



- 1 **ShellTrace v1.0 - A new approach for modelling growth and trace element uptake in marine**
- 2 **bivalve shells: Model verification on pacific oyster shells (*Crassostrea gigas*)**
- 3 *de Winter, Niels J.*
- 4 Analytical, Environmental and Geochemistry (AMGC) Department, Vrije Universiteit Brussel, Pleinlaan 2, 1050
- 5 Brussels, Belgium
- 6





## 7 Abstract

8 Bivalve shells record changes ~~in their~~ environment in the chemical composition of their shells and are  
9 frequently used as paleoclimate archives. However, many studies have shown that physiological changes  
10 related to growth of the bivalve may overprint these chemical tracers. In the present study, a new approach is  
11 presented that models growth and development of bivalve shells without a priori knowledge of the physiology  
12 of the species. The model uses digitized growth increments in a cross section of the shell together with basic  
13 assumptions of the shape of the shell in order to model changes in shell length, thickness, volume, mass and  
14 growth rate at a daily resolution through the lifetime of the bivalve. This approach reconstructs the growth  
15 history of bivalves based on their shell without the need for a culture experiment, paving the way for growth  
16 rate estimations based on fossil bivalve shells. Combination of the growth model with 2D X-Ray Fluorescence  
17 trace element mapping allows the incorporation of trace elements into the shell to be modelled in 3D through  
18 the bivalve's lifetime. This approach yields records of integrated total-shell trace element concentrations and  
19 accumulation rates, which shed light on the rates and mechanisms by which these trace elements are  
20 incorporated into the shells of bivalves. Application of growth and trace element modelling on a set of modern  
21 pacific oyster shells of well-known origin and comparison of model results with conventional trace element  
22 transects highlights the importance of considering heterogeneity in mineralogy, mineralization rates and  
23 chemical composition within the shells of bivalves. These insights lead to a better understanding of the  
24 complexity of trace element concentrations in bivalve shells, which can then be applied as proxies for the  
25 reconstruction of sub-annual changes in palaeoenvironmental conditions over geological timescales.

26



## 1. Introduction

The study of climate and environmental change over geological timescales has yielded various important insights into the dynamics of climate systems on Earth (e.g. Huber et al., 1995; Hesselbo et al., 2000; Zachos et al., 2001; Wang et al., 2001; Sluijs et al., 2006). The knowledge about the sensitivity of Earth's climate and environment that results from these studies is indispensable for the prediction of future changes in Earth's climate. The study of environmental changes relies both on proxy-based palaeoenvironmental reconstructions (e.g. McDermott, 2004; Leng and Marshall, 2004; Zachos et al., 2006; Affek et al., 2008) and climate and environmental modelling based on, and calibrated with, these reconstructions (e.g. Barron et al., 1984; Kutzbach et al., 1989; Claussen et al., 2002; Andrews et al., 2012). Important archives for proxy-based reconstructions of palaeoenvironment on a high temporal resolution are fast-growing carbonate records such as speleothems and the fossil skeletal remains of calcifying organisms such as corals, brachiopods and molluscs (Goreau, 1977; de Winter and Claeys, 2017; Ullmann et al., 2017; Vansteenberghe et al., 2017; de Winter et al., 2017). Mollusc shells have gained much attention in the last decades, because the calcite in these shells has high fossilization potential, their populations are abundant and several studies have shown that chemical proxies in these shells record changes in their environment (e.g. Klein et al., 1997; Schöne et al., 2003; Lazareth et al., 2003; Gillikin et al., 2008). Stable isotope ratios of carbon and oxygen in the calcite shells of bivalve molluscs are almost exclusively precipitated in equilibrium with the surrounding seawater and can thus be used to trace changes in temperature, productivity and salinity on a seasonal scale (Klein et al., 1996; Kirby et al., 1998; Goodwin et al., 2001; Ullmann et al., 2010). However, to disentangle the effects of these parameters and to properly understand changes in the local environment, it is important to apply multi-proxy studies of shell calcite (e.g. Takesue and van Geen, 2004; Ullmann et al., 2013; de Winter et al., 2017). It is for this reason that bivalve sclerochronology studies have focused on understanding the relationships of trace element concentrations in bivalve calcite with their environment (Lorrain et al., 2005; Wanamaker et al., 2008; Freitas et al., 2009; Schöne et al., 2011). Since then, a range of trace element ratios (e.g. Mg/Ca, Sr/Ca, Ba/Ca, Mn/Ca and Li/Mg) have been used as proxies for environmental parameters (Klein et al., 1996a; Lazareth et al., 2003; Carré et al., 2006; Gillikin et al., 2008; Füllenbach et al., 2015; Vihtakari et al., 2017).

A few studies have focused on the development and chemical composition of modern oyster shells and its relation to the environment (e.g. Palmer and Carriker, 1979; Carriker et al., 1980; Lee et al., 2008; Surge and Lohmann, 2008; Ullmann et al., 2010; 2013). These studies have shown that oyster shells are composed mostly of calcite occurring as foliated calcite layers, prismatic calcite and chalky calcite while the myostracum and hinge ligament are made of aragonite (Stenzel, 1963; Palmer and Carriker, 1979). There is some discussion about the role of these calcite mineral phases, whether their precipitation is controlled by environmental conditions and whether changes in the precipitated mineral phase are paced to regular (solar or lunar) cycles (Carriker et al., 1980; Kirby et al., 1998; Surge et al., 2001; Ullmann et al., 2010). It has even been proposed that the mineralization of the chalky calcite phase in oyster shells is mediated by microbial activity (Vermeij, 2014). Beside mineralogy and chemistry of the shell, shell growth rate and dimensions vary widely between



63 individuals in response to several environmental factors such as growth space, substrate, food availability and  
64 amount of predation (Galtsoff, 1964; Palmer and Carriker, 1979; Surge and Lohmann, 2008).

65 All these physiological changes, such as variations in growth and metabolic rate, shell mineralogy and  
66 spawning events, which affect the incorporation of trace elements into the shell of bivalves, complicate the  
67 use of trace element records to complement environmental reconstruction by stable isotope  
68 sclerochronology (Klein et al., 1996b; Gillikin et al., 2005; Immenhauser et al., 2005; Freitas et al., 2006).  
69 Furthermore, several studies have shown that rates by which trace elements are incorporated into bivalve  
70 shells and the degree to which these rates are controlled by environmental factors can be vastly different  
71 between different bivalve species (Reinfelder et al., 1997; Steuber, 1999; Richardson et al., 2004; Carré et al.,  
72 2006). To constrain such variations in physiological parameters on the chemistry of bivalve shells, species-  
73 specific culture experiments are carried out under controlled circumstances so relationships between  
74 environmental parameters and shell chemistry can be precisely determined (e.g. Wang and Fisher, 1996;  
75 Freitas et al., 2006; Gillikin et al., 2006). Such experiments can only be executed on extant bivalves, which  
76 severely limits the potential to apply the acquired proxy transfer functions to reconstruct climate and  
77 environment in pre-Cenozoic times (e.g. de Winter et al., PPP). In this study, a model is introduced that  
78 approximates the development of a range of size parameters in the bivalve shell through ontogenetic age,  
79 based solely on digitized coordinates of recognized annual shell increments in a longitudinal cross section  
80 through the shell. Additionally, the modelled growth development and recruitment pattern in the shell cross  
81 section is then superimposed on an XRF trace element map to model the incorporation of trace elements into  
82 the shell with age. The application of this growth and trace element model is demonstrated on a set of shells  
83 of the modern pacific oyster (*Crassostrea gigas*) with well-known origins and dimensions. Model results are  
84 compared with conventional trace element analyses on line scans through the hinge of the shells as well as  
85 with results from previous bivalve growth studies.

