

1 **A database of paleoceanographic sediment cores from the North Pacific, 1951-2016**

2 Marisa Borreggine¹, Sarah E. Myhre^{1,2*}, K. A. S. Mislan^{1, 3}, Curtis Deutsch¹, Catherine V. Davis⁴

3

4 ¹ School of Oceanography, University of Washington, 1503 NE Boat Street, Box 357940, Seattle, WA

5 98195-7940

6 ² Future of Ice Initiative, University of Washington, Johnson Hall, Room 377A, Box 351360, Seattle, WA

7 98195-1360

8 ³eScience Institute, University of Washington, 3910 15th Ave NE, Box 351570, Seattle, WA 98195

9 ⁴School of Earth, Oceans, and the Environment, University of South Carolina, 701 Sumter Street, Earth &

10 Water Science Building, Room 617, Columbia, SC 29208

11 *Corresponding author

12

13

14

15

16

17

18

19

20

21

22

23

24

25

26

27 **Abstract**

28 We assessed sediment coring, data acquisition, and publications from the North Pacific (north of 30°N)
29 from 1951-2016. There are 2134 sediment cores collected by American, French, Japanese, Russian, and
30 international research vessels across the North Pacific (including the Pacific Subarctic Gyre, Alaskan
31 Gyre, Japan Margin, and California Margin, 1391 cores), Sea of Okhotsk (271 cores), Bering Sea (123
32 cores), and Sea of Japan (349 cores) reported here. All existing metadata associated with these sediment
33 cores are documented here, including coring date, location, core number, cruise number, water depth,
34 vessel metadata, and coring technology. North Pacific sediment core age models are built with isotope
35 stratigraphy, radiocarbon dating, magnetostratigraphy, biostratigraphy, tephrochronology, % opal, color,
36 and lithological proxies. Here, we evaluate the iterative generation of each published age model and
37 provide comprehensive documentation of the dating techniques used, along with sedimentation rates and
38 age ranges. We categorized cores according to the availability of a variety of proxy evidence, including
39 biological (e.g. benthic and planktonic foraminifera assemblages), geochemical (e.g. major trace element
40 concentrations), isotopic (e.g. bulk sediment nitrogen, oxygen, and carbon isotopes), and stratigraphic
41 (e.g. preserved laminations) proxies. This database is a unique resource to the paleoceanographic and
42 paleoclimate communities, and provides cohesive accessibility to sedimentary sequences, age model
43 development, and proxies. The data set is publicly available through PANGAEA at
44 [doi:10.1594/PANGAEA.875998](https://doi.org/10.1594/PANGAEA.875998).

45

46

47

48

49

50

51

52

53 **1 Introduction**

54 Paleoceanographic sediments provide the sedimentary, geochemical, and biological evidence of past
55 Earth system changes. Sediment cores produce robust reconstructions of large oceanographic provinces,
56 and provide insight into earth system mechanistic hypotheses. However, there is not a common repository
57 for paleoceanographic data and publications, and this lack of centralization limits the efficacy of the earth
58 science community to direct research efforts. For the North Pacific, such ongoing mechanistic hypotheses
59 include deep ocean circulation (e.g. Rae et al., 2014; De Pol-Holz et al., 2006), deep-water and
60 intermediate water formation and ventilation (e.g. Knudson and Ravelo, 2015a; Zheng et al., 2000; Cook
61 et al., 2016), and changes in the oceanic preformed nutrient inventories (e.g. Jaccard and Galbraith, 2013;
62 Knudson and Ravelo, 2015b), as well as more regional mechanisms such as sea ice extent (e.g. Max et al.,
63 2012), upwelling intensity (e.g. Di Lorenzo et al., 2008; Hendy et al., 2004), local surface ocean
64 productivity (e.g. Serno et al., 2014; Venti et al., 2017), and terrigenous and marine fluxes of iron (e.g.
65 Davies et al., 2011; Praetorius et al., 2015).

66

67 Paleoceanographers have benefited from the use of large databases of climate data in the past, such as
68 CLIMAP (Climate: Long Range Investigation, Mapping, Prediction), which produced globally resolved
69 temperature records for the Last Glacial Maximum (LGM) and aimed to determine climate system
70 sensitivity from paleoclimate reconstructions (Hoffert & Covey 1992). The PaleoDeepDive project
71 employed a similar approach to the systematic extraction of archival data and constructed a new way to
72 assess and engage with paleobiological data (Zhang, 2015). These projects are examples of international
73 research teams approaching the same limit—extraction and organization of dark data—that arises when
74 creating comprehensive paleo-reconstructions.

75

76 **1.1 Assembling a paleoceanographic database**

77 A clear need exists for high quality paleo-environmental reconstructions to fit the North Pacific into a
78 climate global framework, because the role this enormous ocean basin plays in earth system changes

79 remains relatively unclear in comparison to the Southern and Atlantic Oceans. To address the collective
80 need, we present here a new database of North Pacific paleoceanographic research efforts, along with the
81 broad findings of our census of coring metadata, age model development, and proxy publications. We
82 address the following questions in this manuscript to supplement and provide context for our database:

- 83 1. Where have sediment cores been extracted from the North Pacific seafloor (North of 30°N,
84 including the Pacific Subarctic Gyre, Alaskan Gyre, Japan Margin, and California Margin),
85 the Bering Sea, the Sea of Okhotsk, and the Sea of Japan? What metadata is available in
86 published cruise reports or peer-reviewed investigations, including core name, recovery date,
87 recovery vessel and scientific agency, latitude and longitude, water depth, core length, and
88 coring technology?
- 89 2. For sediment cores with published age models, what lines of evidence were used to develop
90 the chronological age of the sediment, what is the age range from core top to core bottom,
91 and what are the sedimentation rates?
- 92 3. What lines of sedimentary, geochemical, isotopic, and biological proxy evidence are
93 available for each sediment core?
- 94 4. How has the state of North Pacific research efforts and reporting changed since the beginning
95 of paleoceanographic expeditions?

96

97 **1.2 Paleoceanographic age models, proxies, and nomenclature**

98 Marine sedimentary age models tie the sedimentary depth (in meters below sediment surface) to calendar
99 age (ka, thousands of years and/or Ma, millions of years). Not all sedimentary chronologies are of the
100 same quality, and often age models are iteratively refined. Age models are developed with many different
101 dating techniques, which are dependent upon the quality, preservation, and age of the sediments, as well
102 as the investigative priorities of research teams and the proximity of other well-developed sedimentary
103 chronologies. Paleoceanographic proxies, including biological, isotopic, geochemical, and sedimentary
104 observations and measurements, address large thematic questions in the reconstruction of ocean

105 environments, including ocean temperature, paleobiology, seafloor geochemistry, sea ice distribution, and
106 additional sedimentary analyses.

