics: From air pollution to climate change* (2nd ed.). 1-1203 pp., John Wiley & Son, Inc., Hoboken, New Jersey.

- Shindell, D.: Local and remote contributions to Arctic warming, *Geophys. Res. Lett.*, 34 (14), L14704, 2007.
- Sierau, B., Chang, R.Y.W., Leck, C., Paatero, J., Lohmann, U.: Single-particle characterization of the high-Arctic summertime aerosol, *Atmos. Chem. Phys.*, 14 (14), 7409-7430, 2014.
- Updyke, K.M., Nguyen, T.B., Nizkorodov, S.A.: Formation of brown carbon via reactions of ammonia with secondary organic aerosols from biogenic and anthropogenic precursors, *Atmos. Environ.*, 63 (0), 22-31, 2012.
- Währnschimmel, H., MacLeod, M., Hungerbühler, K.: Emissions, Fate and Transport of Persistent Organic Pollutants to the Arctic in a Changing Global Climate, *Environ. Sci. Technol.*, 47 (5), 2323-2330, 2013.
- Wang, B., Laskin, A., Roedel, T., Gilles, M.K., Moffet, R.C., Tivanski, A.V., Knopf, D.A.: Heterogeneous ice nucleation and water uptake by field-collected atmospheric particles below 273 K, *J. Geophys. Res.*, 117, 2012.
- Wang, X., Heald, C.L., Liu, J., Weber, R.J., Campuzano-Jost, P., Jimenez, J.L., Schwarz, J.P., Perring, A.E.: Exploring the observational constraints on the simulation of brown carbon, *Atmos. Chem. Phys.*, 18 (2), 635-653, 2018.
- Weinbruch, S., Wiesemann, D., Ebert, M., Schütze, K., Kallenborn, R., Ström, J.: Chemical composition and sources of aerosol particles at Zeppelin Mountain (Ny Ålesund, Svalbard): An electron microscopy study, *Atmos. Environ.*, 49 (0), 142-150, 2012.
- Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Gustafsson, Ö.: The sources of atmospheric black carbon at a European gateway to the Arctic, *Nat. Commun.*, 7, 12776, 2016.
- Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Semiletov, I.P., Dudarev, O.V., Charkin, A., Shakhova, N., Klimont, Z., Heyes, C., Gustafsson, Ö.: Siberian Arctic black carbon sources constrained by model and observation, *P. Natl. Acad. Sci. USA*, 114 (7), E1054-E1061, 2017.
- Xu, J.W., Martin, R.V., Morrow, A., Sharma, S., Huang, L., Leaitch, W.R., Burkart, J., Schulz, H., Zanutta, M., Willis, M.D., Henze, D.K., Lee, C.J., Herber, A.B., Abbatt, J.P.D.: Source attribution of Arctic black carbon constrained by aircraft and surface measurements, *Atmos. Chem. Phys.*, 17 (19), 11971-11989, 2017.
- Yang, Y., Wang, H., Smith, S.J., Easter, R.C., Rasch, P.J.: Sulfate Aerosol in the Arctic: Source Attribution and Radiative Forcing, *J. Geophys. Res.*, 123 (3), 1899-1918, 2018.
- You, Y., Renbaum-Wolff, L., Carreras-Sospedra, M., Hanna, S.J., Hiranuma, N., Kamal, S., Smith, M.L., Zhang, X., Weber, R.J., Shilling, J.E., Dabdub, D., Martin, S.T., Bertram, A.K.: Images reveal that atmospheric particles can undergo liquid-liquid phase separations, *P. Natl. Acad. Sci. USA*, 109 (33), 13188-13193, 2012.
- Zanutta, M., Laj, P., Gysel, M., Baltensperger, U., Vratolis, S., Eleftheriadis, K., Kondo, Y., Dubuisson, P., Winiarek, V., Kazadzis, S., Tunved, P., Jacobi, H.W.: Effects of mixing state on optical and radiative properties of black carbon in the European Arctic, *Atmos. Chem. Phys.*, 18 (19), 14037-14057, 2018.
- Zangrando, R., Barbaro, E., Zennaro, P., Rossi, S., Kehrwald, N.M., Gabrieli, J., Barbante, C., Gambaro, A.: Molecular Markers of Biomass Burning in Arctic Aerosols, *Environ. Sci. Technol.*, 47 (15), 8565-8574, 2013.