## 86 2. Materials and Methods

### 87 2.1 Specimen acquisition and preparation

88 A set of eight modern pacific oyster (*Crassostrea gigas*) shells were obtained from restaurant Jardin van Gogh  
89 in Brussels, Belgium (<http://www.jardinvangogh.be>). The oysters originate from a cultivation area in coastal  
90 Normandy (France, 49°4.0' N latitude and 1°35.47' W longitude) and were harvested on February 13<sup>th</sup> 2017.  
91 The shells were rinsed with acetone (C<sub>3</sub>H<sub>8</sub>O) and distilled water, cleaned superficially with a soft brush and in  
92 an ultrasonic bath and oven-dried overnight at 50°C. Dried shells were weighed on a digital scales ( $\sigma = 0.01$  g),  
93 their dimensions (shell length, maximum shell width, maximum shell thickness) were measured using digital  
94 callipers ( $\sigma = 0.01$  mm) and their volume was determined by water displacement measurement using a graded  
95 cylinder. All shells were embedded in Araldite® 2020 epoxy resin (Huntsman, Basel, Switzerland), sectioned  
96 longitudinally along their axis of maximum growth using a slow rotating, diamond coated saw ( $\phi = 1$  mm) and  
97 high-grade polished using silicon carbide polishing disks (up to P2400 grain size). Polished shell surfaces were





98 imaged by colour scanning (RGB) using an Epson® 1850 flatbed scanner (Seiko Epson Corp., Nagano, Japan) at  
99 a pixel resolution of 6400 dpi ( $\pm 4 \mu\text{m}$  pixel size).

## 100 2.2 X-Ray Fluorescence measurements

101 Concentrations of calcium (Ca), silicon (Si), magnesium (Mg), strontium (Sr), zinc (Zn), sulphur (S), phosphorous  
102 (P), manganese (Mn) and iron (Fe) were measured on the polished shell surfaces using a Bruker® M4 Tornado  
103 micro-X-ray Fluorescence scanner (Bruker GmbH, Berlin, Germany) equipped with a Rh X-Ray source using  
104 maximum energy settings (50 kV, 600  $\mu\text{A}$ ) with a spot size of 25  $\mu\text{m}$  (Mo  $\text{K}\alpha$ ) and two Silicon Drift detectors.  
105 The XRF setup is described in detail in de Winter and Claeys (2017). The entire shell surface was mapped semi-  
106 quantitatively by XRF scanning in mapping mode using 1 ms integration time per pixel (as described in de  
107 Winter and Claeys, 2017). Spacing between pixels was chosen in such a way that the total amount of pixels per  
108 map was relatively constant ( $\pm 4.0 \cdot 10^6$ ) for all shells while choosing the minimum rectangular area that  
109 contained the entire cross section area. This caused pixel spacing in maps to vary between 25  $\mu\text{m}$  (interlocking  
110 X-Ray spots) and 30  $\mu\text{m}$ . Quantitative XRF line scans were carried out on the dense foliated calcite layers in the  
111 hinges of all shells perpendicular to the growth layers (see Palmer and Carriker, 1979) using the point-by-point  
112 scanning method outlined in de Winter et al. (in review, PPP) with an integration time of 60 seconds per point.  
113 This integration time allowed enough count statistics for the instrument to reach the Time of Stable  
114 Reproducibility (TSR) and provide reproducible trace element concentrations for the elements of interest (de  
115 Winter et al., 2017). All XRF line scans were quantified using the Bruker Esprit® fundamental parameters (FP)  
116 quantification relative to the BAS CRM 393 limestone standard. Errors of reproducibility of  $\mu\text{XRF}$   
117 measurements are generally higher than the instrumental error and depend on the integration time and the  
118 excitation energy of the element (see de Winter and Claeys, 2017; de Winter et al., 2017). Typical  
119 reproducibility errors of  $\mu\text{XRF}$  point measurements are reported in Table 3.

## 120 2.3 Data preparation

121 Concentrations (in  $\mu\text{g/g}$ ) of trace elements were calculated ~~for profiles measured using XRF~~ and plotted using  
122 Grapher™ 8 (Golden Software Inc., Golden, CO, USA) graphing software. Timing of shell deposition was  
123 inferred from annual cyclicity in trace element profiles. Growth increments (lines of simultaneous deposition in  
124 the shell) were digitized on high-resolution colour scans of polished shell cross sections of the shells using the  
125 pen tool in Adobe Illustrator® CC 17.1.0 (Adobe Systems Inc., San Jose, USA). Outlines of the rectangular area  
126 of the cross section mapped by XRF were digitized in the same way. Line coordinates were saved in a SVG-file,  
127 which allowed X-and Y-coordinates of the lines to be extracted and ordered into a comma separated (CSV) file  
128 to be imported into the modelling script (Step 1 in Fig. 1). An example of a shell cross section with traced  
129 growth increments is shown in Figure 2. SVG- and CSV files of growth increments digitized in all shells used in  
130 this study are found in supplementary data 1.

131 XRF map data was processed using Bruker Esprit® software. Maps were subject to a PCA-assisted maximum  
132 likelihood phase analysis using a selection of distinctive elements (Ca, Mg, Sr, P, S and Mn). Minimum phase



area was fixed to 0.05% of the total map area. Phase analysis results were matched with interpreted growth increments and high-resolution colour scans and associated phases were merged. Sum XRF spectra of all pixels in each phase were quantified relative to the BAS CRM 393 standard using Esprit® software. Minimum area of each phase was such that the total integration time contained in all the pixels allowed Time of Stable Reproducibility to be reached for the quantification of the sum spectrum (de Winter et al., 2017). Phase maps were exported as BMP files and oriented in the same way as colour scans with the shell hinge facing left and the inside of the shell facing down (see Fig. 2).

