

1 **Comparison of Vaisala radiosondes RS41 and RS92**  
2 **launched over the oceans from the Arctic to the Tropics**

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19 **Abstract.** To assess the differences between the RS92 radiosonde and its improved  
20 counterpart, the Vaisala RS41-SGP, radiosonde version with a pressure sensor, 36  
21 twin-radiosonde launches were made over the Arctic Ocean, Bering Sea, northwestern

22 Pacific Ocean, and the tropical Indian Ocean during two cruises of the R/V *Mirai* in 2015.

23 The biases, standard deviations, and root mean squares (RMSs) of the differences

24 between the RS41 and RS92 data over all flights and altitudes were smaller than the

25 nominal combined uncertainties of the RS41, except that the RMS of the differences of

26 pressure above 100 hPa exceeded 0.6 hPa. A comparison between daytime and nighttime

27 flights in the tropics revealed that the pressure difference was systematically larger during

28 the day than at night above an altitude of 4.5 km, suggesting that there was some effect of

29 solar heating on the pressure measurements, but the exact reason is unclear. The

30 agreement between the RS41 and RS92 temperature measurements was better than the

31 combined uncertainties. However, there were some noteworthy discrepancies presumably

32 caused by the “wet-bulbing” effect on the RS92 radiosonde and the stagnation of the

33 balloon. Although the median of the relative humidity differences was only a little more than

34 2 % of the relative humidity at all altitudes, the relative humidity of the RS92 was much

35 lower than that of the RS41 at altitudes of about 17 km in the tropics. This dry bias might

36 have been caused by the incomplete solar radiation correction of the RS92, and a

37 correction table for the daytime RS92 humidity was calculated. This study showed that the

38 RS41 measurements were consistent with the specifications of the manufacturer in most

39 cases over both the tropical and polar oceans. However, further studies on the causes of

40 the discrepancies are needed.

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## 42 **1 Introduction**

43 Radiosonde observations are operationally conducted twice a day at about 800 sites  
44 throughout the world. Radiosondes measure temperature, humidity, wind velocities, and  
45 pressure (or height) in the troposphere and stratosphere. They ascend through the  
46 atmosphere attached to balloons filled with helium or hydrogen gas. The data are sent to  
47 the global telecommunication system and are used for data assimilation in real-time  
48 operational weather forecast systems, atmospheric reanalyses, and climate models. In situ  
49 aerological observations are also indispensable for validating satellite-derived  
50 meteorological data (e.g. Fujita et al., 2008), for assessing long-term trends in the upper  
51 atmosphere (e.g. Thorne et al., 2005; Maturilli and Kayser, 2016), and for other  
52 meteorological research, including assimilation experiments and air-sea interaction studies  
53 (e.g. Inoue et al., 2013; 2015; Kawai et al., 2014). Efforts to improve the quality of  
54 radiosonde data have continued to the present time (e.g. Ciesielski et al., 2014; Bodeker et  
55 al., 2016). One consequence of the technological advancements has been the need to  
56 account for accuracy differences following radiosonde upgrades in the long-term  
57 continuous datasets (Wang et al., 2013).

58 The model RS92 radiosonde manufactured by Vaisala Ltd., which was first introduced  
59 in 2003, has been used throughout the world, and it is now being replaced with a successor  
60 model, the RS41 (Table 1). To clarify the differences between the RS41 and RS92  
61 radiosondes, intercomparison experiments have already been carried out at several sites

62 on land from high latitudes to the tropics (Möhl, 2014; Jauhiainen et al., 2014; Jensen et al.,  
63 2016). Jauhiainen et al. (2014) have reported results of comparisons in several countries,  
64 including Finland, the United Kingdom, the Czech Republic, and Malaysia. They reported  
65 that the RS41 radiosonde was a consistent improvement over the RS92 in terms of  
66 reproducibility with respect to temperature and humidity under both day and night  
67 conditions. A different intercomparison study was carried out at a site in Oklahoma, USA, by  
68 Jensen et al. (2016). They showed that the RS92 and RS41 measurements agreed much  
69 better than the manufacturer-specified combined uncertainties. Their results also indicated  
70 that the RS41 measurements of temperature and humidity appeared to be less sensitive to  
71 solar heating than those made with the RS92.

72 The accuracy of the pressure measured with the model RS41-SGP, however, has not  
73 yet been examined, nor has a comparison been made between the RS41 and RS92  
74 radiosondes in the marine atmosphere. Unlike the atmosphere over land, the marine  
75 atmosphere is less affected by topography and the greater temperature variations of the  
76 land surface. As a result, phenomena such as convection and precipitation and their diurnal  
77 cycles over the oceans are different from those over land (e.g. Yang and Slingo, 2001;  
78 Minobe and Takebayashi, 2015). We performed a total of 36 intercomparison flights during  
79 two cruises of R/V *Mirai* of the Japan Agency for Marine-Earth Science and Technology  
80 (JAMSTEC) in 2015. Our observations covered a wide range of latitudes over the oceans,  
81 an important consideration from the standpoint of confirming the performance of the RS41.

82 We describe the cruises and the methodology of the intercomparison observations in Sect.  
83 2. Section 3 shows the results of the comparisons. In Sect. 4, we focus on the data obtained  
84 in the tropics and further discuss the reasons for the differences between the RS41 and  
85 RS92 results. Section 5 is a summary of the study.

## 86 **2 Intercomparison experiment**

### 87 **2.1 Cruises**

88 The intercomparison observations were performed by launching both the RS41 and RS92  
89 radiosondes tied to one balloon (referred to as a “twin-radiosonde” flight) during the  
90 MR15-03 and MR15-04 cruises of R/V *Mirai*. In the case of the MR15-03 cruise, the vessel  
91 departed from Hachinohe, Japan, on 26 August, cruised the Arctic Ocean from 6  
92 September to 3 October (Nishino et al., 2015), and returned to Hachinohe on 21 October.  
93 The twin-radiosonde flights were launched 9 times in the Chukchi Sea, 4 times in the Bering  
94 Sea, and 5 times in the northwestern Pacific (Fig. 1a and Table 2). The MR15-04 cruise  
95 was for tropical meteorological research, and the vessel stayed near 4°04' S, 101°54' E off  
96 Bengkulu, west of Sumatra Island, in the Indian Ocean during 23 November to 17  
97 December for stationary observations, including 16 twin-radiosonde flights (Katsumata et  
98 al., 2015). We also conducted intercomparison observations twice in the western Pacific on  
99 the way from Japan to the site off Sumatra (Fig. 1b and Table 2).

### 100 **2.2 Methods**

101 We used radiosonde models RS92-SGPD and RS41-SGP in this study. Their nominal  
102 accuracies are summarized in Table 1. Whereas the RS41-SG radiosonde used in the  
103 previous studies (Motl, 2014; Jauhiainen et al., 2014; Jensen et al., 2016) derived pressure  
104 from Global Positioning System (GPS) data with no pressure sensor, the RS41-SGP has a  
105 pressure sensor consisting of a silicon capacitor. The pressure and height data analyzed in  
106 this study were measured directly and derived from the hypsometric equation, respectively.  
107 Note that GPS-derived pressure and height were not used, unlike in the previous studies.  
108 Two different DigiCORA systems were used on R/V *Mirai* for the simultaneous RS92 and  
109 RS41 soundings. The receiving system (MW41) used for the RS41 included a processor  
110 (SPS331), processing and recording software (MW41 v2.2.1), GPS antenna (GA20), and  
111 UHF antenna (RB21), which was part of the ASAP sounding station permanently installed  
112 on R/V *Mirai*. The RS41 sensors were calibrated with a new calibrator (RI41) and a  
113 barometer (PTB330). In contrast, we used a previous generation system for the RS92: the  
114 receiving system (MW31) included a processor (SPS311), software (DigiCORA v3.64),  
115 GPS antenna (GA31), and UHF antenna (RM32). The instrumentation was temporarily  
116 placed in or on the aft wheelhouse. The RS92 sensors were calibrated with a calibrator  
117 (GC25) and a PTB330 barometer. Because version 3.61 of DigiCORA was incorrectly used  
118 during the cruises, all RS92 sounding data were simulated with DigiCORA v3.64 after the  
119 cruises.

120 The RS41 and RS92 radiosondes were directly attached to each other with sticky

121 tape (Fig. 2) instead of hanging them from the two ends of a rod (Jensen et al., 2016) to  
122 facilitate the launching operations on the rocking ship deck. The two radiosondes were  
123 hung from a single 350g Totex balloon with the cord of the RS41 radiosonde. The ascent  
124 rates were approximately  $5 \text{ m s}^{-1}$  and  $4 \text{ m s}^{-1}$  during the MR15-03 and MR15-04 cruises,  
125 respectively (Table 2). Whereas nighttime twin-radiosonde flights could be carried out only  
126 once during the MR15-03 cruise owing to operations associated with oceanographic  
127 observations, we performed eight nighttime flights during the MR15-04 cruise (Fig. 1c and  
128 Table 2). In addition information about surface meteorological state, Table 2 lists convective  
129 available potential energy (CAPE), convective inhibition (CIN), and precipitable water (PW)  
130 calculated from RS41 data. CAPE and CIN were calculated for an air parcel corresponding  
131 to an average over the lowest 50 hPa.

132 A number of issues were addressed in post-processing the sounding data. During  
133 flight No. 33 (02:50 UTC on 16 Dec.), the radiosondes oscillated vertically about the  $0^{\circ}\text{C}$   
134 level likely due to icing on the balloon, and hence only the data before the up-and-down  
135 motion were analyzed in this study. In the case of flight No. 9 (05:30 UTC on 16 Sep.), we  
136 delayed the measurement time of the RS41 by 17 s in the analysis because the twin  
137 radiosondes flew horizontally just after launching, and the automatic determinations of the  
138 starting times disagreed between the RS92 and RS41. Because the pressure values  
139 measured with the PTB330 barometer for the calibration of the RS92 had a bias of 0.18  
140 hPa before the launch of the No. 5 radiosondes, we subtracted 0.18 hPa from the observed

141 pressure values of the RS92 No. 1–4 radiosondes when the data were analyzed. The  
142 balloon release detection mode was changed from automatic to manual during the  
143 MR15-04 cruise, and the starting times of the RS92 and RS41 radiosondes during the  
144 MR15-04 cruise generally appeared to differ slightly. Therefore, the measurement times of  
145 all the RS92 radiosonde data during the MR15-04 cruise were delayed by 1.7 s in the  
146 analysis.

### 147 **3 Results**

148 To facilitate comparison with the results of Jensen et al. (2016), we interpolated the RS92  
149 radiosonde profiles to the same time step as the RS41 profiles, and calculated differences  
150 between them at each 10-m vertical grid based on the RS41 radiosonde heights (Fig. 3).  
151 The vertical axis of Fig. 3 is therefore nearly equivalent to the passage of time. The biases,  
152 standard deviations, and root mean square (RMS) differences were all smaller than the  
153 combined uncertainties, except that the RMS differences of pressure above 100 hPa  
154 exceeded 0.6 hPa (Table 3). For temperature and wind speeds, the biases and RMS  
155 differences in our experiments were nearly the same as those of Jensen et al. (2016), but  
156 the differences of pressure and relative humidity were much larger in our study.