Figure Captions

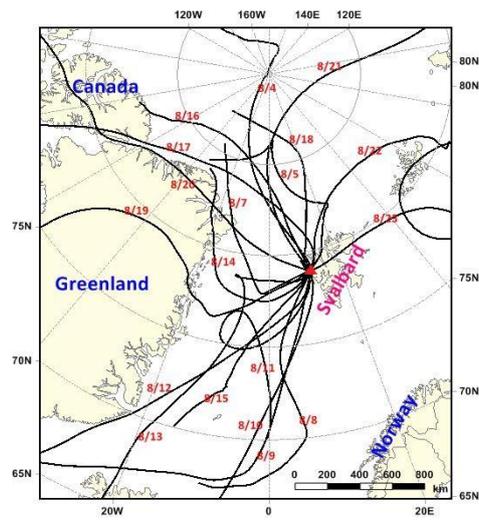


Figure 1 72 h back trajectories of air masses at 500m over Arctic Yellow River Station in Svalbard during 3–26 August 2012, and arriving time was set according to the sampling time

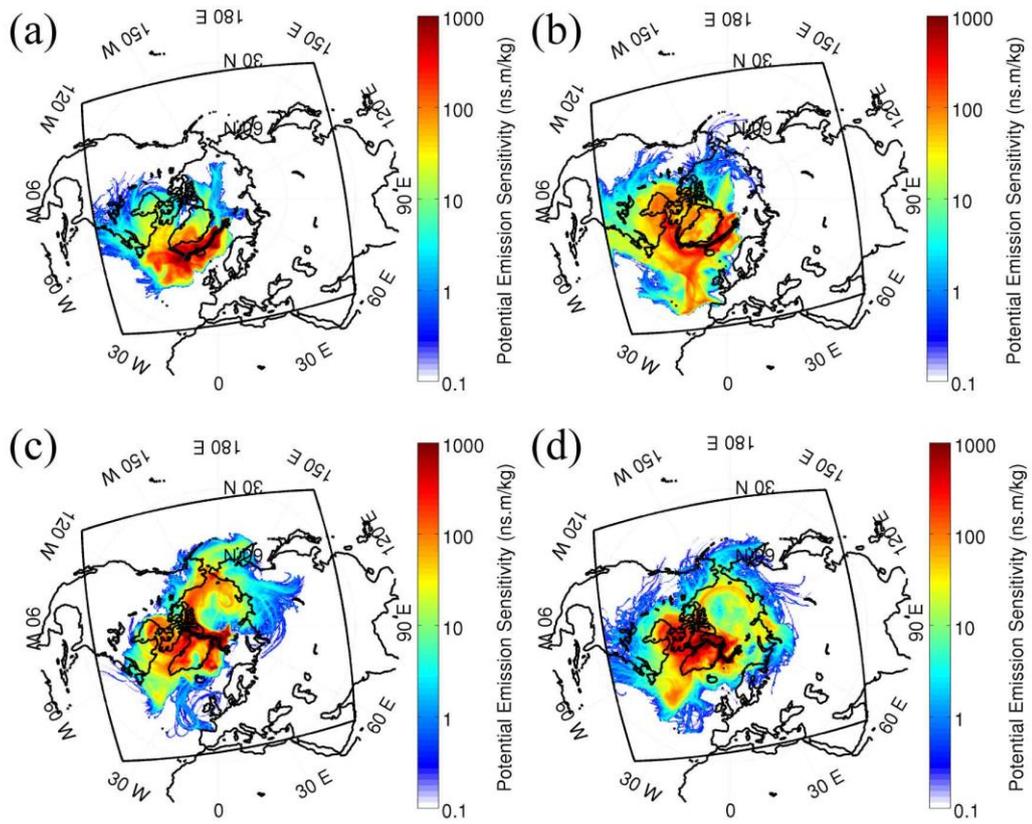


Figure 2 FLEXPART-WRF PES on August 11, 12, 14, and 15, 2012. Black square is showing the WRF domain used to initiate the FLEXPART-WRF simulation.