## 2.4 Modelling approach

A modelling routine was written in the open source computational software package R (R Core Team, 2013) using Microsoft® Visual Studio Code Version 1.10.2. Shell growth and trace element accumulation rates were modelled in six steps, of which **Step 1** is a data preparation step (see above), **Step 2-4** form the growth model and **Step 5-6** make up the trace element model (Fig. 1). The complete R-script used for the model is provided in **supplementary data 2**. Variables used in the modelling process are indicated in **Figure 3**.

146

### 2.4.1 Growth modelling

**Step 2** of the model converts X- and Y-coordinates of all digitized increments to millimetres using the ratio between the real length of the digitized image and the length in pixels. All increments are converted to one cross section matrix (*Digitized cross section*) with a common X-axis with a default step size ( $dx$ ) using linear interpolation between line segments. The resulting cross section is plotted to provide a check on the model progress. From this matrix, the area between each increment and its predecessor is calculated using the formula:

$$\mathbf{F1: } O_i = \int_{x_0}^{x_{end}} Y_{i-1}(x) - Y_i(x) dx \quad (1)$$

in which  $O_i$  is the area between increment  $i$  and increment  $i-1$ ,  $x_0$  and  $x_{end}$  are the extreme values of the range of X coordinates in *Digitized cross section* and  $Y_i$  and  $Y_{i-1}$  are the Y-coordinates of increment  $i$  and  $i-1$  respectively.  $Y_{i-1}$  is always bigger than  $Y_i$  since ~~bivalves build their shell by adding material on the inside of the shell,~~ which faces down in this model. The average shell thickness at each increment is determined using the formula:

$$\mathbf{F2: } T_i = \frac{\sum_{x_0}^{x_{end}} Y_0(x) - Y_i(x)}{x_{end} - x_0} \quad (2)$$

in which  $T_i$  is the average thickness of the shell at increment  $i$ ,  $Y_i(x)$  and  $Y_0(x)$  are the Y-coordinates of increment  $i$  and the top of the shell (increment 0). Total shell ~~length~~ is calculated from the X- and Y-coordinates of the start- and endpoints of the increment (where the increment meets the top or bottom of the shell) and the Pythagorean Theorem following the formula:

$$\mathbf{F3: } L_i = \sqrt{(x_e - x_s)^2 + (Y_e - Y_s)^2} \quad (3)$$



in which  $L_i$  is the length of the shell at increment  $i$ ,  $x_s$ ,  $x_e$ ,  $Y_s$  and  $Y_e$  are the X- and Y-coordinates of the start- and endpoints of the increment  $i$ . The results of these calculations, as well as values for  $x_s$ ,  $x_e$ ,  $Y_s$  and  $Y_e$  are stored in a *Matrix of parameters by increment*.

**Step 3** of the growth model takes *Digitized cross section*, *Matrix of parameters by increment* and a customizable number of increments ( $N$ ) to be interpolated to create a new cross section matrix with  $N-1$  **interpolated sub-increments** between each set of increments (*Sub-incremental cross section*). Interpolation of sub-increments is done by linear interpolation of the Y-coordinate of sub-increments between that of the two increments (see insert in **Figure 3A**) according to the following formula:

$$\text{F4.1: } \left[ Y_t(x) = Y_{i-1}(x) - \frac{n}{N} * (Y_{i-1}(x) - Y_i(x)) \right]_{n=0}^{n=N-1} \Bigg|_{x=x_0}^{x=x_{end}} \quad (4)$$

With

$$\text{F4.2: } t = i - 1 + \frac{n}{N} \quad (5)$$

in which  $Y_t(x)$  is the Y-coordinate of the  $n^{\text{th}}$  sub-increment between increment  $i$  and increment  $i-1$  and  $x_0$  and  $x_{end}$  are the extreme values of range of X-values in *Digitized cross section* (as in **F1**). All calculated values for  $Y_t(x)$  are stored with reference to their sub-increment number ( $t$ ) and X-coordinate in the new *Sub-incremental cross section* matrix. The resulting cross section is plotted to provide a check on the model progress. This new matrix is then used to calculate area between sub-increments, shell thickness and total shell length during deposition of each sub-increment by formula **F1**, **F2** and **F3** respectively. Additionally, using the measured maximum length (*Shell length*) and width (*Shell width*) of the oyster, parameters  $a$  and  $b$  of the ellipse that forms the base of the shell for volume calculations (**Figure 3B**) are calculated according to formulae:

$$\text{F5.1: } a_t = \frac{1}{2} * \frac{W_{max}}{L_{max}} * (x_e - x_s) \quad (6)$$

$$\text{F5.2: } b_t = \frac{1}{2} * (x_e - x_s) \quad (7)$$

in which  $a_t$  and  $b_t$  are the parameters  $a$  and  $b$  of the ellipse that forms the base of the shell at sub-increment  $t$ ,  $L_{max}$  and  $W_{max}$  are the maximum length and width of the oyster shell and  $x_s$  and  $x_e$  are the X-coordinates of the start- and endpoints of the increment  $t$  (as in **F3**). All these parameters are stored in *Matrix of parameters by sub-increment* (**Figure 1**).

**Step 4** takes *Incremental cross section* and the ellipse parameters in the *Matrix of parameters by sub-increment* to calculate the Z-values of the ellipse that forms the base of the shell at each sub-increment (see **Figure 3B**). The Z-value is defined as the distance between the edge of the ellipse and the X-axis through the shell (**Figure 3B**), and is calculated by the following formula, which is an adaptation of the standard formula for ellipsoids:

$$\text{F6.1: } \left( \frac{Z_t(x)}{a_t} \right)^2 + \left( \frac{x_t}{b_t} \right)^2 = 1 \rightarrow$$



$$197 \quad \text{F6.2: } Z_t(x) = \left(\frac{a_t}{b_t}\right) * \sqrt{b_t^2 - x_t^{*2}} \quad (8)$$

198 in which  $Z_t$  is the Z-value (distance from X-axis) of the ellipse at X-coordinate  $x_t^*$  for sub-increment  $t$ ,  $a_t$  and  $b_t$   
 199 are the parameters of the ellipse at sub-increment  $t$  and  $x_t^*$  is the X-coordinate relative to the centre of the  
 200 ellipse, and is calculated by

$$201 \quad \text{F6.3 } x_t^* = x - x_s - b_t \quad (9)$$

202 All Z-values are saved in a matrix (*Z-values* in **Figure 1**) with reference to their increment numbers ( $t$ ) and X-  
 203 coordinates. Then, using the Z-values and the parameters from *Matrix of parameters by sub-increment*, shell  
 204 volume is calculated for each sub-increment. This is done by calculating areas between the sub-increment and  
 205 the top of the shell (sub-increment 0) in a cross sections through the shell perpendicular to the X-axis (in YZ-  
 206 plane, see **Figure 3C**) and multiplying these with the step size in X-direction ( $dx$ ). This is done for every X-value,  
 207 and adding up all volume increments yields an estimate the total volume between the shell sub-increment and  
 208 the base of the shell:

$$209 \quad \text{F7.1: } V_t = \int_{x_0}^{x_{end}} (A_0(x) - A_t(x)) dx \quad (10)$$