107
108 Sediment cores are often represented by their cruise-core unique identifier, which has the general format
109 "cruise name-core number". The cruise number is generally indicative of the research vessel employed
110 and the year of the expedition. For example, L13-81 is the 13th cruise of S.P. Lee in 1981, MR06-04 is the
111 4th cruise of R/V Mirai in 2006, and YK07-12 is the 12th cruise of R/V Yokosuka in 2007. Often the core
112 number will be preceded or followed by a PC (piston core), MC (multicore), TC (trigger core), or GC
113 (gravity core) to signify the coring technology. The sediment core B37-04G is the 4th gravity core from
114 the 37th cruise of R/V Professor Bogorov, and EW9504-11PC is the 11th piston core from the fourth cruise
115 of R/V Maurice Ewing in 1995. However, this nomenclature is not comprehensive. For example, cores
116 affiliated with iterations of the International Ocean Discovery Program are represented by the program
117 abbreviation and their hole number (i.e. ODP Hole 1209A). Cruise name or number is unknown for many
118 cores, and in these cases the core is referred to by number.

119

120 **2 Methods**

121 Here we assembled a database from peer-reviewed publications, publicly available online cruise reports,
122 and print-only cruise reports available through the University of Washington library network. Detailed
123 metadata was reported for cores where it is available, commonly from cruise reports, including water
124 depth (in meters), recovery year, latitude and longitude, coring technology, and core length (in meters).
125 Summary details regarding affiliated research vessels and institutions were gathered from publications or
126 cruise reports. Cruise reports were commonly available for research expeditions affiliated with
127 JAMSTEC, GEOMAR, and Scripps Institution of Oceanography, and less commonly available for older
128 cores. All evidence used in age model development was reported, along with sedimentation rates and the
129 sedimentary age ranges, to provide investigators with the capacity to quickly evaluate specific cores that
130 meet the investigative priorities. For each core, paleoceanographic proxies and associated publications are

131 documented, to provide an efficient resource for assessing the availability and quality of different lines of
132 paleoenvironmental information. In addition, we evaluated the annual number of age models published
133 (using any dating technique), age models published specifically with radiocarbon dating, publications
134 generated, sediment cores collected, research cruises completed, and the mean number of proxies
135 generated per core.

136

137 **3.0 Results**

138 **3.1 Sediment coring and metadata**

139 We documented 2134 sediment cores and 283 marine geology research cruises above 30°N, from 1951 to
140 2016, in the North Pacific, the Bering Sea, the Sea of Japan, and the Sea of Okhotsk (Figure 1, Table 1).

141 The majority of sediment cores were extracted from the Northern Subarctic Pacific (1391 cores), followed
142 by the Sea of Japan (349 cores), the Sea of Okhotsk (271 cores), and the Bering Sea (123 cores). Many of
143 the oldest cores in this oceanographic province came from the central abyssal Pacific and were associated
144 with the Deep Sea Drilling Project. Cores were extracted from the North Pacific from 1951- 2010, and the
145 oldest age models extend to 120,000 ka (Figure 1, Table 1). Metadata associated with sediment cores or
146 marine research cruises are frequently unavailable or omitted from publications affiliated with sediment
147 cores. For example, 495 cores are in the literature without recovery year, 354 sediment cores were
148 published without latitude and longitude, and 642 cores were reported without specifying the coring
149 technology used (Table 1, 2). Moreover, 1210 sediment cores reported in our database were identified in
150 supplemental tables within publications, however no original cruise reports or peer-reviewed publications
151 otherwise report on these cores.

152

153 **3.2 Sediment chronologies**

154 In the North Pacific, 519 marine sediment cores have published age models, and 266 of these
155 chronologies are generated with radiocarbon dating (^{14}C) of planktonic foraminifera, molluscs, or
156 terrigenous material like bark or wood fragments (Figure 2). Radiocarbon dating is the most common

157 chronological technique region-wide (51% of age models incorporate this method). Lead dating (^{210}Pb) is
158 used for 12 sediment chronologies. Many other lines of evidence are used in the North Pacific and
159 marginal seas to develop paleoceanographic age models. These approaches vary regionally and include
160 planktonic foraminifera oxygen isotope stratigraphy, diatom silica oxygen isotope stratigraphy,
161 biostratigraphy, magnetostratigraphy, tephrochronology, chronostratigraphy, carbonate stratigraphy, opal
162 stratigraphy, composition, lithological proxies, the presence of lamination, chlorin content, and color (a^* ,
163 b^* , and L^* values) (Figure 2, Table 3). For example, in the Sea of Japan, lithological proxies such as core
164 laminations are often utilized as chronological proxy evidence, and 12% of local age models incorporate
165 this technique. Tephrochronology is also applied in 51% of Sea of Japan age models due to regional
166 volcanism. In the Bering Sea, peaks in silica are often used, and 13% of the regional age models
167 incorporated this technique. Published sedimentation rates ranged across the North Pacific (0.1-2000
168 cm/ka), Bering Sea (3-250 cm/ka), Sea of Okhotsk (0.7-250 cm/ka), and the Sea of Japan (0.2-74 cm/ka),
169 with the highest rates within the Alaska Current in the North Pacific (up to 2000 cm/ka).

170

171 **3.3 Paleoceanographic proxies from marine sediment cores**

172 From all reported sediment cores in the North Pacific and marginal seas, only 40% of cores have
173 published proxy data (Figure 3, 5). Stable isotope stratigraphy was available for 293 cores, including
174 oxygen, carbon, or nitrogen isotopes ($\delta^{18}\text{O}$, $\delta^{13}\text{C}$, $\delta^{15}\text{N}$). We documented planktonic (236 cores) and
175 benthic (178 cores) foraminiferal oxygen isotopes, planktonic (67 cores) and benthic (77) foraminifera
176 carbon isotopes, and 34 cores with bulk sediment nitrogen isotopes. Of note, 98 cores were available with
177 magnetostratigraphy (Table 4, Figure 4).

178

179 We recorded paleothermometry data for 234 cores, including planktonic foraminifera oxygen isotopes,
180 magnesium/calcium ratios from planktonic foraminifera, and alkenones TEX_{86} , and U^k_{37} (Figure 3). We
181 recorded 425 cores with microfossil biostratigraphy, including foraminifera, diatom, radiolarian, ostracod,
182 silicoflagellate, ebridian, acritarch, coccolithophore, or dinoflagellate assemblages (Figure 3, Table 4).

183 Biostratigraphy utilizes known microfossil assemblages and their corresponding ages to assign a geologic
184 age range to core strata containing assemblages. Geochemical analyses are reported for 151 cores,
185 including measurements of, for example, brassicasterol, magnesium, calcium, molybdenum, cadmium,
186 manganese, uranium, chromium, rhenium, chlorin, titanium, iron, zinc, and beryllium (Figure 3, Table 3).
187 We documented 234 cores with sea ice proxy data, including geochemical biomarker IP25, ice-rafted
188 debris (IRD), and diatom communities (Figure 3). The presence and concentration of these proxies are
189 indicative of contemporary sea ice extent and volume. We recorded 521 sediment cores with lithological
190 analyses, including the documentation of core lamina, sediment density, mass accumulation rates,
191 biogenic opal and barium, silicon/aluminum weight ratios, carbon/nitrogen ratios, inorganic and organic
192 calcium carbonate and carbon, inorganic nitrogen, sulfur, and clay mineral composition (Figure 3).

193

194 **3.4 Research cruise and publication rate**

195 We cataloged 565 peer-reviewed publications and cruise reports and evaluated the progress of
196 paleoceanographic research using a suite of annual metrics of cruise and core metadata and publications.
197 (Figure 5). The state of North Pacific paleoceanographic investigations has evolved incrementally in the
198 65-year history of research in the region, and we characterize the history into two distinct phases (before
199 and after the early 1990s). Cruise reports were not publicly available for every cruise or core, and many
200 cores cited in cruise reports were never published upon in peer-reviewed literature. The majority of cores
201 (1210 cores, or 57% of all cores extracted from the North Pacific) lack any publication.