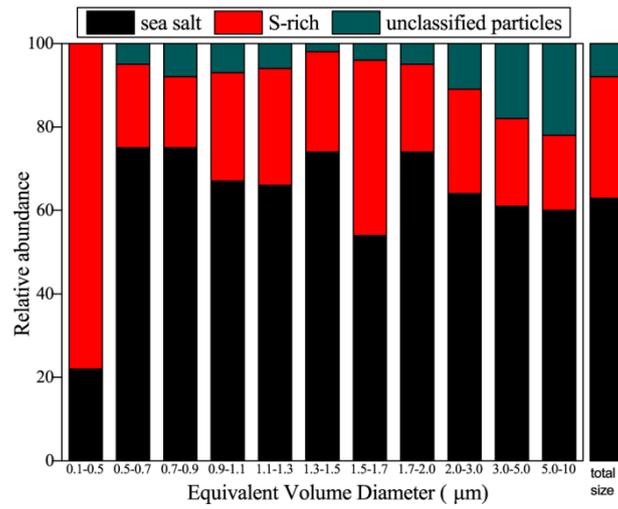


Figure 3 Morphology and relative abundances of typical individual aerosol particles in the 21 analyzed samples.

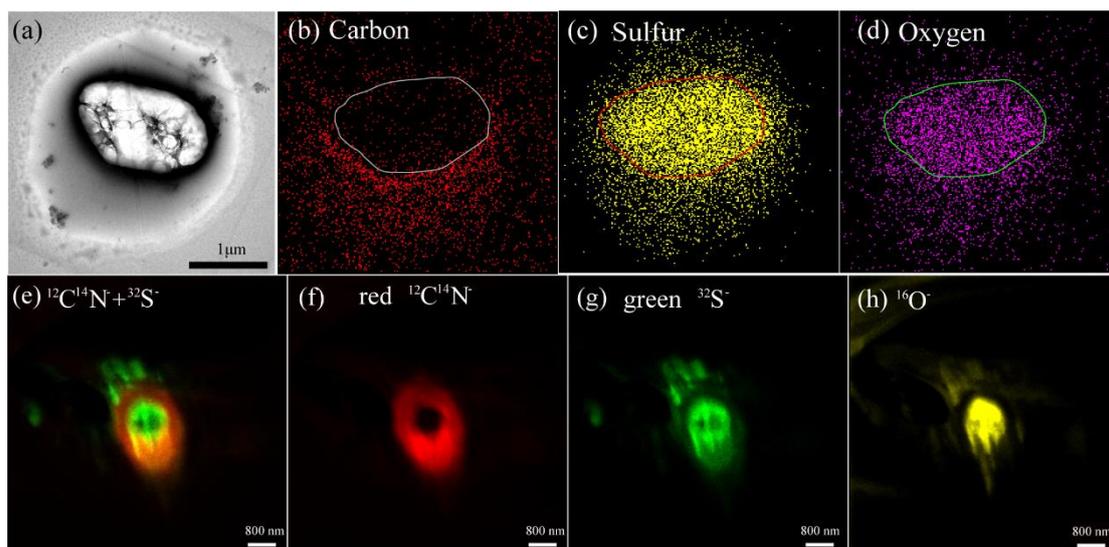


Figure 4 TEM Observations of a secondary particle and NanoSIMS intensity threshold maps of an aerosol particle with sulfate core and OM coating. (a) Bright-field TEM image of an internally mixed particle; (b) elemental carbon (c) sulfur and (d) oxygen maps of the internally mixed particle shown in 1(a); (e) Overlay of $^{12}\text{C}^{14}\text{N}^-$ and $^{32}\text{S}^-$ ion maps in an internally mixed particle; (f) CN^- map (g) S^- (h) O^- secondary ion maps. Ion maps with a set of aerosol particles were shown in Figure S1.