210 in which  $V_t$  is the volume of the shell at increment  $t$  and  $A_0(x)$  and  $A_t(x)$  are the area under increment  $t$  and  
 211 the top of the shell (increment 0) respectively in the cross section in YZ-direction (**Figure 3C**). These areas are  
 212 modelled for every X-value by constructing a circle section through the corresponding point on the sub-  
 213 increment in the XY cross section (centre of the YZ-cross section **Figure 3A**, or point **P1**)

$$214 \quad \text{P1}[x, y_1, z_1] = [x, Y_t(x), 0] \quad (11)$$

215 and the two points where the YZ-cross section intersects the ellipse that forms the base of the shell (see **Figure**  
 216 **3A** and **Figure 3C**):

$$217 \quad \text{P2}[x, y_2, z_2] = [x, Y_{\text{ellipse}}(x), -Z_t(x)] \quad (12)$$

$$218 \quad \text{P3}[x, y_3, z_3] = [x, Y_{\text{ellipse}}(x), Z_t(x)] \quad (13)$$

219 The value  $Y_{\text{ellipse}}(x)$  is the Y-value of the ellipse with respect to the line  $y=0$  (**Figure 3C**), which can be calculated  
 220 by linear interpolation of the slope of the ellipse using the start and end points of the sub-increment ( $x_s$ ,  $x_e$ ,  $Y_s$   
 221 and  $Y_e$ ) and  $x$ :

$$222 \quad \text{F7.2: } Y_{\text{ellipse}} = Y_s + \left(\frac{Y_e - Y_s}{x_e - x_s}\right) * (x - x_s) \quad (14)$$

223 The centre of this circle is the point

$$224 \quad \text{Pc}[x, y_c, z_c] = [x, Y_c, 0] \quad (15)$$

225 its radius  $r$  is equal to the difference between  $Y_t(x)$  and  $Y_c$ , and the circle can be described by the formulae:

$$226 \quad \text{F7.3: } \Delta y^2 + \Delta z^2 = r^2 \rightarrow$$



$$\text{F7.4: } (y_1 - y_c)^2 + (z_1 - z_c)^2 = (y_3 - y_c)^2 + (z_3 - z_c)^2 \rightarrow$$

$$\text{F7.5: } (Y_t(x) - Y_c(x))^2 + (0 - 0)^2 = (Y_{\text{ellipse}}(x) - Y_c(x))^2 + (Z_t(x) - 0)^2 \rightarrow$$

$$\text{F7.6: } Y_c(x) = \frac{(Y_t(x)^2 - Y_{\text{ellipse}}(x)^2 - Z_t(x)^2)}{2 * (Y_t(x) - Y_{\text{ellipse}}(x))} \quad (16)$$

$$\text{F7.7: } r_t(x) = \sqrt{Z_t(x)^2 + (Y_{\text{ellipse}}(x) - Y_c(x))^2} \quad (17)$$

With all parameters known, the area in the YZ-cross section under the sub-increment (between the circle segment and the line  $y=0$ ) can be calculated as the area of the section of the circle above the ellipse plus the area of the rectangle between the ellipse and  $y=0$  (**Figure 3C**). The angle  $\theta$  describing this circle section is equal to:

$$\text{F7.8: } \theta = 2 * \sin^{-1} \left( \frac{Z_t(x)}{r_t(x)} \right) \quad (18)$$

However, if point **P1** lies below the ellipse ( $Y_t(x) < Y_{\text{ellipse}}(x)$ ); in the case of irregular shells that curve upwards during growth), the centre of the circle lies above the shell and the area under the sub-increment is described by subtracting the section of the circle above the ellipse from the area of the rectangle (see **Figure 3C**):

$$\text{F7.9: } A_t = \begin{cases} A_{\text{segment}} + A_{\text{rectangle}} = \frac{1}{2} (r_t(x))^2 * (\theta - \sin \theta) + 2 * Z_t(x) * Y_{\text{ellipse}}(x), & Y_t(x) \geq Y_{\text{ellipse}}(x) \\ A_{\text{rectangle}} - A_{\text{segment}} = 2 * Z_t(x) * Y_{\text{ellipse}}(x) - \frac{1}{2} (r_t(x))^2 * (\theta - \sin \theta), & Y_t(x) < Y_{\text{ellipse}}(x) \end{cases} \quad (19)$$

Net areas are calculated as the differences between the areas under the sub-increment  $t$  and the area under the top of the shell (sub-increment 0), and volumes for sub-increments are calculated by integrating these areas over  $x$  (see formula **F7.1** above). Shell growth rates are then calculated by multiplying the change in volume per sub-increment with *Shell Density* ( $\rho$ ):

$$\text{F8.1: } \Delta M_t = \rho * (V_{t-1} - V_t) \quad (20)$$

and absolute mass accumulation is calculated by simple multiplication of the modelled incremental volume increase of the shell with **Shell Density**:

$$\text{F8.2: } M_t = \rho * V_t \quad (21)$$

## 2.4.2 Trace element modelling

**Step 5** of the model takes the BMP-file of the *Phase map* of the shell and a matrix of the quantified concentrations of all elements of interest in each of the phases as well as their colour values (*Phase data*, **Figure 1**) as input to convert the BMP image to a matrix of phases (*Phase matrix*, **Figure 1**). This matrix is then used to export a table with statistics of the relative abundance of phases in the entire phase map (*Phase statistics*, **Figure 1**). *Phase data* tables used as input to model every shell described in this study are given in **supplementary data 3**.



**Step 6** uses this *Phase matrix* together with *Incremental cross section* to calculate the amount of pixels of each phase that is contained in every sub-increment (*Sub-increment phase matrix*, **Figure 1**). From this data, the concentration of each element in each sub-increment are calculated by multiplying the relative proportion of each phase in the sub-increment by the quantified concentrations of all elements in that phase:

$$\text{F9: } C_t^E = \sum_{p=p_1}^{p=p_n} \frac{S_p}{S_{\text{tot}}} * C_p^E \text{ with } p \text{ in } [p_1, p_2, p_3, \dots, p_n] \quad (22)$$

where  $C_t^E$  is the concentration of element  $E$  in sub-increment  $t$ ,  $p$  is the phase (in  $p_1, p_2, p_3, \dots, p_n$ ),  $S_p$  is the amount of pixels assigned to phase  $p$  in sub-increment  $t$ ,  $S_{\text{tot}}$  is the total amount of pixels in sub-increment  $t$ , and  $C_p$  represents the concentration of element  $E$  in phase  $p$ . The distribution of trace element concentrations in each sub-increment is stored in *Matrix of concentration through time*. This matrix is then multiplied with a smoothed record of modelled mass accumulation and growth rates (see **Step 4**, smoothing occurs via a running average over the mass accumulation and growth rate records and the *Degree of smoothing* is customizable and determines the window size of this running average) to calculate the cumulative accumulation and accumulation rates, respectively, of all (trace) elements through time in the shell:

$$\text{F10.1: } M_t^E = C_t^E * M_t \quad (23)$$

$$\text{F10.2: } \left[ \frac{\partial M}{\partial t} \right]_t^E = C_t^E * N * \Delta M_t \quad (24)$$