202

203 Core recovery rates were high and publication rates were low from 1951-1988, which is a period when
204 expeditions were driven by individual institutions, wherein peer-reviewed publication was not the primary
205 research outcome (Figure 5). Annual rates of cruise completion and sediment core extraction peaked in
206 1965-1968, and this is a consequence of the temporal overlap in peak research efforts by Scripps Institute
207 of Oceanography (1951-1988), Lamont-Doherty Earth Observatory (1964-1975), Oregon State University
208 (1962-1972), and the Deep Sea Drilling Project (1971-1982). Annual rates of publication (peer-reviewed

209 and cruise reports), including those publications with age models, increased around 1995 (Figure 5). In
210 this later period, research cruise efforts were dominated by international research team efforts and
211 resulted in increased peer-reviewed publications, sediment core chronology constructions, and the
212 proliferation of radiocarbon dating. There are 41 cores with very focused investigation (>6 publications),
213 and these archives are primarily located within the California Current (Figure 6). Major peaks in cruising
214 and coring efforts coincided with research cruises by the International Ocean Discovery Program, such as
215 ODP Legs 145 and 146 in 1992 (North Pacific Transect, Cascadia Margin), ODP Leg 167 in 1996
216 (California Margin), and IODP Expedition 323 in 2009 (Bering Sea). Despite the increase in publications
217 around 1995, we observe no distinct temporal trend in the rate of cruise completion and coring effort
218 (Figure 5).

219

220 **4 Discussion**

221 **4.1 Evolving state of North Pacific coring and paleoceanographic approaches**

222 Extensive cruise and research efforts have focused on the marine geology of the North Pacific. Often, the
223 cruise and core metadata from these efforts are unpublished, though they are integral to collaboration,
224 continued research, and publication. Here we present a database with 2134 sediment cores, 283 research
225 cruises, and 565 peer-reviewed publications related to paleoceanographic research (Table 1). We
226 cataloged 519 publications with sedimentary age models, and of those age models, 266 utilized
227 radiocarbon dating, 201 utilized planktonic foraminiferal oxygen isotope stratigraphy, and 129 utilized
228 benthic foraminiferal oxygen isotope stratigraphy (Figure 2, 4). Throughout the North Pacific, Bering
229 Sea, Sea of Okhotsk, and Sea of Japan, the techniques for reconstructing sedimentary age models
230 regionally varied. We documented a community-wide shift away from singular dating techniques toward
231 age models which incorporate several techniques. Multiproxy approaches hold merit through verifying or
232 constraining the results of a singular proxy, and thereby disentangling multiple environmental drivers and
233 providing redundancy in order to create robust records of climate and ocean conditions (Mix et al., 2000;
234 Mann, 2002). Age model development has moved through the last 60 years to more detailed high-

235 resolution age models constructed to investigate millennial and submillennial paleoceanographic
236 variability (Figure 5, 6).

237

238 **4.2 The merit of database management and open-access science**

239 Databases are integral to facilitate efficient, hypothesis-based investigations into earth system
240 mechanisms. Public access to databases facilitate a higher volume of research by a diverse range of
241 scientists (Harnad and Brody 2004). The necessity for databases to encompass a wide array of data over
242 large oceanographic provinces is also largely recognized. Open access tools from PANGAEA support
243 database-dependent research, because hypothesis-based investigations can be more efficient through
244 public data access. For example, content from online databases has contributed to research in atmospheric
245 forcing (e.g. Shaffer et al., 2009), Atlantic Meridional Overturning Circulation (AMOC, e.g. Schmittner
246 and Galbraith, 2008), and Southern Ocean ventilation (e.g. Yamamoto et al., 2015). The metadata in these
247 databases must be thorough, as data is impractical without the affiliated identifier, location, and methods
248 (Lehnert et al., 2000). Database management should be a priority in any field that incorporates the
249 contents of online repositories of knowledge and research. The disconnect between the research goals of
250 the paleoceanographic community and the metadata produced here can be described as a “breakdown”, a
251 limit on the progress of paleoceanographic research (Tanweer et al., 2016). These breakdowns allow us to
252 self-assess and move forward with insight into best practices. Metadata is produced from datasets that are
253 inherently human in design, and therefore are not inerrant. Assessment of the errors in metadata reporting
254 can directly reveal the need for community-wide improvements. As an example, all cores should be
255 reported with latitude and longitude; the absence of this specific metadata significantly impairs further
256 work. The database presented here, as well as others like it, consolidates the research effort of an entire
257 community into an efficient tool for future investigative purposes.

258

259 **4.3 Recommendations for the marine geology community**

260 Ocean sediment records are one of the primary tools for understanding earth's history, and the
261 documentation of these records benefits the entire community of earth scientists. The publications and
262 cruise reports here represent a large body of research completed on North Pacific sediment records,
263 however this may not constitute the entire body of work. There was no preexisting common repository for
264 cruise reports and coring data in the North Pacific, beyond the individual institutional archival processes.
265 This database serves as the most complete archive of the public availability of cruise reports and
266 publications; where available doi and urls are reported in the database. We demonstrate a need for more
267 thorough, accessible documentation of marine geology and paleoceanographic research. In our
268 examination of publications, cruise reports, and notes from research cruises, we gained insights into past
269 inconsistencies in marine sediment record reporting. We recommend a suite actions to ensure efficient,
270 comprehensive sediment core collection, metadata documentation, and the publication of chronologies
271 and proxies. We propose that each publication thoroughly reports metadata on all cores discussed, and
272 their associated cruises, including core unique identifier numbers, cruise name and number, vessel name,
273 geographic coordinates, core recovery date, core length, core recovery water depth (meters below sea
274 level), sampling resolution, sampling volume, core archival repository, and the link (if existing) to public
275 cruise reports. We also suggest summarizing each core's metadata, age model, and publication history of
276 previous publications in the methods section in order to provide a frame of reference for new findings,
277 especially in the context of iterative age model revisions.

278

279 **5 Author Contribution**

280 S. E. Myhre and C. V. Davis initiated the building of this database. M. Borreggine and S. E. Myhre built
281 the database and were joined by K. A. S. Mislán in producing figures and analysis for this manuscript. All
282 authors wrote the manuscript.

283

284 **6 Acknowledgements**

285 The authors would like to thank the University of Washington Library Oceanography collection, namely
286 Louise Richards and Maureen Nolan. We also wish to acknowledge the support for this publication,
287 provided by the University of Washington Purple and Gold Scholarship, the UW School of Oceanography
288 Lowell K. and Alice M. Barger Endowed Scholarship, the Clarence H. Campbell Endowed Lauren
289 Donaldson Scholarship, the UW College of the Environment Student Travel Grant, and NSF Grant OCE-
290 1458967. K.A.S. was supported by the Washington Research Foundation Fund for Innovation in Data-
291 Intensive Discovery and the Moore/Sloan Data Science Environments Project at the University of
292 Washington. The database can be found online at (doi:10.1594/PANGAEA.875998) from PANGAEA.