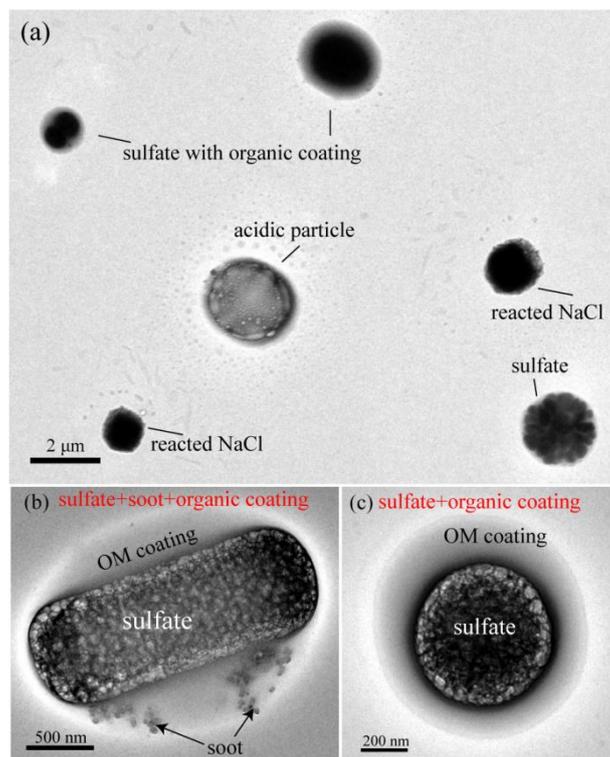


Figure 5 TEM images of individual particles containing sulfate, OM, and soot. (a) Low magnification TEM image showing sulfates, sulfate with OM coating, and reacted NaCl particles. (b) an internally mixed particle of sulfate and soot with OM coating (c) a particle with sulfate core and OM coating.

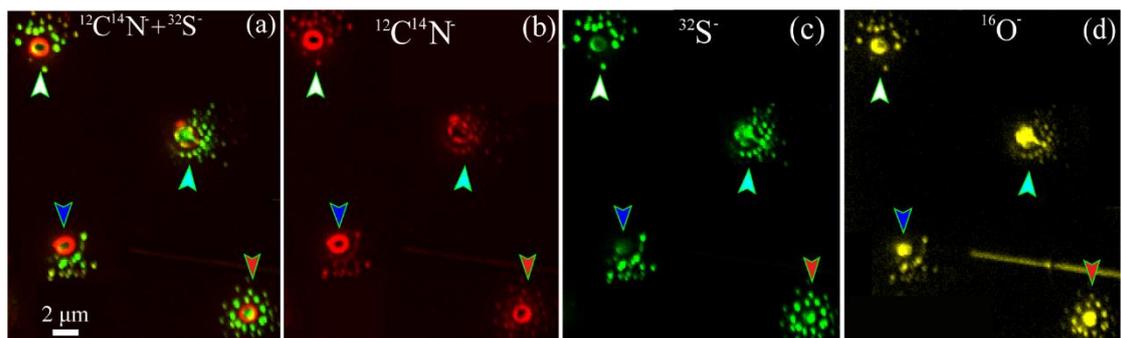


Figure 6 NanoSIMS intensity threshold maps of individual aerosol particles surrounded by satellite particles. (e) Overlay of $^{12}\text{C}^{14}\text{N}^-$ and $^{32}\text{S}^-$ ion maps of individual particles. (f) CN^- (g) S^- (h) O^- maps. Four particles were indicated by white, pink, blue, and red arrows.