Matrices of modelled elemental concentrations (*Matrix of concentrations through time*), cumulative trace element accumulation (*Cumulative elemental mass accumulation*) and accumulation rates (Elemental mass accumulation rate) modelled through the shell's age are exported for further analysis. An overview of all model functions and variables is given in **Table 1**. Exported matrices containing modelling results for every shell featuring in this study are presented in **supplementary data 4**

### 3. Results and discussion

#### 3.1 XRF and shell dimension measurements

Shells of *C. gigas* are highly irregular with considerable differences in shape between individuals, as is evident from measurements of the shell dimensions (**Table 2**) and the colour scans of the shells (**Figure 2** and **supplementary data**). Shell length, width, volume and mass vary considerably between *C. gigas* specimens and estimated age based on proxy records is not a good predictor of shell size. Furthermore, the length-to-width ratio is highly variable between shells, making size development in *C. gigas* hard to predict. Densities of *C. gigas* shells are relatively low ( $\rho = 2.10 \text{ g*cm}^{-3}$ ) compared to the densities of shell-forming minerals such as calcite ( $\rho = 2.71 \text{ g*cm}^{-3}$ ), aragonite ( $\rho = 2.95 \text{ g*cm}^{-3}$ ) and nacre ( $\rho = 2.60 \text{ g*cm}^{-3}$ ). This difference is most likely caused by the presence of porosity in the shell structure, which should be around 23% to account for the difference in shell density.

**Figure 2** shows the result of colour scanning, XRF mapping with phase analysis and a tracing of the growth increments in a longitudinal cross section through one of the *C. gigas* shells. The shell depicted in **Figure 2** is



used as an example for the remainder of the results and discussion, while outcomes for the remaining seven shells are disclosed in in **supplementary data 5**. **Figure 2** shows that phase analysis on the XRF map of oyster shells results in the separation of **four chemically distinct phases** in the cross section. Comparison with the colour scan shows that these phases represent dark foliated calcite layers in the shell (green), light chalky calcite layers in the shell (blue), detrital inclusions **in the edge of the shell** (yellow) and the surrounding resin (red). Trace element concentrations of the foliated and chalky calcite phases in each shell are found in **Table 3**. ~~Mapping and phase analysis in all shells resulted in a distinction between foliated calcite and chalky calcite layers in terms of chemical composition~~ (see **Figure 2** and compare chemical compositions in **Table 3**). The phase map in **Figure 2** also shows that the hinge of the shell consists of **foliated calcite**. Traces of growth increments in the oyster shell show once more that growth patterns in *C. gigas* are highly irregular. **While shell growth always happens by addition of material on the inside of the shell valve (facing down in Figure 2), shell thickness varies strongly throughout the shell and shell extension rates vary both with age and with location in the shell. Furthermore, oyster shells extend away from the shell hinge (to the right in Figure 2) as well as towards the inside of the shell, making the hinge thicker with age (downward and to the left in Figure 2). These shell characteristics complicate the modelling of shell growth and render *C. gigas* an ideal species for rigorous testing of the model presented in this study.**

Results of line scanning through the hinge of the oyster are shown in **Figure 4**. Shells of *C. gigas* are characterized by **periodic** variations in concentrations of strontium (Sr), magnesium (Mg), sulphur (S), iron (Fe), manganese (Mn) and zinc (Zn). Records of silicon (Si) and calcium (Ca) concentrations indicate which parts of the records represent pure shell calcite (high [Ca], low [Si]) and which consist of **calcite diluted with detrital material** (lower [Ca], [Si] > 2000 µg/g, mostly on the outside of the shell, see **Fig. 4**). **Shell increments** used as tracers for growth modelling are generally characterized by decreased Ca and Mg concentrations and increased concentrations of Fe, Mn, Zn and Sr. **Furthermore, records of Sr and Zn show regular cyclicity**, while Fe and Mn records are characterized by sharp increases relative to a stable baseline. The Mg record shows small scale variations inversely related to those in the Zn record. **Periodic variations in the trace element records allow the establishment of an age model for the growth of this oyster shell, as is shown in Figure 4.** **Note that line scanning through the hinge of the shell only allows for the sampling of the last three growth years, as the irregular shape of the oyster shell and the occurrence of chalky calcite further up the hinge prevents the measurement of a complete record through the foliated calcite.** Also note that growth increments used as a basis for growth modelling are not paced to the seasonal cycle. The organisation of **isochronous growth increments** in the colour scan on top of **Figure 4** shows the occurrence of chalky calcite layers embedded between foliated calcite layers in some parts of the shell while these are absent in other parts. This further confirms that multiple types of shell mineral phases (e.g. foliated calcite and chalky calcite) can be precipitated in the shell simultaneously. Since mineral phases are chemically distinct (**Table 3**), this observation warrants the consideration of the growth of both shell phases in an analysis of trace element uptake by oyster shells, showing that simply analysing foliated calcite in the hinge of the shell may not yield a complete understanding of the incorporation of trace elements into the shell.