#### 157 **3.1 Pressure**

158 The pressure difference between the RS41 and RS92 radiosondes increased as the  
159 radiosondes rose to an altitude of about 5 km but averaged an almost constant 0.5–0.6 hPa

160 above that altitude (Fig. 3a). The 90th-percentile line revealed that the sensor-measured  
161 RS41 pressure was lower than the RS92 for more than 90 % of the measurements above 5  
162 km. The percentage of the pressure differences that exceeded the combined uncertainty  
163 (Table 1) was 13.7 % below 100 hPa but 50.9 % above 100 hPa. The bias of pressure  
164 causes the bias of geopotential height (Fig.3b). The height difference increased with the  
165 altitude: The median of the RS41 height was greater than that of the RS92 by  
166 approximately 35 m at an altitude of 15 km, and 100 m at 22 km.

167 We also checked the GPS-derived pressure of the RS41 radiosondes. Figure 4  
168 shows the difference between the RS92 pressure and the RS41 GPS-derived one. The use  
169 of the GPS-derived pressure reduced the bias by approximately 0.2 hPa above an altitude  
170 of 15 km, but there was still a bias of 0.4 hPa or more at most of altitudes. The median of  
171 the difference in Fig.4 was almost the same as in Fig.3a around an altitude of 5 km. The  
172 use of the GPS did not essentially improve the pressure bias. This is different from the  
173 results of Jensen et al. (2016).

### 174 **3.2 Relative humidity**

175 The median of the relative humidity differences peaked at approximately 2 %RH near  
176 10 km (Fig. 3c), a result consistent with the data of Jensen et al. (2016). The humidity  
177 difference was also large near the sea surface in our analysis. For 13.0 % of the  
178 measurements, the absolute value of the difference exceeded 4.0 %RH, which is the

179 combined uncertainty of the RS41-SGP. One noteworthy feature of Fig. 3c is that there  
180 were quite large differences of relative humidity at a height of about 17 km, although the  
181 median difference was less than 0.5 %RH. Figure 5 shows the relationship between the  
182 humidity difference and temperature for each category of relative humidity. During both the  
183 MR15-03 and MR15-04 cruises, the RS41 radiosonde ~~tended to~~ recorded a higher mean  
184 relative humidity ~~than~~ relative to the RS92 for all humidity ranges. The humidity difference  
185 peaked at around  $-40^{\circ}\text{C}$ , a pattern similar to Fig. 17 of Jensen et al. (2016). The  
186 differences were relatively small in the range of  $-50^{\circ}$  to  $-70^{\circ}\text{C}$ , but the RS41 humidity was  
187 much higher than the RS92 at temperatures below  $-80^{\circ}\text{C}$  (Fig. 5b). The atmosphere  
188 associated with temperatures below  $-80^{\circ}\text{C}$  corresponds to the tropopause in the tropics,  
189 where the greatest differences were apparent at altitudes of about 17 km (Fig. 3c).

### 190 3.3 Temperature

191 In the case of temperature, although there was a slight positive bias below an altitude of 10  
192 km, the median of the differences was within  $\pm 0.12^{\circ}\text{C}$  below an altitude of 26 km (Fig. 3d).  
193 The median exceeded  $0.5^{\circ}\text{C}$  above 27 km, but only four flights reached that height, and the  
194 large median was attributable to differences on two of the flights (No. 23 and 24). The  
195 percentages of the temperature difference that exceeded the combined uncertainty were  
196 4.0 % below 16 km and 5.9 % above 16 km. Figure 3d also shows that the standard  
197 deviation of the temperature differences was smaller at altitudes below 16 km, but there

198 were quite large standard deviations near the surface and at altitudes of about 1.3 km and  
199 5.3 km because of some outliers. The extreme temperature difference, which reached  
200 2.75°C at an altitude of 1.27 km, was observed on 10 December in the tropics (Fig. 6a).  
201 The RS92 temperature became much lower than the RS41 just after the radiosondes  
202 passed through a saturated layer into a dry layer. The greater reduction of the RS92  
203 temperature was probably due to the “wet-bulbing” effect mentioned by Jensen et al. (2016),  
204 who indicated that the sequential pulse heating method with relatively long non-heating  
205 periods may not be sufficient to eliminate icing/wetting of the RS92 sensor. A large  
206 temperature difference that was likely caused by the wet-bulbing effect was also observed  
207 in a sounding in the Arctic, although the maximum difference was less than 0.75°C (Fig.  
208 6b).

209 Figure 7 shows the cases of extreme temperature differences that contributed to the  
210 greater standard deviation and cannot be explained by the wet-bulbing effect. For the flight  
211 on 11 December (Fig. 7a), there was a large temperature discrepancy inside the saturated  
212 layer. In that case, the radiosondes were launched in heavy rain, and the ascent rate  
213 dropped to nearly zero at approximately 5.4 km, probably because of rain or snow and  
214 freezing of the balloon. Furthermore, the horizontal wind speed was less than 3.0 m s<sup>-1</sup>  
215 around this altitude. As a result, the temperature sensors were presumably not ventilated  
216 sufficiently. In the case of the flights on 1 and 3 December (Fig. 7b and 7c), the RS41  
217 temperatures were higher than the RS92 by more than 1.0°C near the surface. Because

218 the surface reference air temperatures were close to the RS92 temperatures at the lowest  
219 level, we suspect that the RS41 temperatures were too high. These large temperature  
220 differences lead to enormous discrepancies in CAPE: 864.6 J kg<sup>-1</sup> for No.22, and 1819.0 J  
221 kg<sup>-1</sup> for No.23. Yoneyama et al. (2002) have indicated that ship body heating can affect  
222 radiosonde sensors. However, that effect was restricted to within several tens of meters of  
223 the sea surface in their experiments. Although we cannot completely exclude the possibility  
224 that the temperature sensors of the two RS41 radiosondes were improperly heated by the  
225 body of the ship or direct insolation or improper handling near the surface, the reason for  
226 these large discrepancies remains unclear.

### 227 **3.4 Wind speed**

228 Vertical profiles of the wind speed differences are shown in Fig. 3e and 3f. The percentages  
229 of the differences in the zonal and meridional wind speeds that exceeded 0.5 m s<sup>-1</sup> were  
230 1.9 % and 1.5 %, respectively. Although both the zonal and meridional wind speeds agreed  
231 to within 0.5 m s<sup>-1</sup> for almost all measurements, several spikes can be seen in the standard  
232 deviations and percentiles. In half of all flights, the magnitude of the difference of the  
233 horizontal wind speed exceeded 1.0 m s<sup>-1</sup> for a brief moment. The wind speed data in our  
234 soundings were noisier than those reported by Jensen et al. (2016).

235

### 236 **4 Discussion**

#### 237 **4.1 Day-night differences**

238 Figure 8 compares the differences between daytime (10 flights) and nighttime (8 flights) for  
239 the soundings during the MR15-04 cruise. The median of the pressure difference was  
240 greater in the day than at night above an altitude of 4.5 km (Fig. 8a). The median of the  
241 nighttime differences was close to that of the daytime flights in the Arctic cruise below an  
242 altitude of 15 km, the implication being that the day-night difference might reflect some  
243 effect of solar heating.

244 The median profiles of temperature differences in the day and night were close to  
245 each other, with slightly larger differences in the night at altitudes of 5–15 km (Fig. 8b). The  
246 daytime difference became greater above approximately 24 km, a pattern similar to the  
247 results of Jensen et al. (2016). According to them, the difference in the radiation correction  
248 schemes between the RS92 and RS41 may be the dominant cause of these temperature  
249 differences, particularly at high solar elevation angles and low pressures.

250 The median of the relative humidity difference was larger during the day than at night  
251 from the surface to an altitude of 20 km and was especially large at an altitude of about 17  
252 km (Fig. 8c). The very large difference (RS41 > RS92) in relative humidity around the  
253 tropopause shown in Figs. 3c and 5b occurred in the daytime. This pattern is consistent  
254 with the results of Jauhiainen et al. (2014), who indicated that the difference was largely  
255 due to the dissimilar approaches used to compensate for the heating effect of solar  
256 radiation on the humidity sensor. Similar dry biases were reported for the RS92 radiosonde

257 with the earlier version of DigiCORA (Vömel et al., 2007; Yoneyama et al., 2008), although  
258 the dry bias was generally absent from later observations ~~processed with v3.64 software~~  
259 (Ciesielski et al., 2014; Yu et al., 2015) because the bias due to solar heating was removed  
260 by a correction scheme included in the v3.64 software or developed by Wang et al. (2013).

261 Figure 9 shows the relative difference of relative humidity in the daytime between the RS92  
262 and RS41 radiosondes. The relative difference is defined to be the relative humidity  
263 difference expressed as a percentage of the RS41 relative humidity. The relative difference  
264 was small in the lower troposphere and became greater as the radiosondes rose higher. Its  
265 median peaked at  $-36.9\%$  at an approximate altitude of 19 km. This pattern of the vertical  
266 profile of relative difference is similar to that between the RS92 radiosonde and a reference  
267 instrument shown by Vömel et al. (2007), but the values in Fig. 9 are less than half of those  
268 in Fig. 6 of Vömel et al. (2007) because the RS92 DigiCORA v3.64 and RS41 relative  
269 humidity data are already inherently better.

270 We evaluated how the differences between the two types of radiosonde affected  
271 CAPE, CIN, and PW (Table 4). CAPE tended to be larger when the RS92 was used in the  
272 nighttime. This was due to slightly higher temperature of RS92 near the surface (Fig.8b).  
273 On the other hand, in the daytime the RS41 CAPE was larger the RS92 and the RS41 CIN  
274 was smaller than the RS92. The day-night differences in the CAPE and CIN biases were  
275 caused by the difference in the humidity bias between daytime and nighttime. The  
276 near-surface humidity of the RS41 was larger than that of the RS92 in the daytime (Fig.8c).

277 The larger pressure bias in daytime (Fig.8a), which means to thicken an atmospheric layer  
278 in the RS41 observation, also may contribute to the daytime bias of CAPE. Although the  
279 bias of PW was less than 1.0 mm, and the daytime humidity difference between the RS41  
280 and RS92 affected PW, ~~but the bias of PW was less than 1.0 mm.~~ The ratio of the RS41 to  
281 the RS92 PW was dependent on solar altitude angle (Fig.9), similar to the general shape of  
282 the dependence indicated by Miloshevich et al. (2009) (their Fig.4a), suggesting that the  
283 humidity bias was mainly related with solar heating.

#### 284 **4.2 Humidity correction**

285 Figures 8c and 9 imply that a small dry bias still remains in the RS92 radiosonde  
286 observations. We attempted to correct the RS92 relative humidity obtained during the  
287 MR15-04 cruise by using the RS41 as a reference instrument. However, this is not based  
288 on an assertion that the RS42 measurements must be true values. There is no independent  
289 evidence to judge which radiosonde was more accurate. The RS41 relative humidity was  
290 larger than the RS92 at an altitude between 3-13 km (Fig.8c), suggesting that the RS41  
291 humidity also have a slight moist bias that is unrelated to the radiation correction scheme.  
292 The correction attempted in this subsection is a proposal to bridge the gap in relative  
293 humidity between the RS41 and RS92 radiosondes.

294 We used the cumulative distribution function (CDF) matching method proposed by  
295 Nuret et al. (2008) and Ciesielski et al. (2009) to make the correction. The details of this

296 method can be found in Ciesielski et al. (2009). We first created CDFs of relative humidity  
297 for the RS92 and RS41 using temperature bins of 20°C between +30° and –90°C (10 to  
298 30°C, –10 to 10°C, –30 to –10°C, –50 to –30°C, –70 to –50°C, and –90 to –70°C) using  
299 5hPa radiosonde data in 5%RH intervals. Figure 10 shows the CDFs of the RS92 and  
300 RS41 in the temperature range –90 to –70°C as an example. The frequency of lower  
301 relative humidity was greater for the RS92 in this temperature range, which includes the  
302 tropopause (Fig. 10a). We then, for example, paired the RS92 value of 27.50 %RH at the  
303 71.23th percentile with the corresponding RS41 value at this same percentile. The RS41  
304 relative humidity at the 71.23th percentile was 36.43 %RH, and the difference between  
305 36.43 %RH and 27.50 %RH (= +8.93 %RH) was the bias correction for the RS92 value of  
306 27.5 %RH. Figure 10b shows the bias correction over the entire relative humidity range for  
307 temperatures of –90 to –70°C.