293

294

295

296

297

298

299

300

301

302

303

304

305

306

307

308

309

310

311 **Figure and Tables**

312 **Table 1.** Regional summary of sediment core database for the Bering Sea, North Pacific, Sea of Japan,
 313 and Sea of Okhotsk, including number of cores recovered, the regional percent (%) of cores reported with
 314 latitude and longitude, number of research cruises, total regional publication count, count of cores with no
 315 peer-reviewed publications or cruise report, the range of core recovery years, the regional percent (%) of
 316 cores reported with recovery years, and the range of recovered core water depths (meters below sea level).

Region	Bering Sea	North Pacific	Sea of Japan	Sea of Okhotsk	Total
Count of cruises	24	179	78	31	283
Count of publications	70	306	114	90	565
Count of cores with no publications* or cruise report	61	849	172	128	1210
Number of cores recovered	123	1391	349	271	2134
Percent of regional cores (%) reported with latitude and longitude	91.9	81.9	77.7	91.9	83.4
Range of core recovery years	1957-2009	1951-2009	1957-2010	1972-2009	1951-2010
Percent of regional cores (%) reported with recovery year	87	68.9	96.9	98.5	76.8
Range of recovered core water depths (meters below sea level)	33-3930	21-9585	129-5986	105-8180	21-9585

317 * These cores are listed in large data tables in auxiliary publications, but the original reporting is not publicly
 318 available.

319
 320
 321
 322
 323
 324

325 **Table 2.** Percent of regional cores reported with coring technology, and the number of cores recovered in
 326 the North Pacific and marginal seas by coring technology. Additional reported coring technologies
 327 include the less often utilized Hydrostatic cores, Kasten cores, Asura cores, Pressure cores, and Trigger
 328 cores.

Region	Bering Sea	North Pacific	Sea of Japan	Sea of Okhotsk
% of regional cores reported with drilling technology	75.6	72.1	60.2	68.6
# of Piston cores	33	476	177	67
# of Gravity cores	11	250	15	73
# of Box cores	2	78	1	1
# of Riserless Drilling cores	7	52	0	0
# of Multicores	33	39	1	39
# of Phleger cores	0	15	0	0
# of Jumbo Piston cores	4	10	0	0

329
 330
 331
 332
 333
 334
 335
 336
 337
 338
 339
 340
 341

342 **Table 3.** Summary statistics on core chronology, including the number of cores with radiocarbon dating
 343 (^{14}C) and oxygen isotope stratigraphy of planktonic foraminifera ($\delta^{18}\text{O}$), as well as the regional mean core
 344 top and bottom ages, the number of cores with published sedimentation rate ranges and means, and core
 345 length minimums, maximums, and averages.

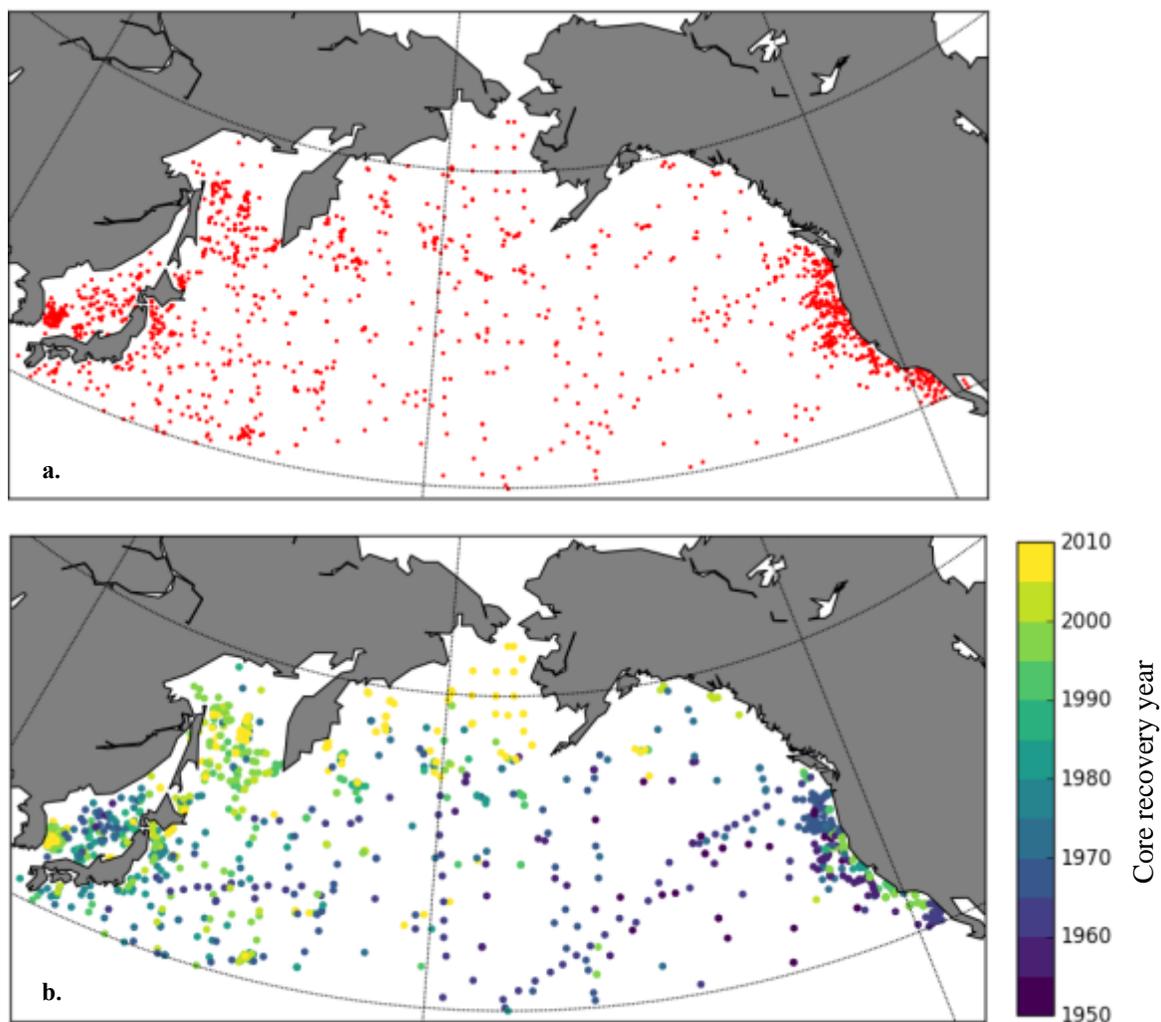
Region	Bering Sea	North Pacific	Sea of Japan	Sea of Okhotsk
^{14}C Dating	28	158	38	42
$\delta^{18}\text{O}$ Stratigraphy, Planktonic foraminifera	20	122	29	32
Mean core top age (Calendar age, ka)	6.0 \pm 9.6	9.5 \pm 61.6	5.9 \pm 28.4	1.1 \pm 2.3
Mean core bottom age (Calendar age, ka)	523.7 \pm 1146.2	6210.8 \pm 20250.7	996.1 \pm 2602.2	153.1 \pm 279.2
Cores with sedimentation rate range	28	97	20	18
Cores with sedimentation rate mean	7	68	9	16
Average core length (m)	56.89	35.7	35.98	0.02
Minimum core length (m)	0.06	0.03	0.24	53.88
Maximum core length (m)	745	1180	555.3	6.50