### 326 3.2 Growth model



327 The output of the growth model applied on cross sections of *C. gigas* shells consists of records of shell length,  
 328 average thickness, volume, mass and growth rates through shell age (**Figure 5**). Tables containing the complete  
 329 records of all these parameters for all shells are given in **supplementary material 4**, and modelled shell  
 330 dimensions at the end of the modelling run are given in **Table 2**. **Figure 5** shows the records for the above  
 331 mentioned shell parameters plotted against age following the **age model based on line scans** through the shell  
 332 hinge. **The results show that, though there is ample variation in size development between individuals, the**  
 333 **development of shell size parameters follow a similar pattern in all the examined shells. Development of shell**  
 334 **length in all modelled shells follows the asymptotic Von Bertalanffy growth model** ( $L = L_{\infty} * e^{-kt}$ ; Von  
 335 Bertalanffy, 1957). Parameters of Von Bertalanffy models ( $k$  and  $L_{\infty}$ ) fitted to shell length records of each shell  
 336 are given in **Table 2**. Results show that, while Bertalanffy curves fit the shell length development very well ( $R^2 >$   
 337 0.90 for most shells except for #3 and #4), Bertalanffy's  $K$  values ( $k$ ) as well as maximum shell lengths ( $L_{\infty}$ ) vary  
 338 strongly. This result is unsurprising for oyster shells, which are known to show large variations in growth rate  
 339 and morphology in response to local environmental constraints on their growth (Galtsoff, 1964; Palmer and  
 340 Carriker, 1979). Curve fitting through a composite of all *C. gigas* shells yields a maximum shell length of 102.34  
 341 mm and the growth curve constant (Bertalanffy's  $K$ ) of 0.99. The values for maximum shell length are  
 342 significantly lower than the value found for sister-taxa *C. virginica* (150 mm; Rothschild et al., 1994), but this  
 343 may be a result of the use of relatively young individuals in this study. The fact that the obtained results seem  
 344 to fit the Von Bertalanffy model well ( $R^2 = 0.60$  for all shells combined, see **Table 3** and **Figure 5**) shows that  
 345 the shell growth results produced by the model are reasonable, because it is known that the Von Bertalanffy  
 346 growth model describes shell length in most bivalves. The values for Bertalanffy's  $K$  fitting the model results  
 347 are quite high compared to most bivalve growth studies (e.g. Bachelet, 1980; MacDonald and Thompson,  
 348 1985; Hart and Chute, 2009), but values greater than 1 are not unheard of in bivalve species that show steep  
 349 growth curves early in life (e.g. Urban, 2000; Richardson et al., 2004). Modelled shell lengths closely resemble  
 350 those measured on the shell, with an average offset of 0.16 mm (0.16% relative to average shell length, see  
 351 **Table 2**) and are in good agreement with shell length measurements of living specimens of *C. gigas* (Diederich,  
 352 2006).

353 The development of other growth parameters shows similar variation within the same pattern of development  
 354 between individuals of *C. gigas*, attesting to the reproducibility of the growth model. For example, the average  
 355 shell thickness of oyster shells is best described by a linear increase in thickness with age (**Figure 5**). Individual  
 356 results show that the initial increase in thickness (slope of the average shell thickness curve) is quite variable,  
 357 but that later in life the different individuals of *C. gigas* converge towards a similar average shell thickness. This  
 358 results in rather variable rates of shell thickness increase between individuals (0.54–1.61 mm/yr, see **Table 2**).  
 359 The convergence of the shell thickness curves at later age suggests that this range is biased by the use in this  
 360 study of relatively young individuals. These differences in the development of shell thickness in oyster shells  
 361 are likely to be a result of spatial constraints on shell growth (Bartol et al., 1999). The agreement between the  
 362 final thicknesses of individuals is quite remarkable given their irregular shell shape and vastly different





proportions of shell length and width (**Table 2**). Maximum thickness (thickness of the thickest part of the shell) is not modelled and therefore cannot be compared with measured values in **Table 2**, but modelled average thicknesses are in agreement with observations in the cross section, and are proportional to measured maximum thickness of the shells.

Modelled shell mass and volume development are best approximated by a polynomial increase with shell age that is in agreement with the linear increase observed in modelled growth rates of *C. gigas*, which is naturally the derivative of shell mass development (**Figure 5**). Modelled shell volume and mass at the end of the shell's lifetime generally underestimate measured volume and mass by 4.2 cm<sup>3</sup> and 9.0 g respectively ( $\pm 21\%$ , see **Table 2**). The most likely reason for this offset is that the assumption of a semi-circular shape of the YZ-plane cross-section through the shell (perpendicular to the growth axis, see **Figure 3C**) underestimates the area of this cross section. In reality, the decrease in shell thickness towards the outer margins of the shell is probably less pronounced. Trends in volume and growth rates are less reproducible between individuals than those in shell length and shell thickness, as is evident from the diverging polynomial fits that fit the model data. This behaviour illustrates the erratic growth of *C. gigas* shells, which is also evident from the shape of their shell (**Figure 2** and **supplementary data 5**). As is shown by the modelled growth rate curves (**Figure 5**), the growth of these oyster shells is characterized by short-lived increases in growth rate followed by periods of slower, more constant shell growth. The implications of these periodic growth spurts punctuating more constant growth rates are also visible in the shell volume curves that often show stepwise increases in shell volume. To a lesser extent, the same periodic growth is seen in the records of shell length and thickness. On a closer examination, periods of faster growth rates can be associated with either contemporary increases in shell length or in shell thickness, but rarely both at the same time. This strongly suggests a control of available growth space on the shape and size development of *C. gigas* shells in competition with other individuals in an oyster reef (e.g. Palmer and Carriker, 1979; Bartol et al., 1999). On the other hand, food availability is known to significantly affect growth rates in bivalve shells (Kerswill, 1949; Coté et al., 1994; Surge and Lohmann, 2008), showing that peaks in growth rate found by the model results in this study may be attributed to short-lived increases in food availability commonly associated with algal blooms in spring and autumn in the region of study (Edwards et al., 2001; Wiltshire et al., 2008). This reliance of shell growth on environmental factors illustrates the potential of these model results to aid in the reconstruction of environmental conditions.

### 3.3 Trace element model

Records of trace element accumulation rates and total shell trace element concentrations that result from trace element modelling are plotted for one of the *C. gigas* shells in **Figure 6** together with concentrations in the hinge of the shell measured using XRF line scanning. Records of accumulation rates of different elements show similar trends during shell growth and correlate with changes in shell growth rates. For some elements (e.g. Zn and S) the total shell concentrations over time resemble concentrations measured in the hinge of the shell, while for other elements (e.g. Mg and Sr) the total shell concentrations show a very different pattern from the measured concentrations in the foliated calcite in the shell hinge. The reason for this difference is that some elements (e.g. Zn and S) have very similar concentrations in the foliated calcite and the chalky



400 calcite layers, whereas these concentration can be very different for other elements (e.g. Mg and Sr; see **Table**  
401 **3**). Since the type of mineral phase deposited during shell growth is not controlled by growth seasonality or  
402 age (Surge et al., 2001; Titschack et al., 2008), differences in the degree of incorporation of mineral phases  
403 over time will result in different total shell concentrations. These differences in total shell concentrations and  
404 concentrations in the shell hinge illustrate the value of the proposed trace element modelling approach, as  
405 concentrations taken up in the shell are better reflected by total shell concentrations than by concentrations in  
406 one of the mineral phases in the shell. Furthermore, in combining the results of trace element modelling with  
407 measurements on the shell hinge of bivalves, it is possible to constrain the relative amount of each mineral  
408 phase that is incorporated into the shell at any given time. This allows the reconstruction of changes in shell  
409 mineralogy and help isolate of the factors that control these changes, which is an important question in the  
410 study of oyster growth (e.g. Currey and Taylor, 2000; Surge et al., 2001; Titschack et al., 2008; Beniash et al.,  
411 2010).

