

Interactive comment on “Simulating oceanic radiocarbon with the FAMOUS GCM: implications for its use as a proxy for ventilation and carbon uptake” by Jennifer E. Dentith et al.

Anonymous Referee #1

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Dentith et al. describe the implementation of radiocarbon (^{14}C , or $\Delta^{14}\text{C}$ when considering the normalized and fractionation-corrected $^{14}\text{C}/^{12}\text{C}$ ratio) into the ocean component of the FAMOUS atmosphere-ocean GCM and present two ^{14}C simulations, one spin-up simulation with constant pre-industrial boundary conditions plus a “historical” simulation forced with transient values of atmospheric CO_2 and $\Delta^{14}\text{C}$ for 1765–2000 CE. The simulation results are compared with water column measurements available for the 1950s and the 1990s, and with ^{14}C records from bivalves and deep-water corals spanning the period 1880–2000 CE. The model results are at least qualitatively in line with observations. FAMOUS simulates radiocarbon in two ways. The “abiotic” approach only considers uptake, transport and radioactive decay of normalized $^{14}\text{C}/^{12}\text{C}$.

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The “biotic” approach also considers isotopic fractionation due to air-sea gas exchange and the isotopic imprint of the biological pump. Comparing both approaches, Dentith et al. find that the “biotic” approach results in slightly elevated $^{14}\text{C}/^{12}\text{C}$ ratios in the deep sea (by about 20‰ in terms of $\Delta^{14}\text{C}$). This result corroborates early findings stating that biological effects on $^{14}\text{C}/^{12}\text{C}$ are much smaller (<10%) than the effects of transport and radioactive decay (Fiadeiro, 1982). In addition, Dentith et al. compare simulated ^{14}C water ages with “true” water ages according to an idealized age tracer. It turns out that for large parts of the oceans, a purely kinematic interpretation of $^{14}\text{C}/^{12}\text{C}$ ages can lead to erroneous conclusions.

The paper is well written, the presentation of the results is mostly clear, and the literature is comprehensive. It may be published if the following issues are amended:

Major issue:

As described in Section 2.3.2, the model setup does not allow for meridional variability of atmospheric $\Delta^{14}\text{C}$. Therefore, the historical simulation misses the strong interhemispheric $\Delta^{14}\text{C}$ gradient in the atmosphere of up to 400‰ (e.g., Fig. 1 in the manuscript) due to nuclear weapons testing. As a consequence, the model may underestimate the actual spatiotemporal ^{14}C gradients in the ocean since the late 1950s. This shortcoming may distort the evolution of bomb ^{14}C transients discussed in Section 3.3. It may also lead to biased post-bomb ^{14}C distributions in subsurface and deeper waters which are discussed in Section 3.2. I would strongly recommend repeating the historical simulation forced with hemispheric averages of atmospheric $\Delta^{14}\text{C}$. The implementation should not be too difficult (the model has to read in three files provided by OCMIP-2).

Further issues and comments (P = page, L = line, SM = supplementary material):

P 3, L 6: The value of the half-life should be consistent with the updated value of 5700 years promoted in the SM.

P 3, L 28: “Indian” should read “Indien”

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P 4, L 28: The models MOM (Toggweiler et al., 1989), LOCH (Mouchet, 2013) and LOVEVLIM (e.g., Menviel et al. 2017) could be added to the list.

P 7, L 28: $(660 / Sc)^{-0.5}$ should read $(Sc / 660)^{-0.5}$ (e.g., Orr et al. 2017, equation (12))

P 7, L 31: Missing reference (I guess it is Wanninkhof 1992)

P 8, L 8 and P 22, L 27: I could not find the paper at GMDD. What is the current status of the manuscript?

P 8, L 12: 14 and 13 should be subscripts

P 8, L 22: “ $\delta^{13}C$ is close to zero”: this is the only case for DIC but not for phytoplankton

P 9, L 10: This is not correct, Stuiver and Pollach (1977) employ 5568 years.

P 9, L 16: What about turbulent diffusion, where is it included?

P 9, L 26–28 and Figure S1: The OCMIP-2 steady state criterion is somewhat different, demanding that 98% of the ocean volume has a $\Delta^{14}C$ drift of less than 0.001‰ per year (see also Aumont et al. 1998). The total ^{14}C inventory may have stabilized while there might be still an ongoing internal redistribution of ^{14}C . Have you checked this?

P 13, L 5–6: “(...) the simulated $\Delta^{14}C$ gradient between the surface ocean and approximately 1000 m depth is shallower than observed (...). This is visible in all regions, except the Northern Hemisphere deep water formation region (NH_DWF) and the North Pacific (NP) (...).” According to Figure 8 the opposite is true (indicating steeper simulated gradients everywhere except for NH_DW and NP).

P 13, L 23: “(...) due to the sparsity of data.” You might wish to compare your results with post-bomb $\Delta^{14}C$ (bottle) data provided by GLODAPv2 (<https://odv.awi.de/data/ocean/glodap-v22019-bottle-data/>).

P 17, L 19: This is an interesting result supporting and quantifying the notion by Fi-

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adeiro (1982) that $^{14}\text{C}/^{12}\text{C}$ can roughly be regarded as radio-conservative tracer in the deep sea.

P 18, L 1–6: The simulated ^{14}C ages reflect local ageing, transport, biogeochemistry, and radioactive decay. On the other hand, the simulated ideal water ages only reflect local ageing and transport. As there is no radioactive decay of the age tracer, the age differences cannot be entirely explained with biogeochemical effects.

P 18, L 15–16: “(...) the water ages generally increase with depth because they are a simple function of advection.” I disagree, there may be considerable mixing leading to nonlinear ageing of water parcels. This is seen in many tracers records.

P 31, Figure 4 (b): I would prefer to see the differences between simulated and observed values, analogously to Figure 2.

Figures 8, 10, 11, 16, S4, S5, and S6: Lines and dots are blurred.

P 38, Figure 11: The line plots should be replaced with Hovmöller diagrams.

SM, P 2: “OCMIP” should read “OCMIP-2”

References:

Aumont, O.; Orr, J. C.; Jamous, D.; Monfray, P.; Marti, O.; Madec, G. A Degradation Approach to Accelerate Simulations to Steady-State in a 3-D Tracer Transport Model of the Global Ocean. *Climate Dynamics* 1998, 14 (2), 101–116.

Fiadeiro, M. E. Three-Dimensional Modeling of Tracers in the Deep Pacific Ocean, II. Radiocarbon and the Circulation. *Journal of Marine Research* 1982, 40, 537–550.

Menviel, L.; Yu, J.; Joos, F.; Mouchet, A.; Meissner, K. J.; England, M. H. Poorly Ventilated Deep Ocean at the Last Glacial Maximum Inferred from Carbon Isotopes: A Data-Model Comparison Study. *Paleoceanography* 2017, 32 (1), 2016PA003024.

Mouchet, A. The Ocean Bomb Radiocarbon Inventory Revisited. *Radiocarbon* 2013,

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55, 1580–1594. Orr, J. C.; Najjar, R. G.; Aumont, O.; Bopp, L.; Bullister, J. L.; Danabasoglu, G.; Doney, S. C.; Dunne, J. P.; Dutay, J.-C.; Graven, H.; et al. Biogeochemical Protocols and Diagnostics for the CMIP6 Ocean Model Intercomparison Project (OMIP). *Geosci. Model Dev.* 2017, 10 (6), 2169–2199.

Orr, J. C.; Najjar, R. G.; Aumont, O.; Bopp, L.; Bullister, J. L.; Danabasoglu, G.; Doney, S. C.; Dunne, J. P.; Dutay, J.-C.; Graven, H.; et al. Biogeochemical Protocols and Diagnostics for the CMIP6 Ocean Model Intercomparison Project (OMIP). *Geosci. Model Dev.* 2017, 10 (6), 2169–2199.

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