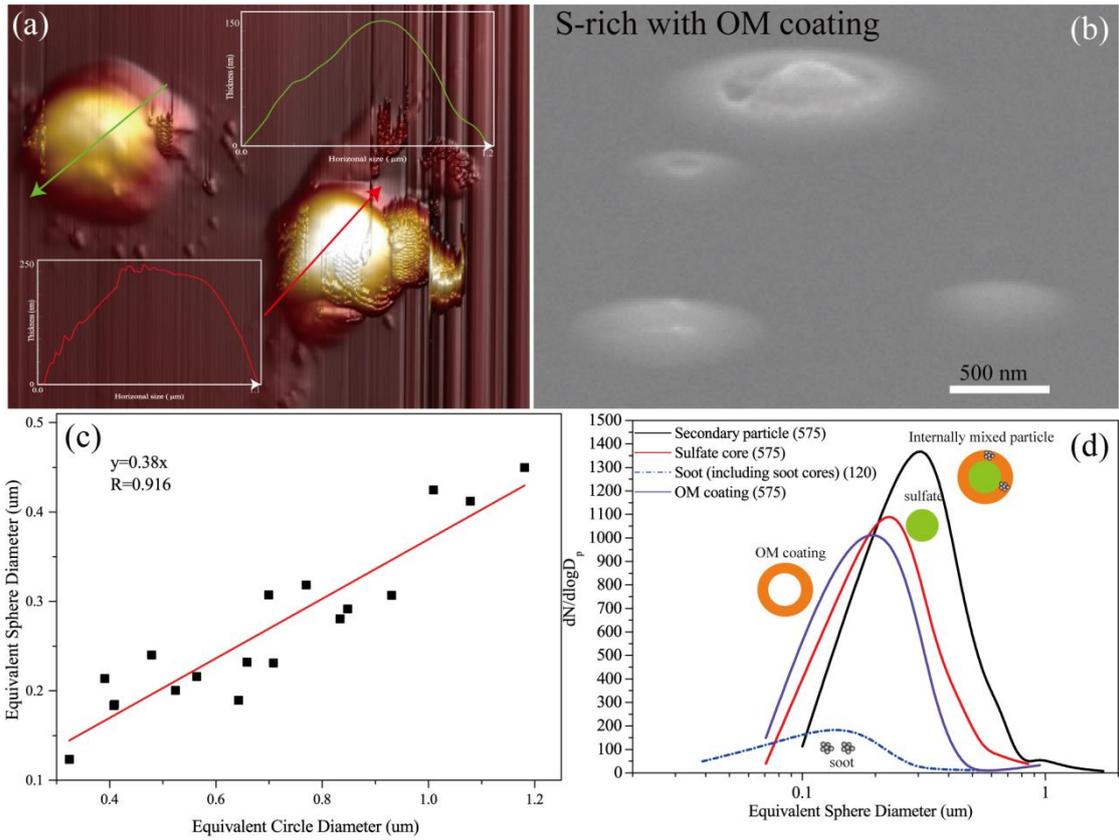


Figure 7 Secondary particles on the substrate. (a) 3-D AFM image of secondary sulfate particles. The colorful arrows represent particles surface properties of the particle section. (b) SEM image of S-rich with OM coating obtained from 75° tilt of the SEM specimen stage (c) The near linear relationships between ECD and ESD based on S-rich particles with thick OM coating by Atomic force microscopy. (d) Size distribution of individual particle with OM coating and sulfate cores based on the estimated ESD diameter from TEM image. Sizes of soot particles are equal to the equivalent circle diameter.