308 Table 5 shows the daytime bias correction for the entire range of temperatures and  
309 relative humidities. The correction was seldom more than 5 %RH when the RS92  
310 temperature exceeded –60°C. The correction was large for RS92 radiosonde values in the  
311 range 15–50 %RH and temperatures of –80°C, with a maximum of +8.93 %RH. This  
312 pattern is similar to that of the correction table for the RS80 radiosonde in the daytime  
313 reported by Ciesielski et al. (2010) (their Fig. 7b), but the values in Table 5 are much smaller.  
314 We corrected the daytime RS92 relative humidity values obtained during the MR15-04  
315 cruise using Table 5. The correction value for an arbitrary RS92 measurement can be

316 obtained by linear two-dimensional interpolation using Table 5 and the RS92 temperature  
317 and relative humidity. Figure 11 shows median profiles of the differences between the RS92  
318 and RS41 radiosondes before and after the correction. Although the median of the  
319 magnitude of the differences still exceeded 2.0 %RH around 120, 150, and 560 hPa, most  
320 of the medians were within  $\pm 1.0$  %RH. The mean of the relative humidity difference of the  
321 5hPa interval data was  $-2.02$  %RH if no correction was made; this difference was reduced  
322 to  $-0.01$  %RH after the correction.

## 323 **5 Conclusions**

324 To examine differences between the RS41 and RS92 radiosondes, a total of 36  
325 twin-radiosonde flights were performed over the Arctic Ocean, Bering Sea, northwestern  
326 Pacific Ocean, and the tropical Indian Ocean during two cruises of R/V *Mirai* in 2015. We  
327 used the model RS41-SGP radiosonde, which has a pressure sensor, unlike previous  
328 studies that used the RS41-SG, which has no pressure sensor.

329 The biases, standard deviations, and RMS of the differences between the RS41 and  
330 RS92 over all flights and heights were smaller than the nominal combined uncertainties of  
331 the RS41, except that the RMS differences of pressure above 100 hPa exceeded 0.6 hPa.  
332 Whereas the biases and the RMS differences of temperature and wind speeds were close  
333 to those reported by Jensen et al. (2016), the differences of pressure and relative humidity  
334 were greater in our experiments. The pressure difference increased as the radiosondes

335 rose higher; the median and mean were 0.5–0.6 hPa at altitudes above 5 km. This pressure  
336 difference corresponded to a geopotential height difference of more than 35 m above an  
337 altitude of 15 km. A comparison between daytime and nighttime flights in the tropics  
338 revealed that the pressure difference was systematically larger in the day than at night at  
339 altitudes above 4.5 km, the suggestion being that there was some effect of solar heating on  
340 the pressure measurements. The exact reason, however, is unclear.

341         The RS41 and RS92 temperature measurements in general agreed better than the  
342 combined uncertainties, but there were some noteworthy exceptions. One possible reason  
343 for the noteworthy discrepancies is the wet-bulbing effect described by Jensen et al. (2016).  
344 In a dry layer just above a saturated layer, the RS92 temperature sensor was cooled too  
345 much by evaporation. The RS41 temperature appeared to be less sensitive to this  
346 wet-bulbing effect. This phenomenon was confirmed in both the tropics and Arctic. During  
347 heavy rain and weak wind conditions, the stagnation of the balloon probably suppressed  
348 the ventilation around the temperature sensors, the result being an extreme temperature  
349 difference.

350         The median of the relative humidity differences at all altitudes was only a little more  
351 than 2 %RH. However, there were quite large differences at an altitude of about 17 km.  
352 These large differences occurred in the daytime around the tropical tropopause, where the  
353 temperature was below  $-80^{\circ}\text{C}$ . The reason for this dry bias may be that there was some  
354 remnant of the error of the RS92 radiosonde solar radiation correction. The differences in

355 humidity affected the calculation of CAPE, CIN, and PW, and we confirmed the day-night  
356 difference of these variables. We attempted to correct the RS92 relative humidity data  
357 obtained in the daytime during the MR15-04 cruise by using the CDF matching method,  
358 and the corrected RS92 relative humidity agreed well with the RS41 values.

359 Our results showed that measurements with the RS41 radiosonde satisfied the  
360 performance specifications of the manufacturer in most cases over both the tropical and  
361 polar oceans. The RS41 temperature and humidity sensors appeared to be unaffected by  
362 the solar radiation correction error and the wet-bulbing effect. Some concerns, however,  
363 remain. Specifically, the reasons for the pressure bias in the upper layer and the two cases  
364 of extreme temperature discrepancies that occurred below an altitude of several hundred  
365 meters are unknown. Further experiments will be necessary to address these issues, and  
366 users should be cognizant of these concerns.

## 367 **6 Data availability**

368 The sounding dataset and the ship-observed surface meteorology are expected to be  
369 released just two years after the cruises (October 2017 for the MR15-03, and December  
370 2017 for the MR15-04) from the website of the Data Research System for Whole Cruise  
371 Information (DARWIN) in JAMSTEC (<http://www.godac.jamstec.go.jp/darwin/e>) in accord  
372 with the cruise data policy of JAMSTEC.

373

374 Author contributions

375 All co-authors contributed to designing the experiments and preparing for the observation  
376 cruises. Y. Kawai, M. Katsumata, and K. Oshima participated in the R/V *Mirai* cruises and  
377 carried out the radiosonde soundings. K. Oshima reprocessed the RS92 data. M.  
378 Katsumata calculated CAPE, CIN, and PW. Y. Kawai mainly analyzed the data and  
379 prepared the manuscript with contributions from all co-authors.

380

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388 **References**

389 Bodeker, G. E., Bojinski, S., Cimini, C., Dirksen, R. J., Haeffelin, M., Hannigan, J. W., Hurst,  
390 D. F., Leblanc, T., Madonna, F., Maturilli, M., Mikalsen, A. C., Philpona, R., Reale, T.,  
391 Siedel, D. J., Tan, D. G. H., Thorne, P. W., Vömel, H., and Wang, J.: Reference upper-air  
392 observations for climate: From concept to reality, B. Am. Meteorol. Soc., 97, 123-135,

393 doi:10.1175/BAMS-D-14-00072.1, 2016.

394 Ciesielski, P. E., Johnson, R. H., and Wang, J: Correction of humidity biases in Vaisala  
395 RS80-H sondes during NAME, *J. Atmos. Ocean. Tech.*, 26, 1763-1780,  
396 doi:10.1175/2009JTECHA1222.1, 2009.

397 Ciesielski, P. E., Chang, W.-M., Huang, S. -C., Johnson, R. H., Jou, B. J. -D., Lee, W. -C.,  
398 Lin, P. -H., Liu, C. -H., and Wang, J.: Quality-controlled upper-air sounding dataset for  
399 TIMREX/SoWMEX: Development and corrections, *J. Atmos. Ocean. Tech.*, 27,  
400 1802-1821, doi:10.1175/2010JTECHA1481.1, 2010.

401 Ciesielski, P. E., Yu, H., Johnson, R. H., Yoneyama, K., Katsumata, M., Long, C. N., Wang,  
402 J., Loehrer, S. M., Young, K., Williams, S. F., Brown, W., Braun, J., and Van Hove, T.:  
403 Quality-controlled upper-air sounding dataset for DYNAMO/CINDY/AMIE: Development  
404 and corrections, *J. Atmos. Ocean. Tech.*, 31, 741-764, doi:10.1175/JTECH-D-13-00165.1,  
405 2014.

406 Fujita, M., Kimura, F., Yoneyama, K., and Yoshizaki, M.: Verification of precipitable water  
407 vapor estimated from shipborne GPS measurements, *Geophys. Res. Lett.*, 35, L13803,  
408 doi:10.1029/2008GL033764, 2008

409 Inoue, J., Enomoto, T., and Hori, M. E.: The impact of radiosonde data over the ice-free  
410 Arctic Ocean on the atmospheric circulation in the Northern Hemisphere, *Geophys. Res.*  
411 *Lett.*, 40, 864-869, doi:10.1002/grl.50207, 2013.

412 Inoue, J., Yamazaki, A., Ono, J., Dethloff, K., Maturilli, M., Neuber, R., Edwards, P., and

413 Yamaguchi, H.: Additional Arctic observations improve weather and sea-ice forecasts for  
414 the Northern Sea Route, *Sci. Rep.*, 5, 16868, doi:10.1038/srep16868, 2015.

415 Jauhiainen, H., Survo, P., Lehtinen, R., and Lentonen, J.: Radiosonde RS41 and RS92 key  
416 differences and comparison test results in different locations and climates. *TECO-2014,*  
417 *WMO Technical Conference on Meteorological and Environmental Instruments and*  
418 *Methods of Observations, Saint Petersburg, Russian Federation, 7–9 July 2014, P3(16),*  
419 *2014.*

420 Jensen, P. M., Holdridge, D. J., Survo, P., Lehtinen, R., Baxter, S., Toto, T., and Johnson, K.  
421 L.: Comparison of Vaisala radiosondes RS41 and RS92 at the ARM Southern Great  
422 Plains site, *Atmos. Meas. Tech.*, 9, 3115-3129, doi:10.5194/amt-9-3115-2016, 2016.

423 Katsumata, M., and coauthors: R/V Mirai Cruise Report MR15-04, *Cruise Rep.*, Japan  
424 Agency for Marine-Earth Science and Technology, Yokosuka, Japan, 241.pp, 2015.  
425 (Available from  
426 [http://www.godac.jamstec.go.jp/catalog/data/doc\\_catalog/media/MR15-04\\_all.pdf](http://www.godac.jamstec.go.jp/catalog/data/doc_catalog/media/MR15-04_all.pdf))

427 Kawai, Y., Tomita, H., Cronin, M. F., and Bond, N. A.: Atmospheric pressure response to  
428 mesoscale sea surface temperature variations in the Kuroshio Extension: In situ  
429 evidence, *J. Geophys. Res. Atmos.*, 119, 8015-8031. doi:10.1002/2013JD021126, 2014.

430 Maturilli, M., and Kayser, M.: Arctic warming, moisture increase and circulation changes  
431 observed in the Ny-Ålesund homogenized radiosonde record, *Theor. Appl. Climatol.*,  
432 doi:10.1007/s00704-016-1864-0, 2016.

433 Miloshevich, L. M., Vömel, H., Whiteman, D. N., and Leblanc, T.: Accuracy of assessment  
434 and correction of Vaisala RS92 radiosonde water vapor measurement, J. Geophys. Res.,  
435 114, D11305, doi:10.1029/2008JD011565, 2009.

436 Minobe, S., and Takebayashi, S.: Diurnal precipitation and high cloud frequency variability  
437 over the Gulf Stream and over the Kuroshio, *Clim. Dyn.*, 44, 2079-2095,  
438 doi:10.1007/s00382-014-2245-y, 2015.

439 Motl, M.: Vaisala RS41 trial in the Czech Republic, *Vaisala News*, 192, 14-17, 2014.

440 Nishino, S., and coauthors: R/V Mirai Cruise Report MR15-03, Cruise Rep., Japan Agency  
441 for Marine-Earth Science and Technology, Yokosuka, Japan, 297.pp, 2015. (Available  
442 from  
443 [http://www.godac.jamstec.go.jp/catalog/data/doc\\_catalog/media/MR15-03\\_leg1\\_all.pdf](http://www.godac.jamstec.go.jp/catalog/data/doc_catalog/media/MR15-03_leg1_all.pdf))

444 Nuret, M., Lafore, J.-P., Bock, O., Guichard, F., Agusti-Panareda, A., N'Gamini, J.-B., and  
445 Redelsperger, J.-L.: Correction of humidity bias for Vaisala RS80-A sondes during the  
446 AMMA 2006 observing period, *J. Atmos. Ocean. Tech.*, 25, 2152-2158,  
447 doi:10.1175/2008JTECHA1103.1, 2008.

448 Thorne, P. W., Parker, D. E., Tett, S. F. B., Jones, P. D., McCarthy, M., Coleman, H., and  
449 Brohan, P.: Revisiting radiosonde upper air temperatures from 1958 to 2002, *J. Geophys.*  
450 *Res.*, 110, D18105, doi:10.1029/2004JD005753, 2005.

451 Vömel, H., Selkirk, H., Miloshevich, L., Valverde-Canossa, J., Valdés, J., Kyrö, E., Kivi, R.,  
452 Stolz, W., Peng, G., and Diaz, J. A.: Radiation dry bias of the Vaisala RS92 humidity

453 sensor, *J. Atmos. Ocean. Tech.*, 24, 953-963, doi:10.1175/JTECH2019.1, 2007.

454 Wang, J., Zhang, L., Dai, A., Immler, F., Sommer, M., and Vömel, H.: Radiation dry bias  
455 correction of Vaisala RS92 humidity data and its impacts on historical radiosonde data, *J.*  
456 *Atmos. Ocean. Tech.*, 30, 197-214, doi:10.1175/ JTECH-D-12-00113.1, 2013.

457 Yang, G.-Y., and Slingo, J.: The diurnal cycle in the tropics, *Mon. Wea. Rev.*, 129, 784-801,  
458 doi:10.1175/1520-0493(2001)129<0784:TDCITT>2.0.CO;2, 2001.

459 Yoneyama, K., Hanyu, M., Sueyoshi, S., Yoshiura, F., and Katsumata, M.: Radiosonde  
460 observation from the ship in the tropical region, *Report of Japan Marine Science and*  
461 *Technology Center*, 45, 31-39, 2002. (Available from  
462 [http://www.jamstec.go.jp/res/ress/yoneyamak/PDFs/Yoneyama-etal\\_2002\\_JAMSTECR.](http://www.jamstec.go.jp/res/ress/yoneyamak/PDFs/Yoneyama-etal_2002_JAMSTECR.pdf)  
463 pdf)

464 Yoneyama, K., Fujita, M., Sato, N., Fujiwara, M., Inai, Y., and Hasebe, F.: Correction for  
465 radiation dry bias found in RS92 radiosonde data during the MISMO field experiment,  
466 *SOLA*, 4, 13-16, doi:10.2151/sola.2008-004, 2008.

467 Yu, H., Ciesielski, P. E., Wang, J., Kuo, H.-C., Vömel, H., and Dirksen, R.: Evaluation of  
468 humidity correction methods for Vaisala RS92 tropical sounding data, *J. Atmos. Ocean.*  
469 *Tech.*, 32, 397–411, doi:10.1175/JTECH-D-14-00166.1, 2015.

470

471

472 **Table 1.** Nominal accuracies of the radiosondes according to the manufacturer.

473

		RS41-SGP	RS92-SGPD
Weight		113 g	280 g
Combined uncertainty in sounding (2-sigma confidence level (95.5 %) cumulative uncertainty)	Temperature	0.3°C < 16 km 0.4°C > 16 km	0.5°C
	Relative humidity	4 %RH	5 %RH
	Pressure		1.0 > 100 hPa 0.6 < 100 hPa
Reproducibility in sounding (standard deviation of differences in twin soundings)	Temperature <sup>a</sup>	0.15°C > 100 hPa 0.30°C < 100 hPa	0.2°C > 100 hPa 0.3°C 100–20 hPa 0.5°C < 20 hPa
	Relative humidity <sup>a</sup>		2 %RH
	Pressure		0.5 > 100 hPa 0.3 < 100 hPa
	Wind speed		0.15 m/s
Wind direction <sup>b</sup>		2°	

474 <sup>a</sup> Ascent rate above 3 m s<sup>-1</sup>

475 <sup>b</sup> Wind speed above 3 m s<sup>-1</sup>

476 **Table 2.** Date, position (latitude and longitude), surface meteorological state (pressure, temperature, relative humidity, wind direction, and  
 477 wind speed), CAPE, CIN, and PW when each twin-radiosonde was launched. Line under UTC time denotes nighttime.  
 478

Cruise	No.	Date	Time (UTC)	Time (LT)	Lat. (°N)	Lon. (°E)	Pressure (hPa)	Temp. (°C)	RH (%)	Wind dir. (°)	Wind speed (m s <sup>-1</sup> )	RS41 Maximum height (m)	Mean ascent rate (m s <sup>-1</sup> )	RS41 CAPE (J kg <sup>-1</sup> )	RS41 CIN (J kg <sup>-1</sup> )	RS41 PW (mm)
15-03	1	27 Aug.	23:30	9:30	40.17	149.94	1011.7	15.9	69	23	7.1	26,734	4.06	0	NA	14.3
	2	28 Aug.	23:30	9:30	42.42	153.41	1010.7	14.0	70	306	11.2	23,328	4.42	0.6	1.5	11.3
	3	29 Aug.	23:30	9:30	44.83	157.19	1004.2	12.1	93	289	11.6	21,607	4.45	0	NA	31.2
	4	31 Aug.	23:32	10:32	49.93	165.75	999.6	10.9	93	275	5.6	19,380	4.74	3.8	4.5	24.0
	5	2 Sep.	23:30	11:30	55.49	175.34	1000.4	10.3	97	155	7.8	13,617	4.68	3.7	0	22.9
	6	4 Sep.	23:32	11:32	63.43	-172.92	1008.6	9.0	81	294	3.6	23,554	5.06	0.2	0.7	19.9
	7	8 Sep.	5:30	18:30	71.05	-166.94	1015.9	1.3	83	342	6.7	22,872	5.22	2.6	0.4	8.4
	8	12 Sep.	23:30	13:30	72.48	-156.29	1009.8	-0.1	96	91	9.3	21,243	5.36	0.1	0	12.8
	9	16 Sep.	5:30	19:30	72.34	-156.18	1015.1	-1.7	86	46	5.4	22,298	5.33	0	0.2	7.7
	10	24 Sep.	23:31	12:31	73.21	-157.80	993.2	0.7	95	170	9.8	25,309	5.12	0	0	13.1
	11	28 Sep.	17:31	6:31	74.37	-166.57	987.8	-1.4	92	164	8.6	23,291	5.18	9.4	0.3	6.8
	12	28 Sep.	23:30	12:30	74.47	-168.18	982.0	-0.9	70	167	11.2	22,811	5.26	0	NA	6.4
	13	29 Sep.	5:30	18:30	74.00	-168.76	979.9	-2.3	80	210	9.9	19,338	5.25	47.8	1.1	4.6
	14	30 Sep.	<u>11:30</u>	0:30	70.38	-168.76	993.2	-2.1	89	282	7.0	19,897	5.16	0	NA	5.1
	15	30 Sep.	23:30	12:30	68.06	-168.83	1008.6	1.8	69	296	7.1	22,613	5.17	25.2	1.0	5.3
	16	4 Oct.	23:30	12:30	60.74	-167.78	1011.4	8.1	100	186	14.3	19,498	4.77	0.3	0	20.6
	17	11 Oct.	23:30	11:30	53.64	178.82	1006.8	6.3	90	10	3.8	25,051	5.17	0.7	0.4	14.5

	18	17 Oct.	23:30	9:30	41.79	154.88	1019.8	12.0	64	177	2.9	25,928	5.21	0	NA	9.2
15-04	19	10 Nov.	5:38	14:38	23.57	136.76	1011.6	26.7	83	357	3.3	25,395	3.78	1309.0	5.6	42.1
	20	11 Nov.	5:39	14:39	19.21	134.81	1011.6	28.0	81	72	8.1	26,589	4.04	1558.5	4.6	42.6
	21	30 Nov.	8:29	15:29	-4.08	101.89	1006.2	28.5	75	202	4.2	22,184	3.95	630.9	22.8	59.8
	22	1 Dec.	5:30	12:30	-4.05	101.89	1008.1	28.4	79	298	2.7	26,510	4.27	2228.8	3.4	60.4
	23	3 Dec.	5:29	12:29	-4.07	101.89	1008.5	28.0	82	275	4.2	28,867	4.35	3008.1	3.7	63.0
	24	5 Dec.	2:30	9:30	-4.07	101.88	1009.5	26.0	92	254	1.9	28,016	4.07	645.1	15.9	64.6
	25	5 Dec.	<u>17:45</u>	0:45	-4.09	101.89	1008.6	27.4	86	80	1.3	26,822	3.98	1531.4	1.0	64.7
	26	6 Dec.	<u>20:26</u>	3:26	-4.07	101.91	1005.8	27.9	85	139	6.2	27,518	3.97	1393.3	23.0	63.9
	27	8 Dec.	<u>14:29</u>	21:29	-4.08	101.89	1010.5	27.9	82	126	3.0	26,965	4.26	1357.2	0.8	63.4
	28	9 Dec.	2:28	9:28	-4.05	101.89	1010.0	27.4	81	298	1.9	27,123	4.32	979.2	9.6	66.8
	29	10 Dec.	<u>17:27</u>	0:27	-4.04	101.89	1009.1	27.0	87	6	1.4	24,650	4.40	1324.6	0.3	63.3
	30	11 Dec.	<u>14:20</u>	21:20	-4.05	101.87	1008.0	25.5	98	5	10.3	15,050	6.62	162.5	86.9	78.4
	31	13 Dec.	<u>20:28</u>	3:28	-4.06	101.89	1006.1	28.1	77	324	6.2	20,798	3.57	887.1	12.5	60.0
	32	15 Dec.	5:28	12:28	-4.05	101.90	1007.9	27.6	82	339	8.6	23,698	4.25	1229.5	1.5	61.5
	33	16 Dec.	2:50	9:50	-4.06	101.89	1010.3	25.0	94	310	5.2	4,803	2.48	0	0.1	54.3
	34	16 Dec.	<u>14:22</u>	21:22	-4.06	101.89	1010.1	26.2	90	11	7.9	21,629	4.48	1030.4	0.4	57.6
	35	17 Dec.	5:28	12:28	-4.05	101.90	1008.2	28.2	72	278	1.4	21,607	3.61	379.5	24.1	48.2
	36	17 Dec.	<u>20:27</u>	3:27	-5.17	101.41	1007.2	28.2	79	303	6.0	24,944	3.70	2035.6	2.7	59.8

479 **Table 3.** Biases, RMS differences, and standard deviations (SDs) of the variables between  
 480 the RS92 and RS41 radiosondes. The bias is the mean of RS92 – RS41 differences.  
 481

Variable	Total		MR15-03 (Subarctic – Arctic)		MR15-04 (Subtropics – Tropics)	
	Bias	RMS SD	Bias	RMS SD	Bias	RMS SD
	Temperature (°C)	+0.04	0.17	+0.01	0.15	+0.06
$P_{RS92} > 100\text{hPa}$		0.17		0.15		0.18
Temperature (°C)	-0.01	0.22	-0.10	0.27	+0.05	0.18
$P_{RS92} < 100\text{hPa}$		0.22		0.25		0.17
Pressure (hPa)	+0.52	0.67	+0.41	0.58	+0.64	0.76
$P_{RS92} > 100\text{hPa}$		0.42		0.40		0.41
Pressure (hPa)	+0.55	0.67	+0.57	0.61	+0.53	0.71
$P_{RS92} < 100\text{hPa}$		0.38		0.21		0.47
Relative humidity (%RH)	-0.89	3.14	-0.50	2.14	-1.26	3.86
		3.01		2.08		3.64
Zonal wind speed (m s <sup>-1</sup> )	-0.0017	0.18	+0.0027	0.17	-0.0059	0.18
		0.18		0.17		0.18
Meridional wind speed (m s <sup>-1</sup> )	-0.0051	0.17	+0.0104	0.18	-0.0199	0.16
		0.17		0.18		0.15

482

483

484 Table 4. Biases and standard deviations of CAPE, CIN and PW between the RS92 and  
 485 RS41 radiosondes. The bias is the mean of RS92 – RS41 differences. Values in  
 486 parentheses are the statistics without the two outliers shown in Fig. 7b-c (Flight No. 22  
 487 and No. 23).

488

	MR15-03			MR15-04 Daytime			MR15-04 Nighttime		
	RS41 Mean	Bias	SD	RS41 Mean	Bias	SD	RS41 Mean	Bias	SD
CAPE (J kg <sup>-1</sup> )	5.3	-0.9	1.8	1196.9 (841.5)	-331.7 (-75.4)	614.7 (222.4)	1215.3	111.1	94.8
CIN (J kg <sup>-1</sup> )	0.8	0.8	1.9	9.2 (10.6)	1.1 (1.0)	4.4 (5.0)	16.0	-0.2	1.3
PW (mm)	13.2	-0.2	0.3	56.3 (55.0)	-0.9 (-0.6)	1.1 (1.0)	63.9	0.1	0.5

489

490

491 **Table 5.** Bias correction table of relative humidity that was created by matching the CDFs  
 492 from the RS92 data to the RS41 data (%RH) based on the daytime data obtained during  
 493 the MR15-04 cruise.

494

	$\leq -80^{\circ}\text{C}$	$-60^{\circ}\text{C}$	$-40^{\circ}\text{C}$	$-20^{\circ}\text{C}$	$0^{\circ}\text{C}$	$\geq 20^{\circ}\text{C}$
2.5 %RH	1.84	0	-0.42	0	0	0
7.5	0.50	2.35	0.50	0.25	0.36	0
12.5	4.12	2.14	3.24	1.15	0.79	0
17.5	6.47	3.13	2.31	1.43	1.00	0
22.5	7.14	3.33	2.86	1.67	1.67	0
27.5	8.93	1.67	4.09	2.50	1.82	0
32.5	8.13	2.50	4.23	3.00	0.88	0
37.5	7.31	2.50	4.33	2.92	4.17	1.67
42.5	6.25	4.06	4.38	2.73	3.75	0.63
47.5	7.50	5.00	2.50	2.78	2.08	4.17
52.5	5.00	5.50	4.17	2.65	1.67	2.14
57.5	0	4.50	5.00	4.09	2.00	1.25
62.5	0	5.00	2.22	5.00	2.76	2.50
67.5	0	5.00	0	4.44	0.80	0.49
72.5	0	0	0	3.27	1.60	1.25
77.5	0	0	0	3.38	1.35	1.44
82.5	0	0	0	2.50	1.45	1.36
87.5	0	0	0	3.00	1.73	0.91
92.5	0	0	0	2.50	0.90	0.56
97.5	0	0	0	0	0	0

495

496

497

## Figure Captions

498 **Figure 1.** Positions of the twin-radiosonde launches during the (a) MR15-03 cruise, and (b)  
499 MR15-04 cruise. (c) Time-latitude diagram of the launches. Black and red dots represent  
500 daytime and nighttime soundings, respectively.

501

502 **Figure 2.** Photographs of (upper) the RS92 and RS41 radiosondes directly attached to  
503 each other and (lower) a launch on R/V *Mirai*.

504

505 **Figure 3.** Vertical profiles of the median (black), 25–75th percentile (green), 10–90th  
506 percentile (gray), and mean  $\pm$  standard deviation (cyan) of all differences between the  
507 RS92 and RS41 observations (RS92 – RS41) for (a) pressure, (b) geopotential height,  
508 (c) relative humidity, (d) temperature, (e) zonal wind, and (f) meridional wind.

509

510 **Figure 4.** As in Fig.3a, but for between the RS41 GPS-derived and RS92 pressures (RS92  
511 – RS41).

512

513 **Figure 5.** Mean difference in relative humidity between the RS92 and RS41 radiosondes  
514 (RS92 – RS41) as a function of the RS41 temperature for relative humidity ranges of  
515 0–20 % (blue), 20–40 % (red), 40–60 % (green), 60–100 % (black), and 0–100 % (gray).

516

517 **Figure 6.** Vertical profiles of the RS41 temperature (red), RS92 temperature (blue), RS41  
518 relative humidity (magenta), and RS92 relative humidity (cyan). (a) Flight No. 29  
519 launched at 1727 UTC on 10 December 2015 in the tropics, and (b) Flight No. 9 launched  
520 at 0530 UTC on 16 September 2015 in the Arctic.

521

522 **Figure 7.** As Fig. 6, but for (a) Flight No. 30 launched at 1420 UTC on 11 December 2015,  
523 (b) Flight No. 22 launched at 0530 UTC on 1 December 2015, and (c) Flight No. 23  
524 launched at 0529 UTC on 3 December 2015. All launches in the tropics.

525

526 **Figure 8.** Differences between the RS92 and RS41 radiosonde (RS92 – RS41) results for  
527 daytime (blue) and nighttime (red) flights during the MR15-04 cruise for (a) pressure, (b)  
528 temperature, and (c) relative humidity.

529

530 **Figure 9.** The ratio of the RS41 to the RS92 PW as a function of solar altitude angle. Blue  
531 and red dots represent soundings in the MR15-03 and MR15-04 cruises, respectively.

532

533 **Figure 910.** Relative difference between the RS92 and RS41 relative humidity obtained  
534 during the daytime on the MR15-04 cruise (blue dots, %). Relative difference is defined  
535 as the relative humidity difference expressed as a percentage of the RS41 relative  
536 humidity. Green line denotes the median of the relative difference. Lower panel shows an

537 enlargement of part of the upper panel.

538

539 **Figure 4011.** (a) CDFs of relative humidity for the RS92 (bold dashed line) and RS41 (bold  
540 solid line) data in the temperature range of  $-90$  to  $-70^{\circ}\text{C}$ . The daytime data obtained  
541 during the MR15-04 cruise were used. Thin solid lines illustrate the CDF-matching  
542 technique (see text). (b) Bias correction of relative humidity for the same temperature  
543 range.

544

545 **Figure 4112.** Medians of the relative humidity difference between the RS92 and RS41  
546 radiosondes obtained during the daytime on the MR15-04 cruise. Blue and black lines  
547 show the profiles before and after the bias correction of the RS92 data.

548

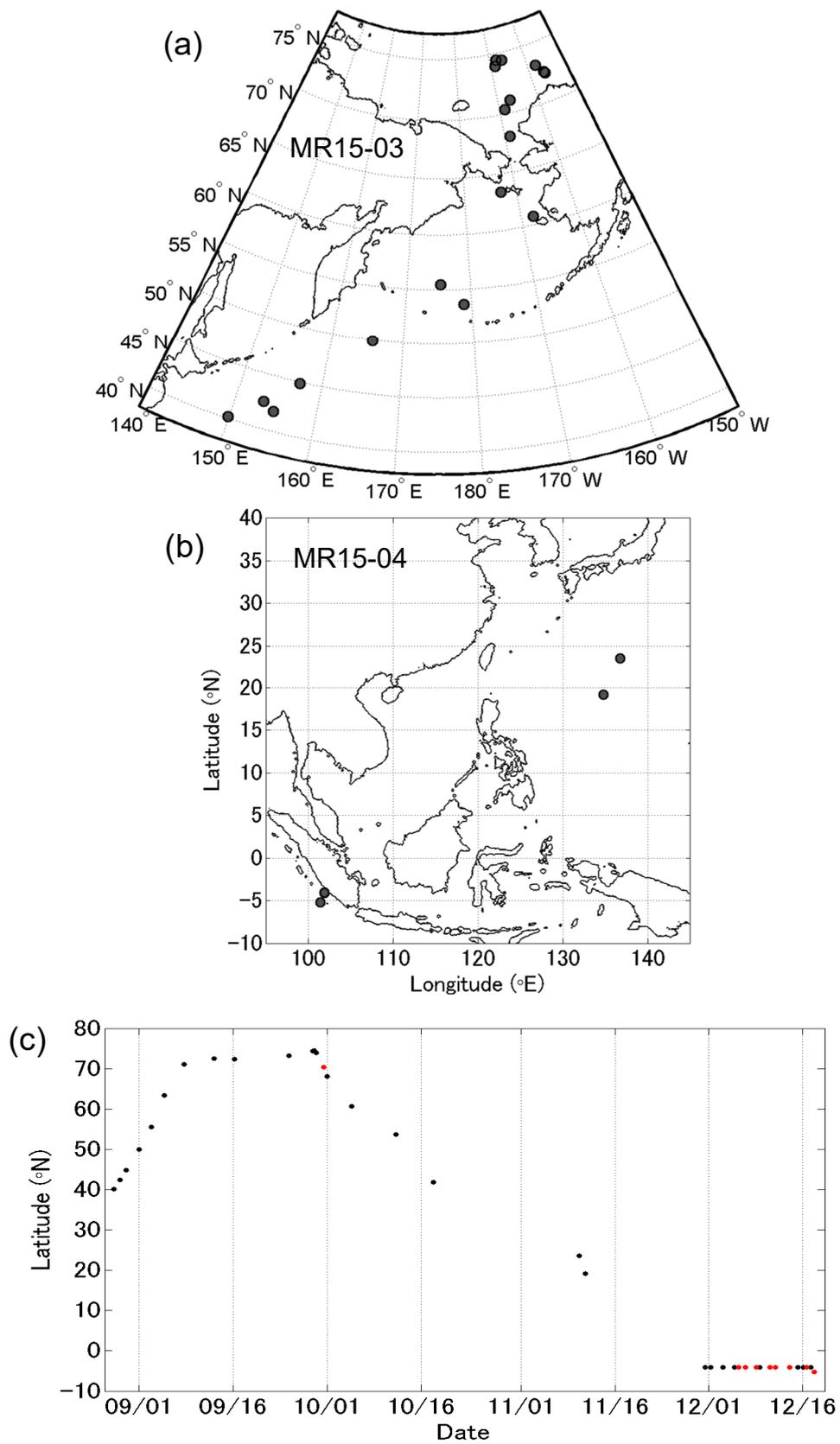


Fig. 1. Positions of the twin-radiosonde launches during the (a) MR15-03 cruise, and (b) MR15-04 cruise. (c) Time-latitude diagram of the launches. Black and red dots represent daytime and nighttime soundings, respectively.



Fig.2. Photographs of (upper) the RS92 and RS41 radiosondes directly attached to each other and (lower) a launch on R/V *Mirai*.

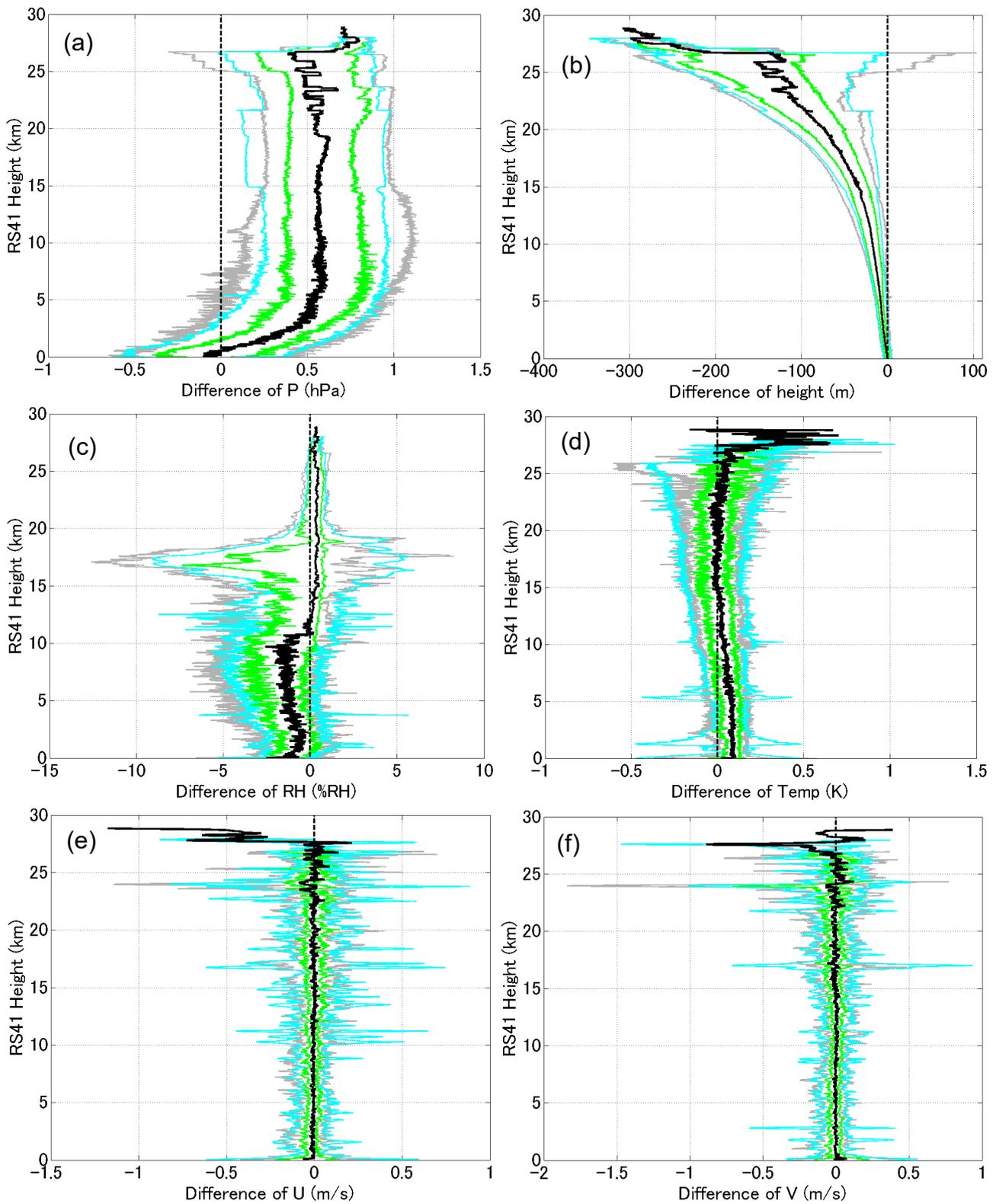


Fig.3. Vertical profiles of the median (black), 25–75th percentile (green), 10–90th percentile (gray), and mean  $\pm$  standard deviation (cyan) of all differences between the RS92 and RS41 observations (RS92 – RS41) for (a) pressure, (b) geopotential height, (c) relative humidity, (d) temperature, (e) zonal wind, and (f) meridional wind.

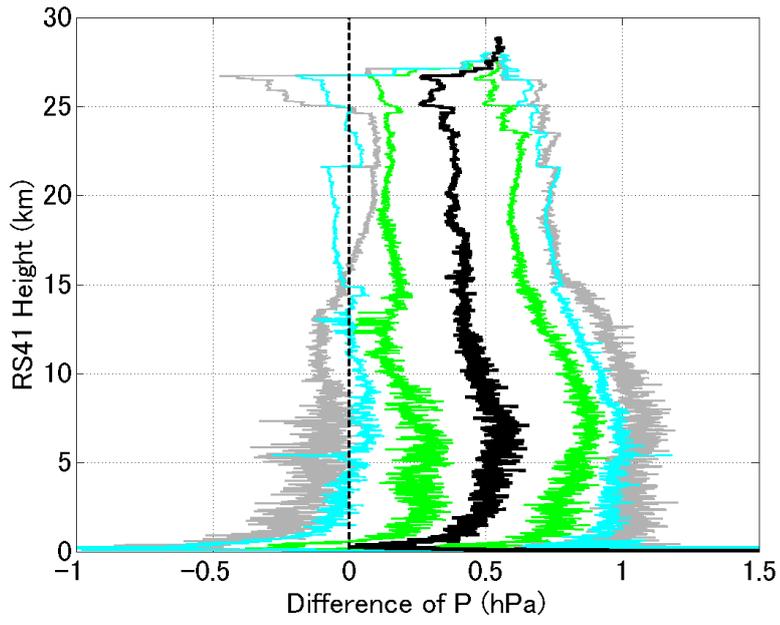


Fig.4. As in Fig.3a, but for between the RS41 GPS-derived and RS92 pressures (RS92 – RS41).

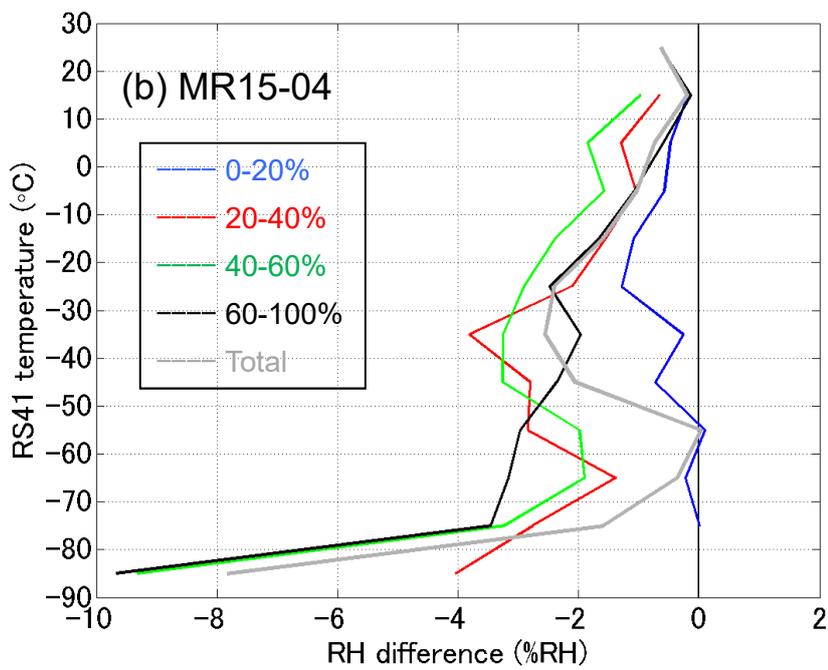
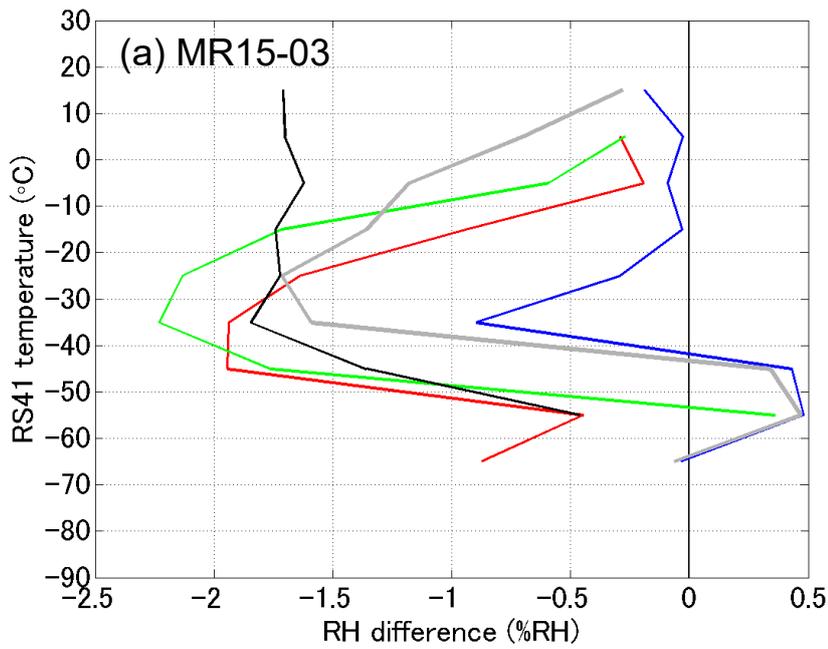


Fig.5. Mean difference in relative humidity between the RS92 and RS41 radiosondes (RS92 – RS41) as a function of the RS41 temperature for relative humidity ranges of 0–20 % (blue), 20–40 % (red), 40–60 % (green), 60–100 % (black), and 0–100 % (gray).

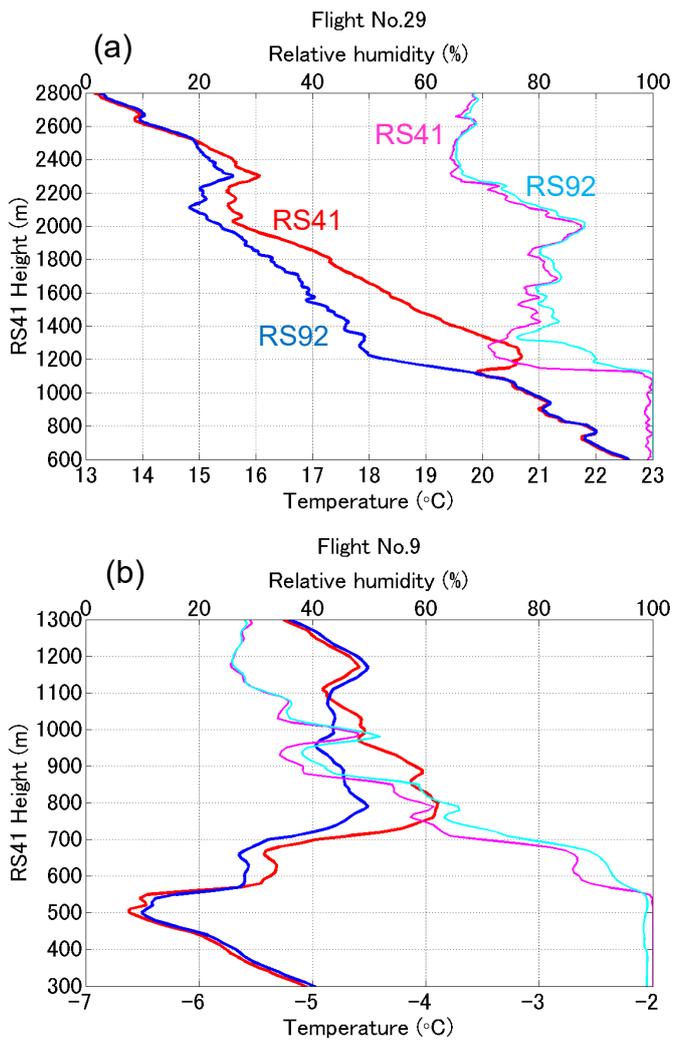


Fig.6. Vertical profiles of the RS41 temperature (red), RS92 temperature (blue), RS41 relative humidity (magenta), and RS92 relative humidity (cyan). (a) Flight No. 29 launched at 1727 UTC on 10 December 2015 in the tropics, and (b) Flight No. 9 launched at 0530 UTC on 16 September 2015 in the Arctic.