346
 347
 348
 349
 350
 351
 352
 353
 354

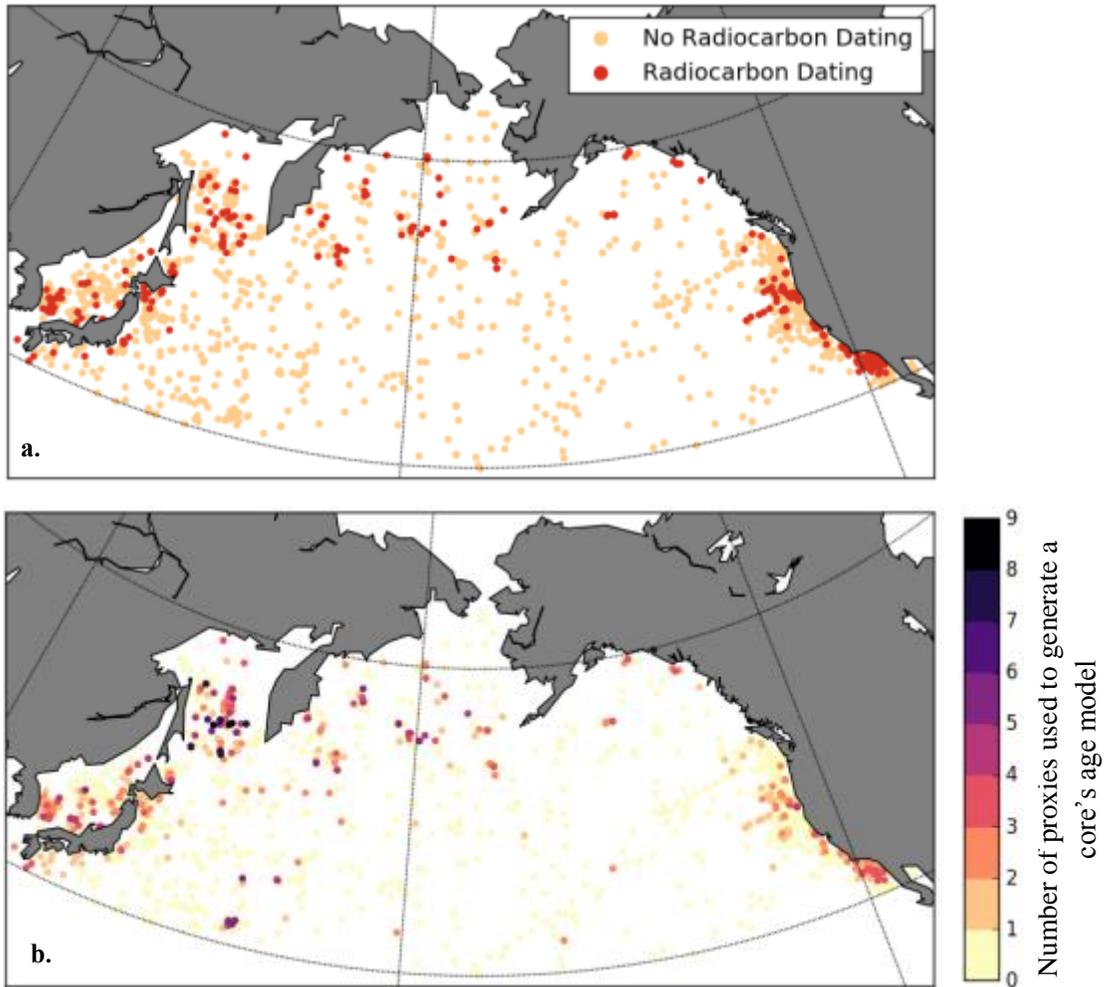
355 **Table 4.** Regional summary of isotopic, geochemical, biological, and sedimentary proxies. Benthic and
 356 planktonic isotopic analysis is for all cores, including, but not limited to, isotope analysis used in core
 357 chronology development.

Region	Bering Sea	North Pacific	Sea of Japan	Sea of Okhotsk
$\delta^{18}\text{O}_{\text{benthic}}$	19	118	15	25
$\delta^{18}\text{O}_{\text{planktonic}}$	20	139	45	32
$\delta^{13}\text{C}_{\text{benthic}}$	7	57	7	6
$\delta^{13}\text{C}_{\text{planktonic}}$	6	50	10	1
$\delta^{15}\text{N}$	5	16	12	1
Alkenone U^{K}_{37}	7	64	15	12
TEX_{86}	0	1	0	0
Foraminiferal Biostratigraphy	12	82	22	8
Foraminiferal Abundance	8	53	41	18
Diatom Biostratigraphy	35	87	49	25
Coccolithophore Biostratigraphy	2	9	4	0
Ostracod Biostratigraphy	1	1	2	0
Geochemical Proxies (Mg, Mo, Cd, Mn, U, Cr, Re, Ca, T, Fe, Mg, Zn, Be, Chlorin)	1	189	46	27
% Opal	21	38	20	20
Total Inorganic Carbon (%)	16	204	49	26
Total Organic Carbon (%)	16	142	62	26

358
 359
 360
 361
 362
 363
 364
 365
 366
 367
 368
 369



370
 371 **Figure 1.** Location and recovery year of marine sediment cores from the North Pacific and marginal seas.
 372 **a.** Sediment cores in the North Pacific (above 30°N) published latitudes and longitudes (354 additional
 373 cores were documented in either cruise reports or peer-reviewed publications without latitude and
 374 longitude). **b.** Sediment cores published with an associated core recovery year, ranging from 1951-2010,
 375 and this age range corresponds with the color bar (495 cores have been published in either cruise reports
 376 or peer-reviewed publications without recovery year).
 377



378

379 **Figure 2.** Age model development for sediment cores in the North Pacific and marginal seas. **a.** Cores

380 with age models that have been constructed with radiocarbon dating (^{14}C). **b.** Number of

381 paleoceanographic proxies used to generate each core's age model.