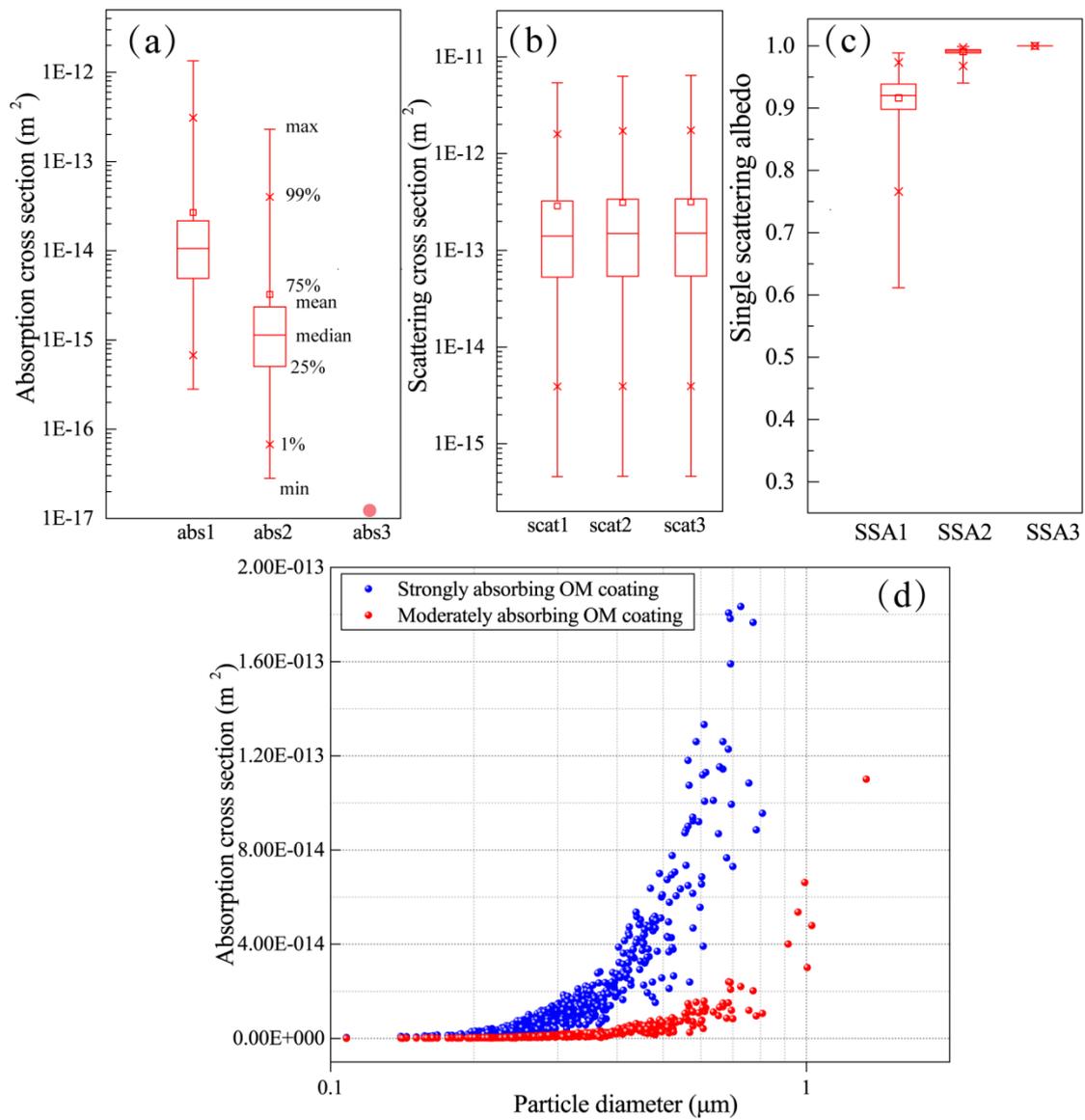


Figure 8 Optical properties of Box-and-whisker plots showing optical parameters of all analysed particles assuming sulfate core and BrC shell (not considering soot cores in the particles). (a) Scattering cross section (b) Absorption cross section (c) Single scattering albedo. Top to bottom markers in the box-and-whisker represent max, 99%, 75%, mean, median, 25%, 1%, min values. (d) Absorption cross section along with particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating.