

Atmos. Chem. Phys. Discuss., author comment AC1  
<https://doi.org/10.5194/acp-2021-24-AC1>, 2021  
© Author(s) 2021. This work is distributed under  
the Creative Commons Attribution 4.0 License.

## Reply on RC1

Keiichiro Hara et al.

---

Author comment on "Characterization of aerosol number size distributions and their effect on cloud properties at Syowa Station, Antarctica" by Keiichiro Hara et al., Atmos. Chem. Phys. Discuss., <https://doi.org/10.5194/acp-2021-24-AC1>, 2021

---

Reply to comments from Reviewer#1:

We would like to thank your helpful comments to improve our manuscript. All comments are responded and addressed in the current revise. Details are listed as follows.

Comment from Referee: Unfortunately, the article in its present appearance has some weak points. It is regrettable that in several issues the reasoning is barely traceable and inadequately supported by the presented data evaluation. This is especially obvious concerning the claimed link between ozone hole → NPF → CCN → cloud properties → cloud amount. I think it is worth the effort addressing this weakness and considering a more in-depth analysis.

Reply from Author: When aerosol particles derived from NPF enhanced by ozone hole are grown to critical size for CCN activation, these particles can play important roles in the link. It is certain that model simulation is one of useful method to understand atmospheric processes as pointed by you, but we still have many uncertain parameters and knowledge such as aerosol size distributions, specific chemical forms of the condensable vapors, and the concentrations of the vapors in the Antarctic free troposphere. Because we must put many assumptions with high uncertainty into the model calculation, we do not use model simulation in the revise manuscript. Instead, we attempted to estimate aerosol lifetime from coagulation sink in the Antarctic free troposphere to ascertain this link among ozone hole, NPF, CCN ability, and cloud amount. Details are explained in each comment as follows.

Comment from Referee: Chapters 2.2.2, lines 112-119: To be honest, I do not understand the mathematics behind your estimate regarding J5. Why do you need the term Coag.S10-20N10-20, but then neglecting the condensation sink Cond5-10N5-10? Moreover, it is not even clear in which way you calculated the condensation sink shown in Figure 6d. Please specify input values for eq. (2) and in addition show the measured total particle concentrations in Figure 6.

Reply from Author: More explanation to estimate the formation rate of J5 was added into the revised manuscript. In the revised manuscript, calculation procedures were in accordance with Dal Maso et al. (2002). Also, we added CN concentrations in the revised figure, which were measured simultaneously with SMPS at Syowa Station. Also, measurement procedure was added into the section of 2.1.

Comment from Referee: Chapters 2.3: Back trajectory analyses are a crucial tool in this

study, so the authors should provide more details: Why did you rely on the NCEP meteorology data set? The GDAS dataset has a higher resolution and is more accurate in general! Why an initial starting point of 500 m above ground has been chosen, well above the aerosol measuring point? With hysplit (using GDAS input), it is possible to start trajectory ensembles from different high levels. This option could be useful to assess the reliability of the back trajectory analysis. Another point: 5-days back trajectory may not be sufficient to address the origin of particles in the accumulation mode. Finally, does the chosen 1500 m boundary level between FT and BL refer to height above ground?

Reply from Author: It is true that GDAS dataset has a higher resolution than NCEP-reanalysis, but GDAS dataset is available since January 2005. In this study, the measurement period was February 2004 – December 2006, so that we used NCEP-reanalysis dataset to keep same quality during the measurement periods. As pointed by you, 5-day backward trajectory is slightly short (not sufficient) to know the origins of aerosol particles in the accumulation mode. Considering uncertainty of the trajectory analysis in troposphere and high uncertainty at lower altitudes, we set initial height of 500m corresponding to upper boundary layer over Syowa and calculation period of 5 days. In this study, we used results of the backward trajectory mostly to identify air mass origins and pathway for nucleation modes and modal structures. Therefore, we can compare the relation and identification by 5-day backward trajectory. To divide air mass origins, we chose altitude of 1500 m, because height of top of boundary layer was distributed below 1400 m through the year, which were observed by the tethered balloon measurements in 2005 at Syowa Station (Hara et al., 2011a, 2013).

Comment from Referee: Chapters 3.2, lines 180-181: Is there any evidence from your data that in this case sea-salt originated from the snow surface? This should be specified, otherwise a reference is needed.

Reply from Author: Direct evidence was already published in previous works (other papers, e.g., Hara et al., 2012, 2020). We added more explanation and the specific references about sea-salt aerosols originated from the snow surface in the revised manuscript. Simultaneous measurements of aerosol constituents showed that sea-salt aerosols in ultrafine – coarse modes were released from sea-ice area during winter – spring (e.g., Hara et al., 2012, 2020).

Comment from Referee: Chapters 3.2, lines 185-186: In this case, 10-days back trajectories could be beneficial!

Reply from Author: As mentioned above, we did not calculate 10-day backward trajectory.

Comment from Referee: Chapters 3.3, line 223: "Particularly, fresh nucleation mode appeared only in end-August..." you mean fresh nucleation mode without aged nucleation mode?

Reply from Author: In the sentence of "Particularly, fresh nucleation mode appeared only in end-August...", we focused on fresh nucleation mode ( $D < 10$  nm), because fresh nucleation mode was direct evidence of recent NPF.

Comment from Referee: Chapters 3.3, lines 259-264: Please delete this sentence as well as the (R1) and (R2), because it is (very) basic textbook knowledge.

Reply from Author: Reactions (R1 and R2) were removed in the revised manuscript. The sentences immediately before R1 and R2 were modified in the revised manuscript.

Comment from Referee: Chapters 3.3, lines 266-267: This is not visible in Fig. 6g! I daresay that UV radiation is roughly comparable in Oct/Nov and December. Please provide numbers of the measured difference.

Reply from Author: Seasonal variation of solar radiation at Syowa (not shown in Figure) was clearly lower in October – November than that in December. Nevertheless, UV amount in October – November was higher or comparable than that in December. Especially, UV amount in shorter wavelength in October – November was obviously higher

than that in December. We added short explanation and comparison of values of UV amounts in this period into the revised manuscript.

Comment from Referee: Chapters 3.3, lines 270-279 and Chapter 3.4: This part is rather speculative and barely convincing! I agree that enhanced UV-radiation under ozone depletion conditions may potentially have an impact on NPF and CCN concentration during this period. However, in my view it is hard to believe that those NPF events and their subsequently grow to CCN relevant diameters will last more than around 2 months in the FT and will then have any significant influence on cloud properties in Dec/Jan! First, you estimated particle lifetime solely based on coagulation sink. Is there some evidence, that such a simplification is adequate? Moreover, consider that comparable photochemical processes provoking NPF in BL and FT surely also proceed during Dec/Jan (note comparable UV radiation and even more prominent DMS emissions). Thus, I cannot realize that NPF happened more than 2 months before in FT could have any significant impact on CCN concentrations and cloud properties in Dec/Jan. If at all, only detailed model simulation may give a robust answer concerning this conclusion.

Reply from Author: As stated above, we decide not to use model simulation because of many unknown and uncertain parameters in the Antarctic free troposphere. The most important issue of this point is which aerosol particles derived from NPF in the Antarctic free troposphere and boundary layer can be survived to be grown to the critical diameter for CCN activation, or not. Furthermore, Williams et al. (2002) presented clearly that lifetime of aerosol particles with size of  $D_p < 60$  nm in free troposphere was controlled by coagulation loss. In the revised manuscript, therefore, we focused on the aerosol lifetime by coagulation loss in troposphere. To our knowledge, number size distributions of aerosol particles with size with  $D_p < 100$  nm are still unknown in the Antarctic free troposphere through the year. Therefore, we assumed that aerosol mixing ratios (i.e., number in respective size bins) in the free troposphere were as same as those at surface. Then, aerosol lifetime (e-folding time) was estimated using coagulation sink in each size. Vertical variations of the e-folding time showed clearly vertical gradient with longer lifetime at higher altitudes in all sizes ( $D_p = 1 - 50$  nm). The e-folding time at upper troposphere (ca. 8.3 km, 300 hPa) was 3 - 4 times longer than that at surface. Under the conditions with the assumption above (same mixing ratios), the e-folding time of aerosol particles larger than 20-30 nm exceed 30 days in middle free troposphere. Although size distributions of aerosol particles with size of  $D_p < 100$  nm are still unknown in the Antarctic free troposphere, size distribution of aerosol particles with size larger than 300 nm was available over Syowa (1 - 2 order lower in free troposphere than at surface; Kizu et al., 2010). If vertical gradient of aerosol number concentrations and size distributions of aerosol particles smaller than 100 nm was similar to that of aerosols with  $D_p > 300$  nm under the background conditions, the longer e-folding time in the free troposphere is expected. Consequently, we conclude that the e-folding time of 1 - 2 months might be available for aerosol particles with size of  $D_p > 30$  nm in the Antarctic free troposphere, with vertical gradient of the number concentrations of pre-existing particles. This discussion was added into the revised manuscript.

Comment from Referee: Chapters 3.4, line 293, Figures 8, and 9b: Does the calculated significance level refer to an ANOVA variance test, meaning that the corresponding distributions are significantly different on this level.

Reply from Author: ANOVA variance test was made for the revised manuscript. The result was added into the revised figure. Also, ANOVA test showed significant difference ( $p = 0.0125$ ).

Comment from Referee: Figure 5: Please specify the black lines and dots in the figure caption.

Reply from Author: Explanation of black lines and dots were added into the caption. The lines and dots showed the data numbers (n in y-axis).

All corrected parts and sentences were marked by red characters in the revised manuscript.

Please also note the supplement to this comment:

<https://acp.copernicus.org/preprints/acp-2021-24/acp-2021-24-AC1-supplement.pdf>