#### 412 **4. Conclusions and outlook**

413 This study proposes a new method of modelling the growth, **development** and trace element incorporation in  
414 bivalve shell based on the location of growth increments in a cross section of the shell. The advent of a  
415 working model that can independently constrain growth and trace element uptake rates would greatly benefit  
416 the field of bivalve sclerochronology by **providing independent control on shell growth rates, which influence**  
417 **the expression of geochemical proxies in the shell**. This development is especially interesting for studies  
418 dealing with extinct bivalve species for which there are no modern analogues. The basic assumptions of the  
419 model render it applicable on all bivalve species with the same general shape and growth direction. Growth  
420 modelling following this numerical approach yields curves of shell development with age that resemble growth  
421 curves established via *in vivo* measurements and allows the discussion of differences in growth and  
422 development within and between bivalve species. **The present modelling approach allows the comparison of**  
423 **growth and development of bivalve shells on a sub-annual scale without a priori knowledge about growth**  
424 **rates in the species, opening up the comparison of proxy records in fossil bivalves with records of growth rate**  
425 **derived by applying this model. This allows the discussion of the applicability of trace element concentrations**  
426 **as direct tracers of environmental change as opposed to being controlled by physiological processes related to**  
427 **shell growth.**

428 The combination of growth modelling with 2D trace element XRF mapping allows the projection of trace  
429 element distribution to a 3D model of shell volume to numerically model the total shell concentration and  
430 accumulation of trace elements into bivalve shells. **Comparison between modelled total shell trace element**  
431 **content and concentrations measured along the growth axis in the shell hinge following a conventional**  
432 **measurement protocol reveals different patterns in trace element concentrations.** This shows that  
433 conventional trace element profiles through the shell hinge, recording only a small part of the shell, are not  
434 always representative for total shell concentrations and that modelling these concentrations may shed more  
435 light on the incorporation of trace elements into bivalve calcite. Further research should therefore consist of  
436 applying this modelling approach in other bivalve studies to compare modelled and measured trace element



concentrations. According to the results presented here, studies focusing on establishing trace element proxy transfer functions could benefit from basing their regressions on total shell trace element concentrations rather than measurements in the shell hinge in their attempts to isolate environmental controls on trace element concentrations in bivalve calcite.

441

#### 442 Code availability

The R script of the ShellTrace model used in this paper was published in the open source research data repository Zenodo (<http://doi.org/10.5281/zenodo.817258>). The complete script used for the ShellTrace model in this publication will be made available by means of an R package in the CRAN repository (<https://cran.r-project.org>), and the script is given in **supplementary data 2**.

447

#### 448 Acknowledgments

Thanks go to restaurant Jardin van Gogh in Brussels (<http://www.jardinvangogh.be>) for kindly supplying the author with the pacific oysters of which the shells were used for this research project, and for disclosing information about the origin of the bivalves. The author is financed by a personal PhD fellowship from IWT Flanders (IWT700). Thanks go to the Hercules foundation Flanders for acquisition of XRF instrumentation (grant HERC1309) and VUB Strategic Research for support of the AMGC research group. The author declares that there are no conflicts of interest.

455

#### 456 References

- Affek, H.P., Bar-Matthews, M., Ayalon, A., Matthews, A., Eiler, J.M., 2008. Glacial/interglacial temperature variations in Soreq cave speleothems as recorded by “clumped isotope” thermometry. *Geochimica et Cosmochimica Acta* 72, 5351–5360.
- Andrews, T., Gregory, J.M., Webb, M.J., Taylor, K.E., 2012. Forcing, feedbacks and climate sensitivity in CMIP5 coupled atmosphere-ocean climate models. *Geophysical Research Letters* 39.
- Bachelet, G., 1980. Growth and recruitment of the tellinid bivalve *Macoma balthica* at the southern limit of its geographical distribution, the Gironde estuary (SW France). *Marine Biology* 59, 105–117.
- Barron, E.J., Washington, W.M., 1984. The role of geographic variables in explaining paleoclimates: Results from Cretaceous climate model sensitivity studies. *Journal of Geophysical Research: Atmospheres* 89, 1267–1279.
- Bartol, I.K., Mann, R., Luckenbach, M., 1999. Growth and mortality of oysters (*Crassostrea virginica*) on constructed intertidal reefs: effects of tidal height and substrate level. *Journal of Experimental Marine Biology and Ecology* 237, 157–184.
- Beniash, E., Ivanina, A., Lieb, N.S., Kurochkin, I., Sokolova, I.M., 2010. Elevated level of carbon dioxide affects metabolism and shell formation in oysters *Crassostrea virginica*. *Marine Ecology Progress Series* 419, 95–108.
- Carré, M., Bentaleb, I., Bruguier, O., Ordinola, E., Barrett, N.T., Fontugne, M., 2006. Calcification rate influence on trace element concentrations in aragonitic bivalve shells: evidences and mechanisms. *Geochimica et Cosmochimica Acta* 70, 4906–4920.
- Carriker, M.R., Palmer, R.E., Sick, L.V., Johnson, C.C., 1980. Interaction of mineral elements in sea water and shell of oysters (*Crassostrea virginica* (Gmelin)) cultured in controlled and natural systems. *Journal of experimental marine biology and ecology* 46, 279–296.
- Claussen, M., Mysak, L., Weaver, A., Crucifix, M., Fichet, T., Loutre, M.-F., Weber, S., Alcamo, J., Alexeev, V., Berger, A., others, 2002. Earth system models of intermediate complexity: closing the gap in the spectrum of climate system models. *Climate dynamics* 18, 579–586.
- Côté, J., Himmelman, J.H., Claereboudt, M.R., 1994. Separating effects of limited food and space on growth of the giant scallop *Placopecten magellanicus* in suspended culture. *MARINE ECOLOGY-PROGRESS SERIES* 106, 85–85.
- Currey, J.D., Taylor, J.D., 1974. The mechanical behaviour of some molluscan hard tissues. *Journal of Zoology* 173, 395–406.
- de Winter, N.J., Claeys, P., 2016. Micro X-ray fluorescence (μXRF) line scanning on Cretaceous rudist bivalves: A new method for reproducible trace element profiles in bivalve calcite. *Sedimentology*. doi:10.1111/sed.12299
- de Winter, N., Sinnesael, M., Makarona, C., Vansteenberge, S., Claeys, P., 2017. Trace element analyses of carbonates using portable and micro-X-ray fluorescence: Performance and optimization of measurement parameters and strategies. *Journal of Analytical Atomic Spectrometry* 32, 1211–1223.
- de Winter, N.J., Goderis, S., Dehairs, F., Jagt, J.W.M., Fraaije, R.H.B., Van Malderen, S.J.M., Vanhaecke, F., Claeys, P., Tropical seasonality in the late Campanian (Late Cretaceous): Comparison between multiproxy records from three bivalve taxa from Oman. *Palaeogeography, Palaeoclimatology, Palaeoecology*, in review.