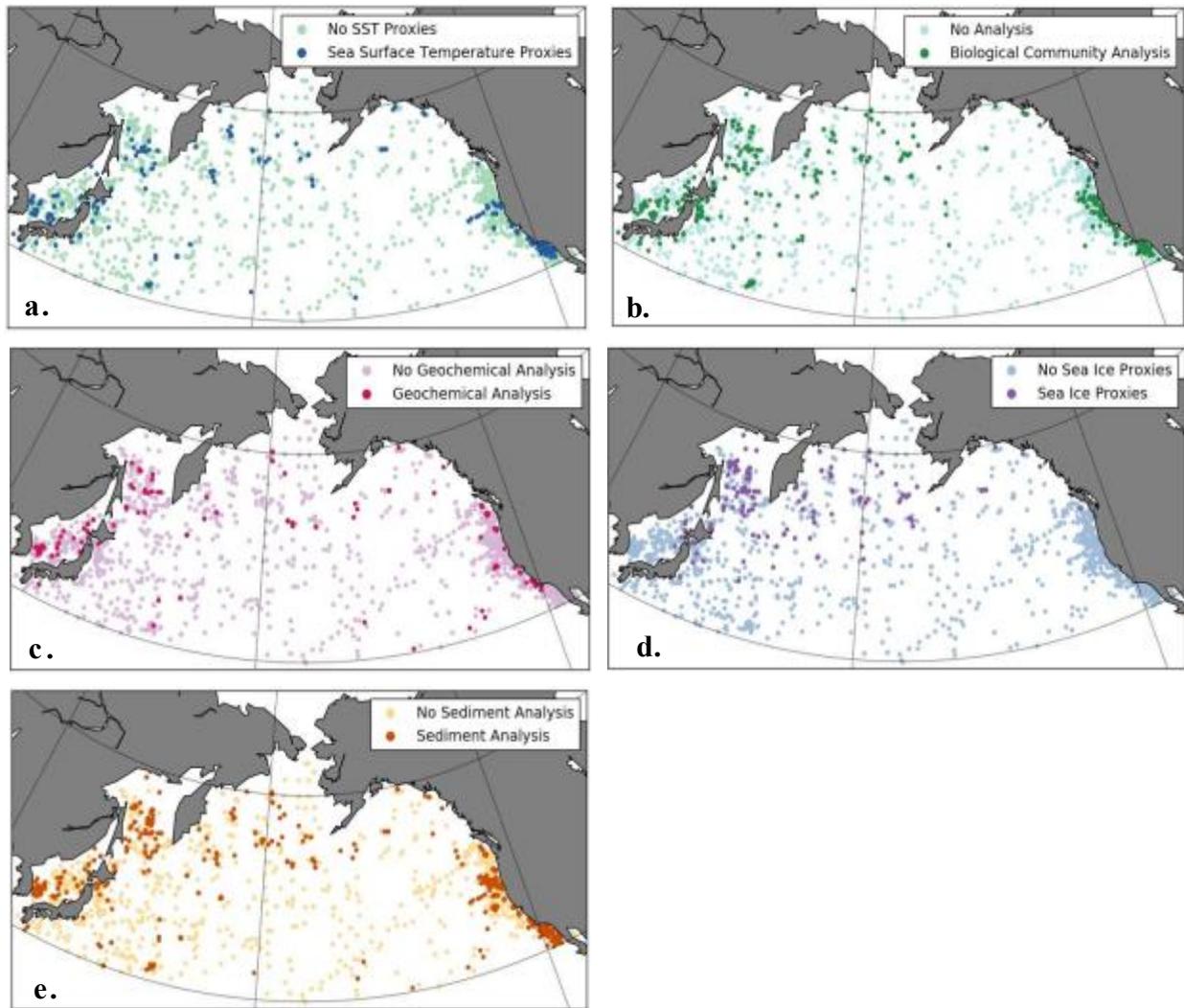
382

383

384

385

386



387

388

Figure 3. Published paleoceanographic proxies in the North Pacific. a. sea surface temperature

389

reconstructions including planktonic foraminifera oxygen isotopes ($\delta^{18}O_p$), magnesium/calcium

390

measurements, TEX_{86} , and U^{k}_{37} alkenones b. biostratigraphy of microfossils, including foraminifera,

391

diatoms, radiolarians, ostracods, silicoflagellates, ebridians, acritarchs, coccolithophores, and

392

dinoflagellates c. geochemical proxy analysis, including trace metals such as brassicasterol, magnesium,

393

calcium, molybdenum, cadmium, manganese, uranium, chromium, rhenium, chlorine, titanium, iron, zinc,

394

and beryllium d. presence of sea ice proxies including geochemical biomarker IP25, ice-rafted debris

395

(IRD), and sea ice related diatom communities e. analysis of lithological core proxies, including

396

measurements of core lamina, biogenic opal and barium, silicon/aluminum weight ratio, sulfur, inorganic

397 and organic carbon weight and mass accumulation rates, mass accumulation rates of various elements,
398 inorganic nitrogen, carbon to nitrogen ratios, sediment density, and clay mineral composition

399

400

401

402

403

404

405

406

407

408

409

410

411

412

413

414

415

416

417

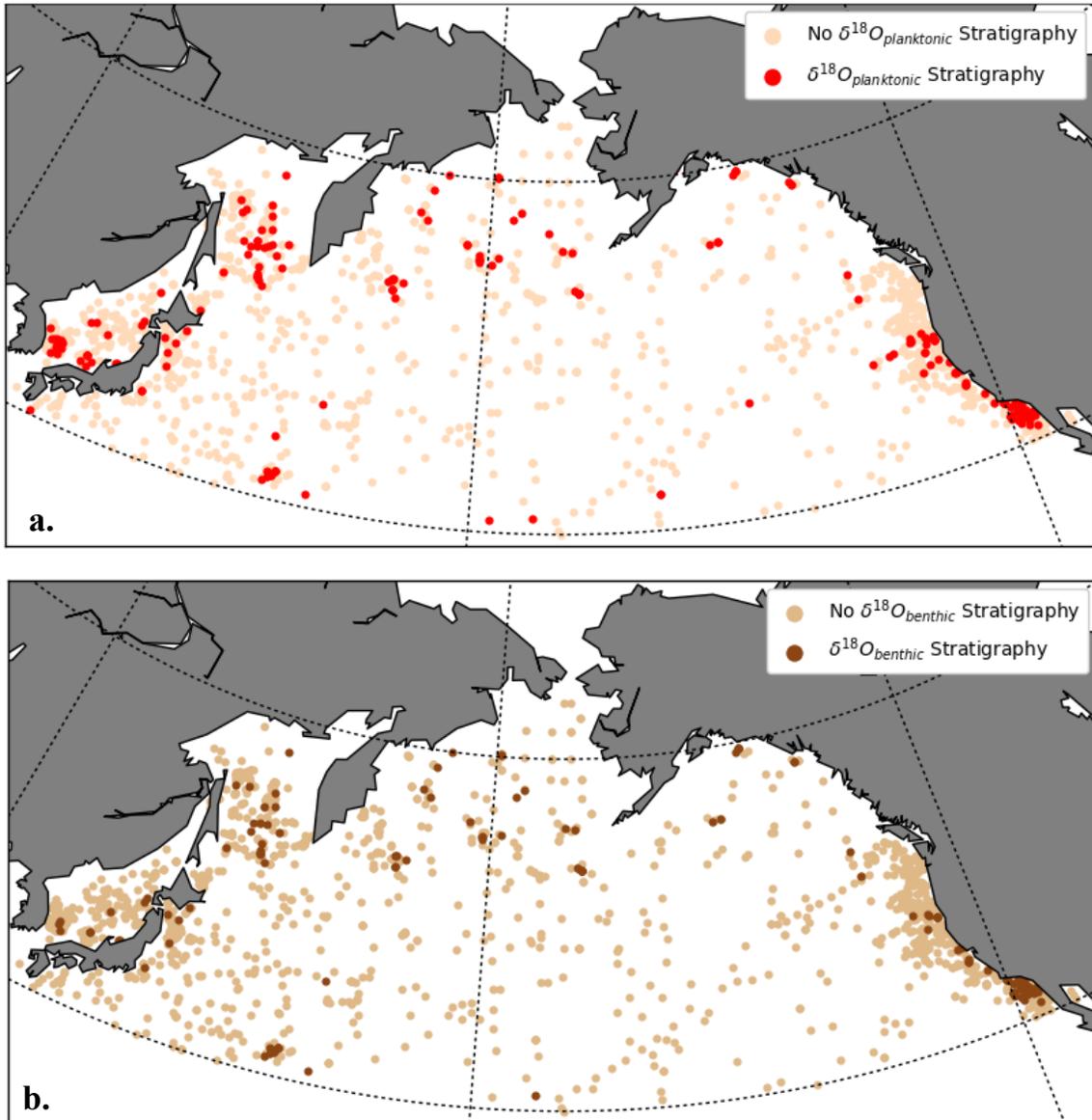
418

419

420

421

422



424

425 **Figure 4.** Published planktonic and benthic foraminiferal oxygen isotope stratigraphy for sediment core