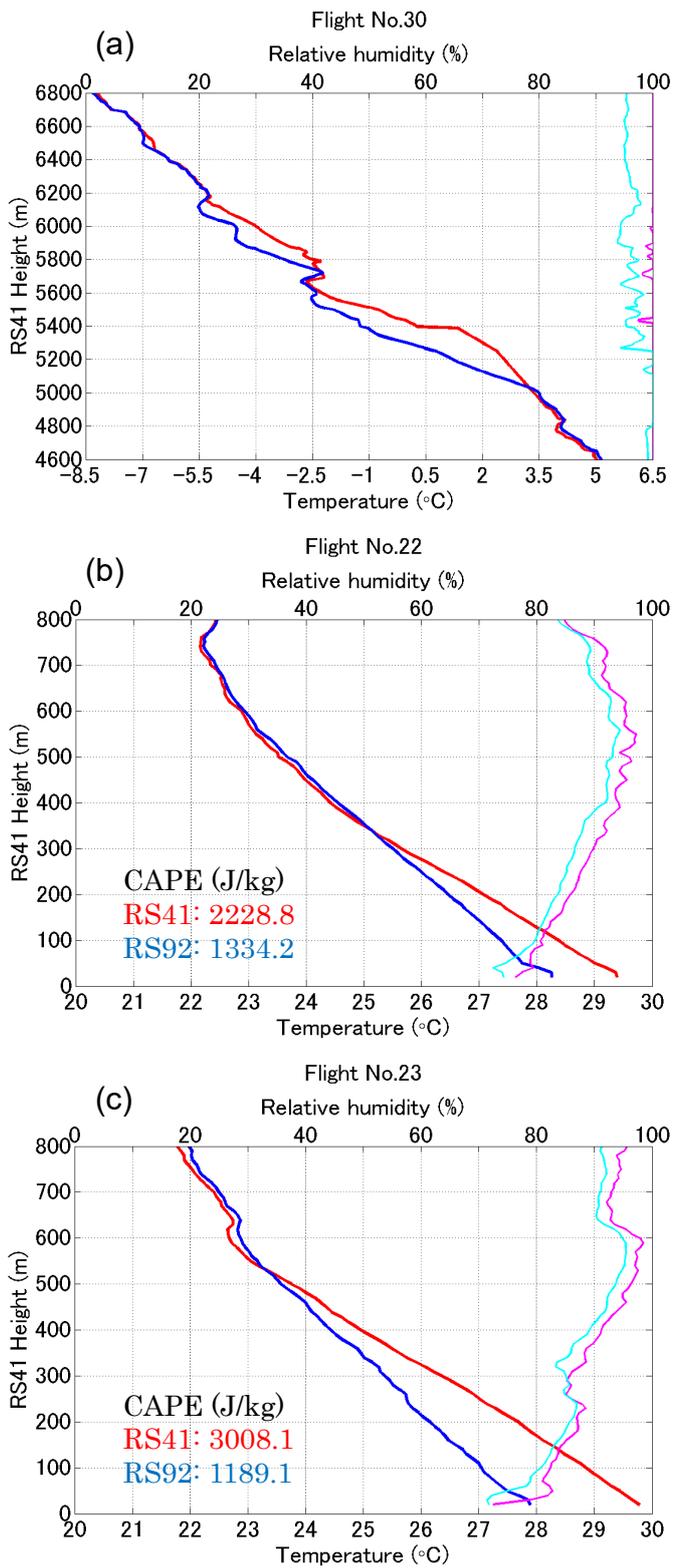


Fig.7. As in Fig. 6, but for (a) Flight No. 30 launched at 1420 UTC on 11 December 2015, (b) Flight No. 22 launched at 0530 UTC on 1 December 2015, and (c) Flight No. 23 launched at 0529 UTC on 3 December 2015. All launches in the tropics.

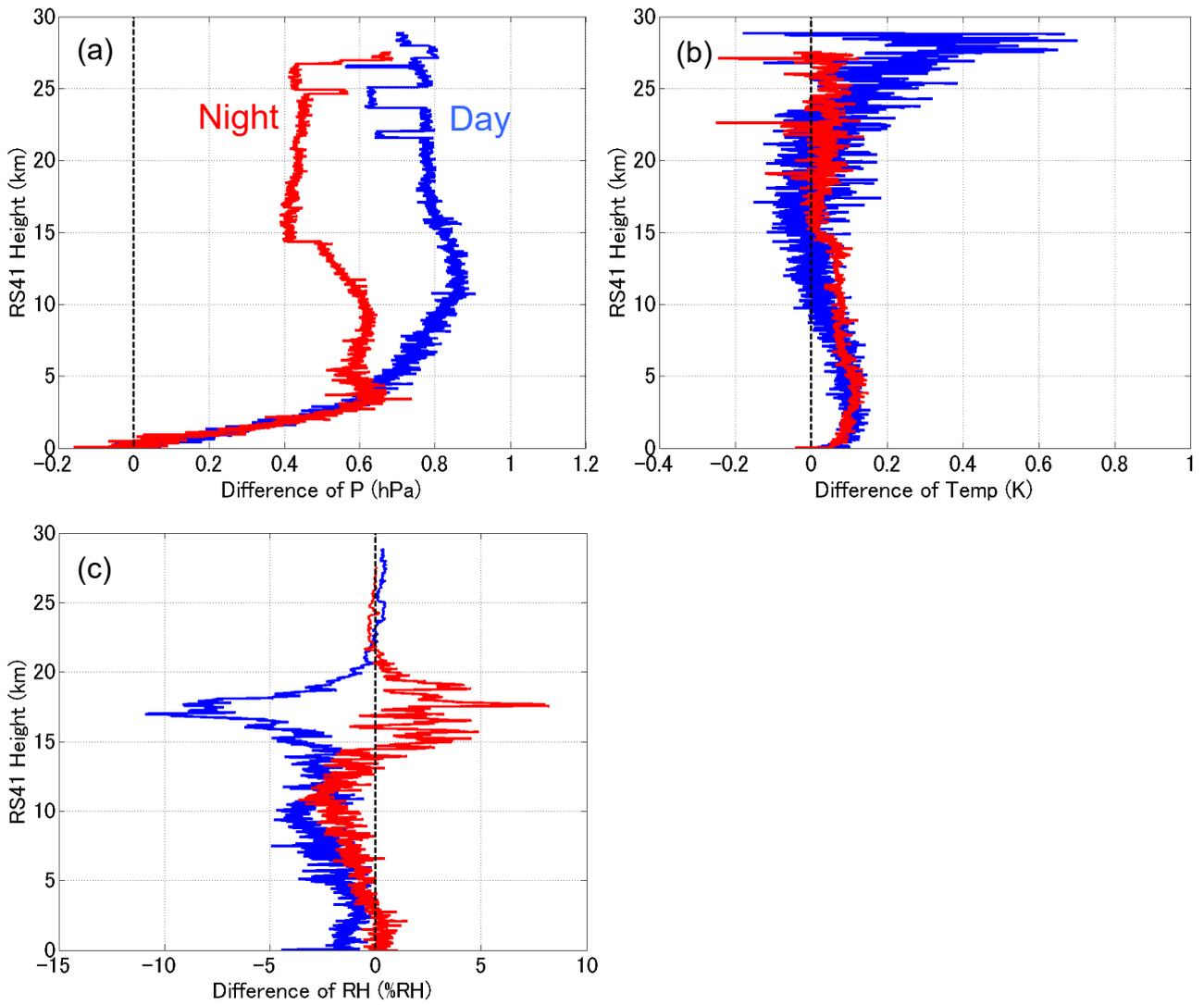


Fig.8. Differences between the RS92 and RS41 radiosonde (RS92 – RS41) results for daytime (blue) and nighttime (red) flights during the MR15-04 cruise for (a) pressure, (b) temperature, and (c) relative humidity.

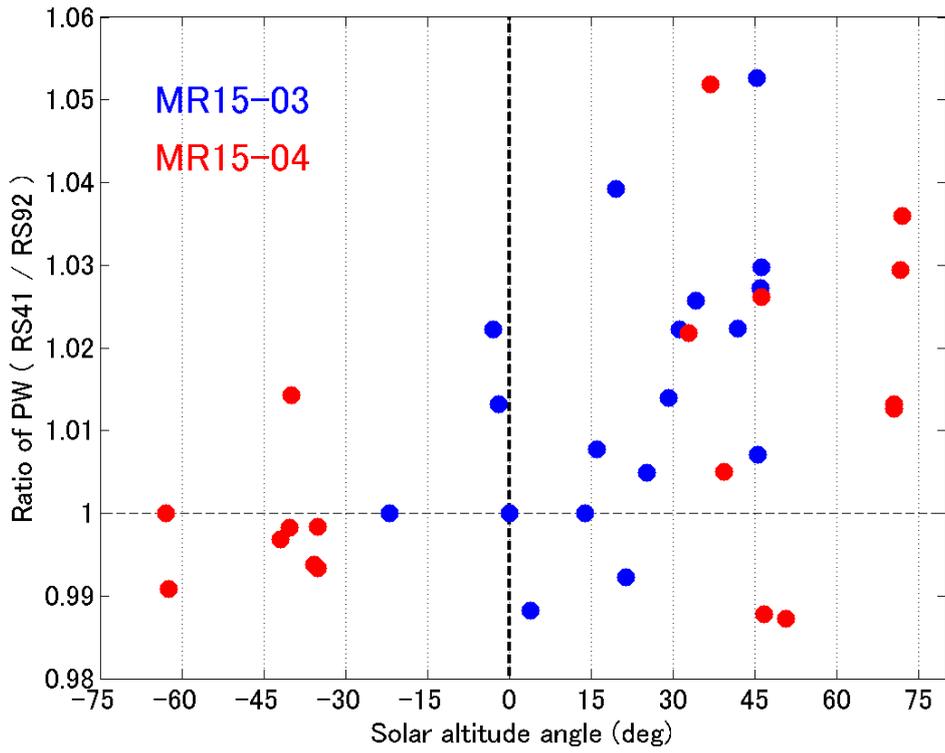


Fig. 9. The ratio of the RS41 to the RS92 PW as a function of solar altitude angle. Blue and red dots represent soundings in the MR15-03 and MR15-04 cruises, respectively.

