



EGUsphere, author comment AC1  
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## Reply on CC1, regarding tides other than M2

Richard Ray

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Author comment on "Technical note: On seasonal variability of the M<sub>2</sub> tide" by Richard D. Ray, EGU sphere, <https://doi.org/10.5194/egusphere-2022-252-AC1>, 2022

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As Haidong Pan notes, there are many tidal constituents in addition to (the usually dominant) M<sub>2</sub> that could be examined for seasonal variability. For lunar tides like O<sub>1</sub> and N<sub>2</sub>, any analysis of seasonality should begin by checking for the presence of nearby spectral lines, in the manner laid out in my Note for M<sub>2</sub>. In each case there are small astronomical constituents within the relevant tidal group, as well as compound tides and climate-driven lines at frequencies 1 or 2 cpy away from the central constituent. One simply needs to take care not to overlook an important contributor, such as a compound tide (of which there are many possible).

The solar tides, as Haidong Pan points out, are more problematic. For S<sub>2</sub> there is semiannual modulation from K<sub>2</sub>, but also annual modulation from T<sub>2</sub> and R<sub>2</sub>. Harmonic analysis is probably not recommended in this case, but some rough results can be obtained by a response analysis (e.g., Cartwright, 1968). With a response analysis, the gravitational parts of K<sub>2</sub>, T<sub>2</sub>, and R<sub>2</sub> can be approximately determined from estimates of the mean S<sub>2</sub> tide. Any residual modulations seen in S<sub>2</sub> can then presumably be attributed to seasonal climate variability. But this is only a "rough" approach, because it overlooks the radiational forcing of S<sub>2</sub>, caused by loading by the S<sub>2</sub> atmospheric tide, which itself has significant seasonal variability. The radiational forcing of the S<sub>2</sub> ocean tide has been studied by Arbic (2005) and others. In the end, even with a response analysis, we may isolate a seasonal S<sub>2</sub> signal, but it may not be clear whether it is originating in the ocean (e.g., from seasonal stratification) or in the atmosphere (from air tides).

The diurnal K<sub>1</sub> is just as difficult, if not more so. Again a response analysis may be used to determine the gravitational parts of the neighboring constituents P<sub>1</sub> and psi<sub>1</sub>. But the S<sub>1</sub> constituent, 1 cpy from K<sub>1</sub>, is almost wholly radiational, with significant temporal variability. In monthly estimates of K<sub>1</sub>, modulations from S<sub>1</sub> would be difficult to untangle from other seasonal changes.

For these reasons, analysis of lunar tides is more straightforward. Understanding their seasonality can still be difficult and even the spectral approach requires care. For example, O<sub>1</sub> seasonality was briefly examined at station Lusi (China) in the course of a review of coastal tides (Ray et al., 2011). Small semiannual variations in O<sub>1</sub> (of order 1%) will be induced by the linear constituent tau<sub>1</sub>, 2 cpy away, but monthly estimates of O<sub>1</sub> at Lusi revealed significantly larger variations than that. Unfortunately, in our 2011 discussion, we overlooked the possible presence of the compound tide MP<sub>1</sub> (which coincides with tau<sub>1</sub>). Because both M<sub>2</sub> and P<sub>1</sub> are large along the China coastline, the compound MP<sub>1</sub> is almost certainly responsible for most of the semiannual oscillation seen in monthly estimates of O<sub>1</sub> at Lusi. Nonetheless, there is still a significant annual variation

in O1, which is probably induced by climatic changes in ocean stratification, in the same manner as seen for M2 in that region (Kang et al., 2002).

Finally, seasonality of the compound KO2 could indeed mimic seasonality in M2. Presumably, in general, the effect is very small, but it is potentially at work in shallow-water regions with large diurnal tides and small M2.

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