426 age models. a. analysis of planktonic foraminiferal oxygen isotopes b. analysis of benthic foraminiferal

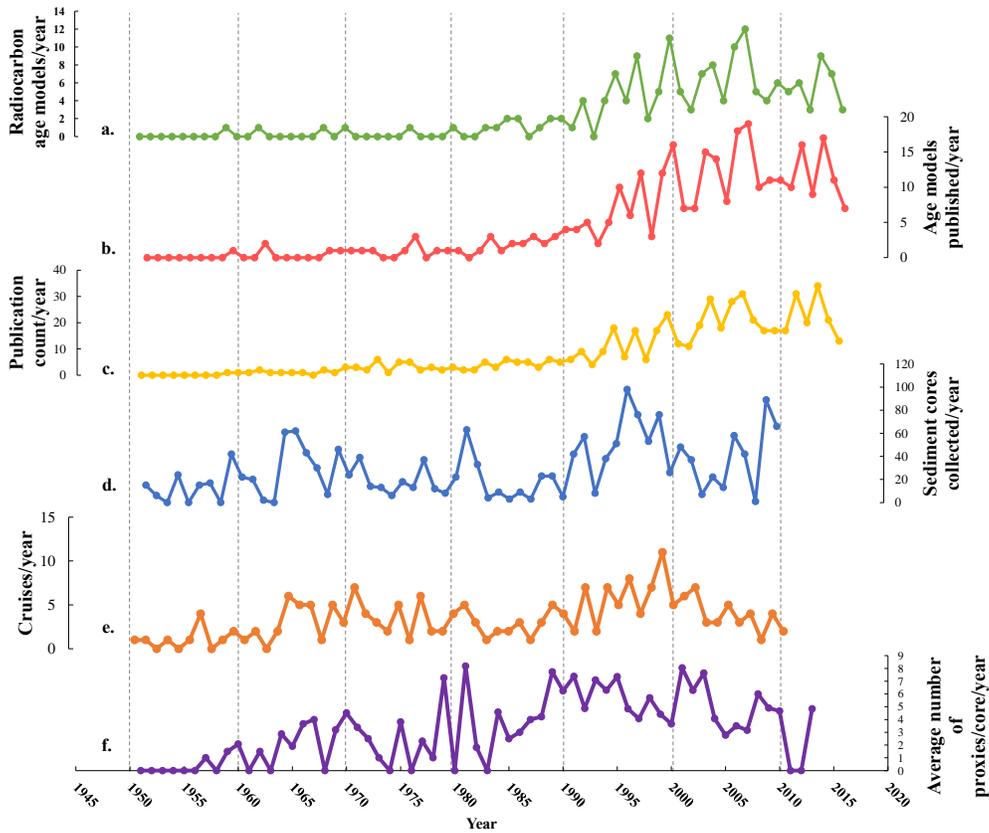
427 oxygen isotopes

428

429

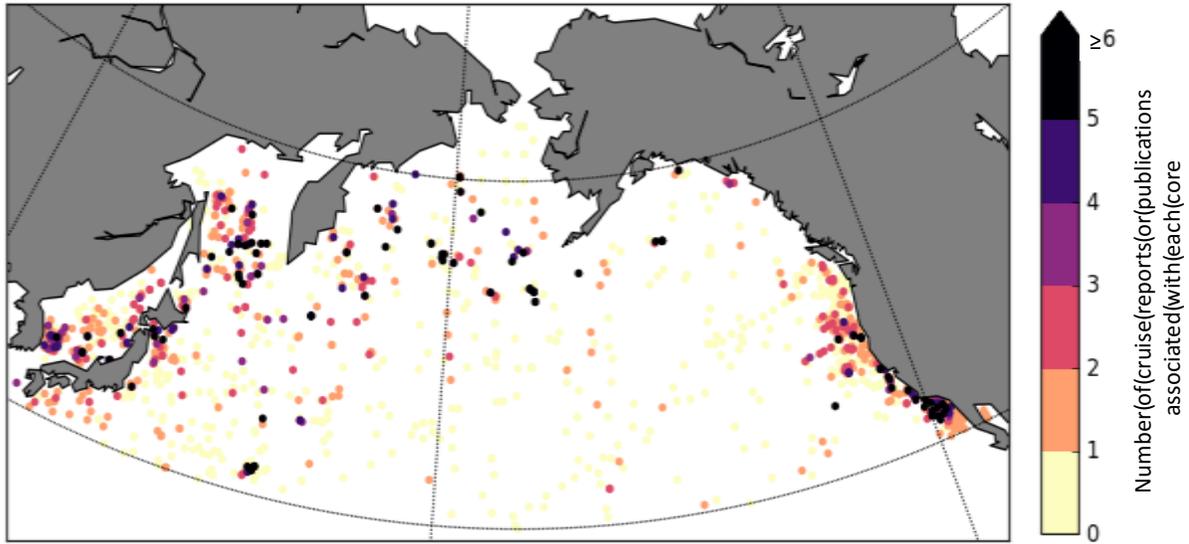
430

431
432
433
434



435

436 **Figure 5.** North Pacific and marginal seas marine geology cruise and paleoceanographic research
437 progress through time, wherein the lead-lag timing of cruise reporting and core publication is assumed for
438 the most recent years. We utilize peer-reviewed publications to locate cores, and there is a lag between
439 publication and core extraction. **a.** Annual number of age models published using radiocarbon dating. **b.**
440 Annual number of all age models published. **c.** Total annual number of publications and cruise reports. **d.**
441 Annual number of cores collected. **e.** Marine research expedition count per year. **f.** Annual mean number
442 of proxy analyses published per core.



443

444 **Figure 6.** Number of affiliated publications and cruise reports for each core. Maximum publication count
445 for an individual core is 23.

446

447

448

449

450

451

452

453

454

455

456

457

458

459 7 **References**

460

461 Cook, M. S., Ravelo, A. C., Mix, A., Nesbitt, I. M., and Miller, N. V. 2016. Tracing subarctic Pacific
462 water masses with benthic foraminiferal stable isotopes during the LGM and late Pleistocene. *Deep*
463 *Sea Research Part II: Topical Studies in Oceanography*. 125: 84-95. doi:10.1016/j.dsr2.2016.02.006.

464 CLIMAP Project Members. 1981. Seasonal reconstruction of the earth's surface at the last glacial
465 maximum. Geological Society of America, Map and Chart Series 36. 18 pp.

466 <https://www.ncdc.noaa.gov/paleo/metadata/noaa-ocean-2516.html>.