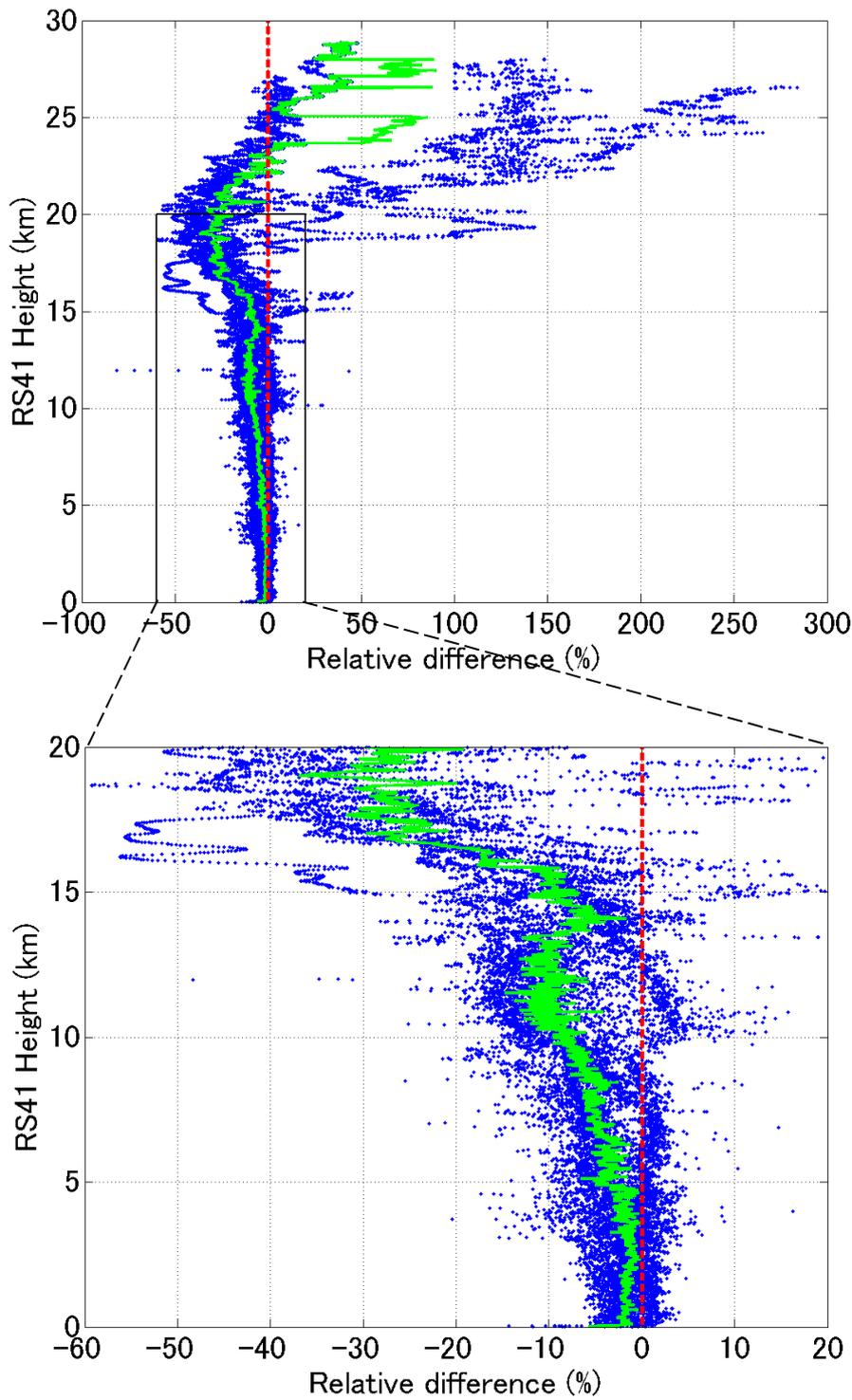


Fig. 910. Relative difference between the RS92 and RS41 relative humidity obtained during the daytime on the MR15-04 cruise (blue dots, %). Relative difference is defined as the relative humidity difference expressed as a percentage of the RS41 relative humidity. Green line denotes the median of the relative difference. Lower panel shows an enlargement of part of the upper panel.

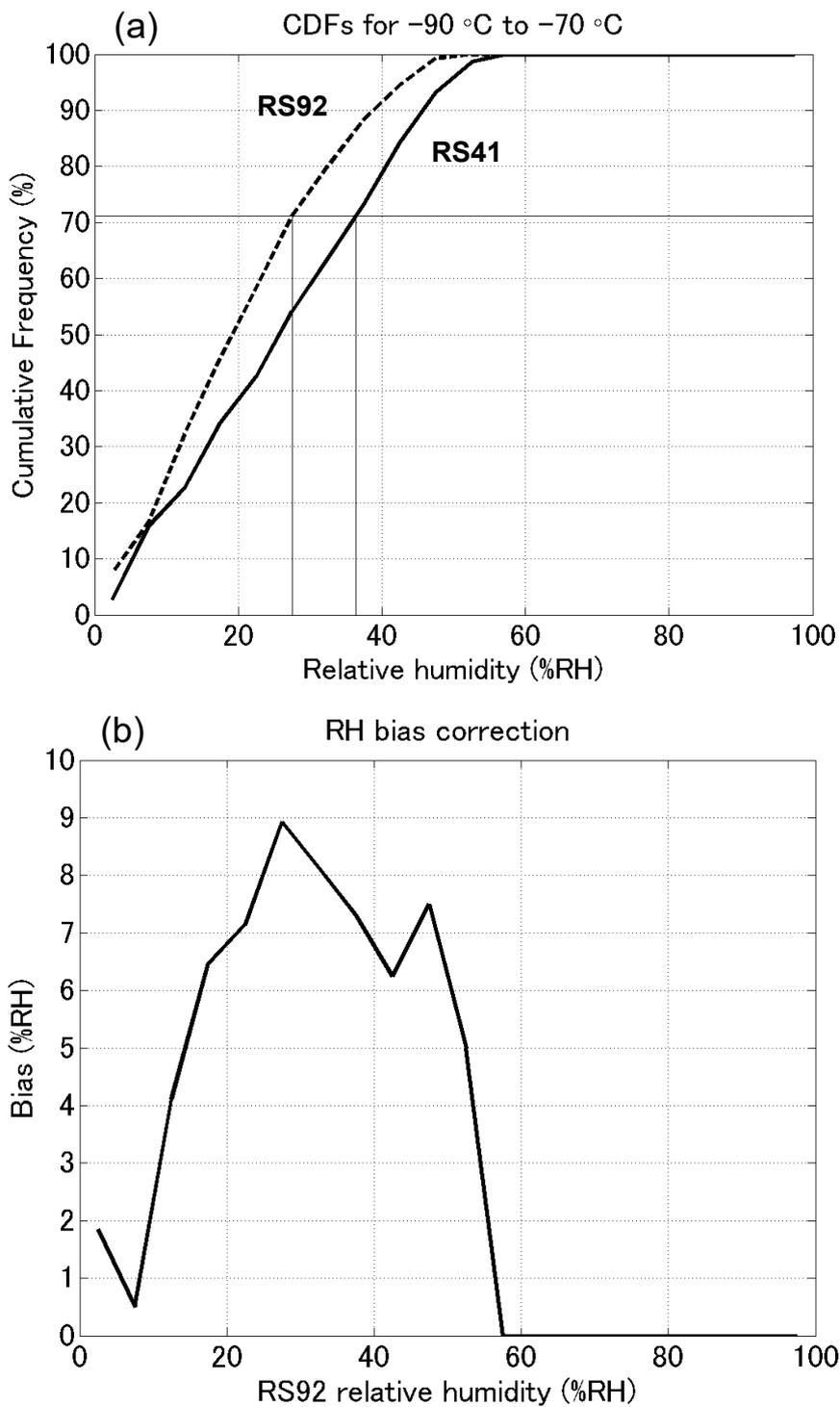


Fig. 1011. (a) CDFs of relative humidity for the RS92 (bold dashed line) and RS41 (bold solid line) data in the temperature range of  $-90$  to  $-70^{\circ}\text{C}$ . The daytime data obtained during the MR15-04 cruise were used. Thin solid lines illustrate the CDF-matching technique (see text). (b) Bias correction of relative humidity for the same temperature range.

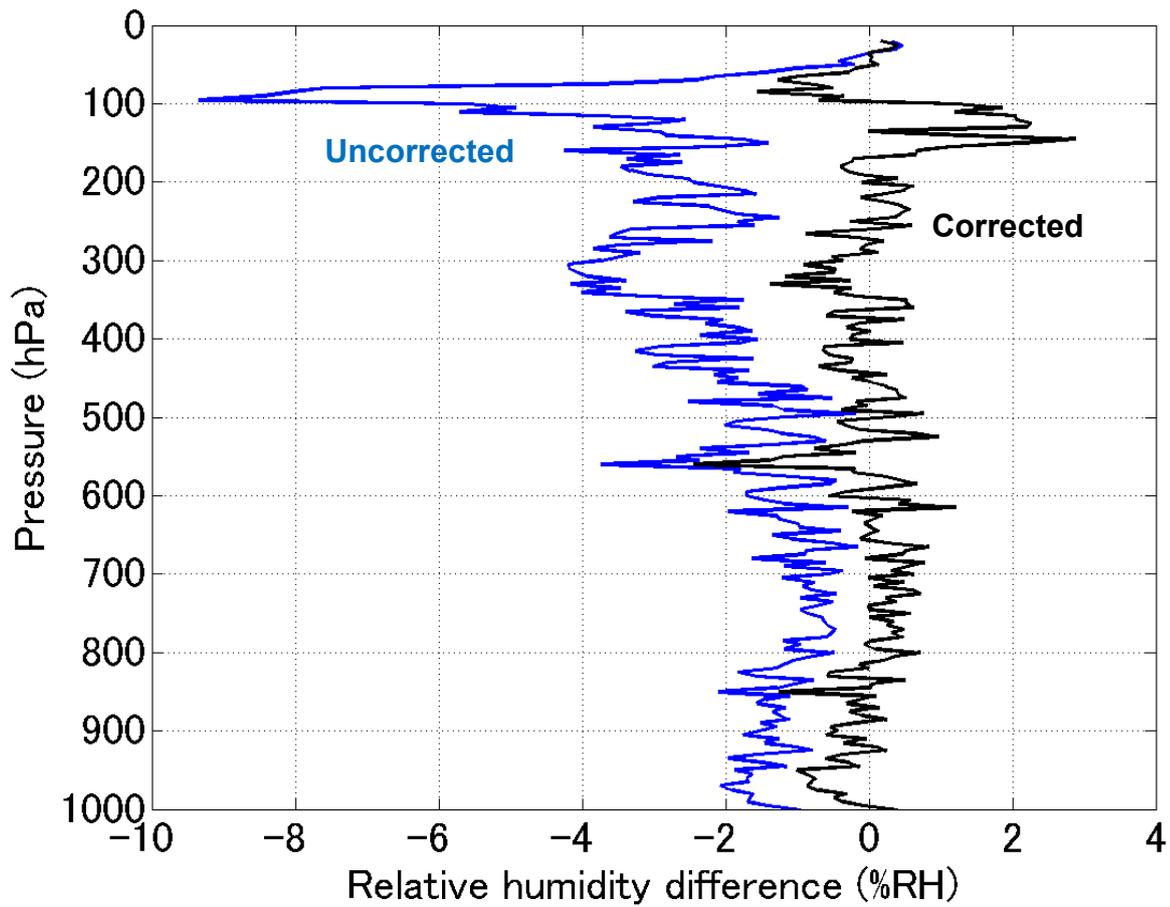


Fig. 4112. Medians of the relative humidity difference between the RS92 and RS41 radiosondes obtained during the daytime on the MR15-04 cruise. Blue and black lines show the profiles before and after the bias correction of the RS92 data.