467 Davies, M. H., A. C. Mix, J. S. Stoner, J. A. Addison, J. Jaeger, B. Finney, and J. Wiest. 2011. The
468 deglacial transition on the southeastern Alaska Margin: Meltwater input, sea level rise, marine
469 productivity, and sedimentary anoxia. *Paleoceanography* 26. doi:10.1029/2010pa002051.

470 Di Lorenzo, E., N. Schneider, K. M. Cobb, P. J. S. Franks, K. Chha, A. J. Miller, J. C. McWilliams, S. J.
471 Bograd, H. Arango, E. Curchitser, T. M. Powell, and P. Rivière. 2008. North Pacific Gyre Oscillation
472 links ocean climate and ecosystem change. *Geophysical Research Letters* 35.

473 doi:10.1029/2007GL032838.

474 GTN-P. 2015. Global Terrestrial Network for Permafrost metadata for permafrost boreholes (TSP) and
475 active layer monitoring (CALM) sites. doi:10.1594/PANGAEA.842821.

476 Harnad, S. and Brody, T. 2004. Comparing the impact of open access (OA) vs. non-OA articles in the
477 same journals. *D-lib Magazine* 10. Retrieved from: <https://eprints.soton.ac.uk/260207/>.

478 Hartmann, Jörg; Moosdorf, Nils. 2012. Global Lithological Map Database v1.0 (gridded to 0.5° spatial
479 resolution). doi:10.1594/PANGAEA.788537.

480 Hendy, I. L., T. F. Pedersen, J. P. Kennett, and R. Tada. 2004. Intermittent existence of a southern
481 Californian upwelling cell during submillennial climate change of the last 60 kyr.

482 *Paleoceanography* 19. doi:10.1029/2003pa000965.

483 Hoffert, M. I., and C. Covey. 1992. Deriving global climate sensitivity from palaeoclimate
484 reconstructions. *Nature* 360: 573–576. doi:10.1038/360573a0.

485 Jaccard, S. L., and E. D. Galbraith. 2013. Direct ventilation of the North Pacific did not reach the deep
486 ocean during the last deglaciation. *Geophysical Research Letters* 40: 199–203.

487 doi:10.1029/2012gl054118.

488 Knudson, K. P., and A. C. Ravelo. 2015a. North Pacific Intermediate Water circulation enhanced by the
489 closure of the Bering Strait. *Paleoceanography*. 30: 1287-1304. doi:10.1002/2015PA002840.

490 Knudson, K. P., and A. C. Ravelo. 2015b. Enhanced subarctic Pacific stratification and nutrient
491 utilization during glacials over the last 1.2 Myr. *Geophysical Research Letters*. 42: 9870-9879.

492 doi:10.1002/2015GL066317.

493 Lehnert, K., Y. Su, C. H. Langmuir, B. Sarbas, and U. Nohl. 2000. A global geochemical database
494 structure for rocks. *Geochemistry, Geophysics, Geosystems* 1. doi:10.1029/1999gc000026.

495 Le Moigne, Frédéric AC. 2013. Global database of surface ocean particulate organic carbon export fluxes
496 diagnosed from the ²³⁴Th technique. doi:10.1594/PANGAEA.809717.

497 Mann, M. E. 2002. Climate Reconstruction: The Value of Multiple Proxies. *Science* 297: 1481–1482.
498 doi:10.1126/science.1074318.

499 Max, L., J.-R. Riethdorf, R. Tiedemann, M. Smirnova, L. Lembke-Jene, K. Fahl, D. Nürnberg, A. Matul,
500 and G. Mollenhauer. 2012. Sea surface temperature variability and sea-ice extent in the subarctic
501 northwest Pacific during the past 15,000 years. *Paleoceanography* 27. doi:10.1029/2012PA002292.

502 Mix, A. C., E. Bard, G. Eglinton, L. D. Keigwin, A. C. Ravelo, and Y. Rosenthal. 2000. Alkenones and
503 multiproxy strategies in paleoceanographic studies. *Geochemistry, Geophysics, Geosystems* 1.
504 doi:10.1029/2000gc000056.

505 Pol-Holz, R. D., O. Ulloa, L. Dezileau, J. Kaiser, F. Lamy, and D. Hebbeln. 2006. Melting of the
506 Patagonian Ice Sheet and deglacial perturbations of the nitrogen cycle in the eastern South Pacific.
507 *Geophysical Research Letters* 33. doi:10.1029/2005gl024477.

508 Praetorius, S. K., A. C. Mix, M. H. Walczak, M. D. Wolhowe, J. A. Addison, and F. G. Prahl. 2015.
509 North Pacific deglacial hypoxic events linked to abrupt ocean warming. *Nature* 527: 362–366.
510 doi:10.1038/nature15753.

511 Rae, J. W. B., M. Sarnthein, G. L. Foster, A. Ridgwell, P. M. Grootes, and T. Elliott. 2014. Deep water
512 formation in the North Pacific and deglacial CO₂ rise. *Paleoceanography* 29: 645–667.
513 doi:10.1002/2013PA002570.

514 Schmittner, A., and E. D. Galbraith. 2008. Glacial greenhouse-gas fluctuations controlled by ocean
515 circulation changes. *Nature* 456: 373–376. doi:10.1038/nature07531.

516 Serno, S., G. Winckler, R. F. Anderson, C. T. Hayes, H. Ren, R. Gersonde, and G. H. Haug. 2014. Using
517 the natural spatial pattern of marine productivity in the Subarctic North Pacific to evaluate
518 paleoproductivity proxies. *Paleoceanography* 29: 438–453. doi:10.1002/2013PA002594.

519 Shaffer, G., S. M. Olsen, and J. O. P. Pedersen. 2009. Long-term ocean oxygen depletion in response to
520 carbon dioxide emissions from fossil fuels. *Nature Geoscience* 2: 105–109. doi:10.1038/ngeo420.

521 Tanweer, A., B. Fiore-Gartland, and C. Aragon. 2016. Impediment to insight to innovation: understanding
522 data assemblages through the breakdown–repair process. *Information, Communication & Society* 19:
523 736–752. doi:10.1080/1369118x.2016.1153125

524 Venti, N. L., K. Billups, and T. D. Herbert. 2017. Paleoproductivity in the northwestern Pacific Ocean
525 during the Pliocene-Pleistocene climate transition (3.0–1.8 Ma). *Paleoceanography* 32: 92–103.
526 doi:10.1002/2016PA002955.

- 527 Yamamoto, A., A. Abe-Ouchi, M. Shigemitsu, A. Oka, K. Takahashi, R. Ohgaito, and Y. Yamanaka.
528 2015. Global deep ocean oxygenation by enhanced ventilation in the Southern Ocean under long-term
529 global warming. *Global Biogeochemical Cycles* 29: 1801–1815. doi:10.1002/2015gb005181.
- 530 Zhang, C. 2015. DeepDive: a data management system for automatic knowledge base construction.
531 Thesis. Retrieved from: <http://cs.stanford.edu/people/czhang/zhang.thesis.pdf>.
- 532 Zheng, Y., A. V. Geen, R. F. Anderson, J. V. Gardner, and W. E. Dean. 2000. Intensification of the
533 Northeast Pacific oxygen minimum zone during the Bølling-Allerød Warm Period. *Paleoceanography*
534 15: 528–536. doi:10.1029/1999pa000473.