- Diederich, S., 2006. High survival and growth rates of introduced Pacific oysters may cause restrictions on habitat use by native mussels in the Wadden Sea. *Journal of Experimental Marine Biology and Ecology* 328, 211–227.
- Edwards, M., John, A.W.G., Johns, D.G., Reid, P.C., 2001. Case history and persistence of the non-indigenous diatom *Coscinodiscus wailesii* in the north-east Atlantic. *Journal of the Marine Biological Association of the UK* 81, 207–211.
- Freitas, P.S., Clarke, L.J., Kennedy, H., Richardson, C.A., Abrantes, F., 2006. Environmental and biological controls on elemental (Mg/Ca, Sr/Ca and Mn/Ca) ratios in shells of the king scallop *Pecten maximus*. *Geochimica et Cosmochimica Acta* 70, 5119–5133.
- Freitas, P.S., Clarke, L.J., Kennedy, H., Richardson, C.A., others, 2009. Ion microprobe assessment of the heterogeneity of Mg/Ca, Sr/Ca and Mn/Ca ratios in *Pecten maximus* and *Mytilus edulis* (bivalvia) shell calcite precipitated at constant temperature. *Biogeosciences Discussions* 6, 1267.
- Füllenbach, C.S., Schöne, B.R., Mertz-Kraus, R., 2015. Strontium/lithium ratio in aragonitic shells of *Cerastoderma edule* (Bivalvia)—A new potential temperature proxy for brackish environments. *Chemical Geology* 417, 341–355.
- Galtsoff, P.S., 1964. The American Oyster: US Fish and Wildlife Service. *Fishery Bulletin* 64, 480.
- Gillikin, D.P., Lorrain, A., Bouillon, S., Willenz, P., Dehairs, F., 2006. Stable carbon isotopic composition of *Mytilus edulis* shells: relation to metabolism, salinity,  $\delta^{13}\text{C}$  DIC and phytoplankton. *Organic Geochemistry* 37, 1371–1382.
- Gillikin, D.P., Lorrain, A., Navez, J., Taylor, J.W., André, L., Keppens, E., Baeyens, W., Dehairs, F., 2005. Strong biological controls on Sr/Ca ratios in aragonitic marine bivalve shells. *Geochemistry, Geophysics, Geosystems* 6.
- Gillikin, D.P., Lorrain, A., Paulet, Y.-M., André, L., Dehairs, F., 2008. Synchronous barium peaks in high-resolution profiles of calcite and aragonite marine bivalve shells. *Geo-Marine Letters* 28, 351–358.
- Goodwin, D.H., Flessa, K.W., Schöne, B.R., Dettman, D.L., 2001. Cross-calibration of daily growth increments, stable isotope variation, and temperature in the Gulf of California bivalve mollusk *Chione cortezi*: implications for paleoenvironmental analysis. *Palaaios* 16, 387–398.
- Goreau, T.J., 1977. Coral skeletal chemistry: physiological and environmental regulation of stable isotopes and trace metals in *Montastrea annularis*. *Proceedings of the Royal Society of London B: Biological Sciences* 196, 291–315.
- Hart, D.R., Chute, A.S., 2009. Verification of Atlantic sea scallop (*Placopecten magellanicus*) shell growth rings by tracking cohorts in fishery closed areas. *Canadian Journal of Fisheries and Aquatic Sciences* 66, 751–758.
- Hesselbo, S.P., Gröcke, D.R., Jenkyns, H.C., Bjerrum, C.J., Farrimond, P., Bell, H.S.M., Green, O.R., 2000. Massive dissociation of gas hydrate during a Jurassic oceanic anoxic event. *Nature* 406, 392–395.
- Huber, B.T., Hodell, D.A., Hamilton, C.P., 1995. Middle–Late Cretaceous climate of the southern high latitudes: stable isotopic evidence for minimal equator-to-pole thermal gradients. *Geological Society of America Bulletin* 107, 1164–1191.
- Immenhauser, A., Nägler, T.F., Steuber, T., Hippler, D., 2005. A critical assessment of mollusk  $18\text{O}/16\text{O}$ , Mg/Ca, and  $44\text{Ca}/40\text{Ca}$  ratios as proxies for Cretaceous seawater temperature seasonality. *Palaeogeography, Palaeoclimatology, Palaeoecology* 215, 221–237. doi:10.1016/j.palaeo.2004.09.005
- Kerswill, C.J., 1949. Effects of water circulation on the growth of quahaugs and oysters. *Journal of the Fisheries Board of Canada* 7, 545–551.
- Kirby, M.X., Soniat, T.M., Spero, H.J., 1998. Stable isotope sclerochronology of Pleistocene and Recent oyster shells (*Crassostrea virginica*). *Palaaios* 13, 560–569.
- Klein, R.T., Lohmann, K.C., Thayer, C.W., 1996a. Bivalve skeletons record sea-surface temperature and  $\delta^{18}\text{O}$  via Mg/Ca and  $18\text{O}/16\text{O}$  ratios. *Geology* 24, 415–418.
- Klein, R.T., Lohmann, K.C., Thayer, C.W., 1996b. Sr/Ca and  $13\text{C}/12\text{C}$  ratios in skeletal calcite of *Mytilus trossulus*: Covariation with metabolic rate, salinity, and carbon isotopic composition of seawater. *Geochimica et Cosmochimica Acta* 60, 4207–4221.
- Klein, R.T., Lohmann, K.C., Kennedy, G.L., 1997. Elemental and isotopic proxies of paleotemperature and paleosalinity: Climate reconstruction of the marginal northeast Pacific ca. 80 ka. *Geology* 25, 363–366.
- Kutzbach, J.E., Guetter, P.J., Ruddiman, W.F., Prell, W.L., 1989. Sensitivity of climate to late Cenozoic uplift in Southern Asia and the American West: Numerical experiments. *Journal of Geophysical Research: Atmospheres* 94, 18393–18407.
- Lazareth, C.E., Vander Putten, E., André, L., Dehairs, F., 2003. High-resolution trace element profiles in shells of the mangrove bivalve *Isochnomon ephippium*: a record of environmental spatio-temporal variations? *Estuarine, Coastal and Shelf Science* 57, 1103–1114.
- Lee, S.W., Kim, Y.M., Kim, R.H., Choi, C.S., 2008. Nano-structured biogenic calcite: A thermal and chemical approach to folia in oyster shell. *Micron* 39, 380–386. doi:10.1016/j.micron.2007.03.006
- Leng, M.J., Marshall, J.D., 2004. Palaeoclimate interpretation of stable isotope data from lake sediment archives. *Quaternary Science Reviews* 23, 811–831.
- Lorrain, A., Gillikin, D.P., Paulet, Y.-M., Chauvaud, L., Le Mercier, A., Navez, J., André, L., 2005. Strong kinetic effects on Sr/Ca ratios in the calcitic bivalve *Pecten maximus*. *Geology* 33, 965–968.
- MacDonald, B.A., Thompson, R.J., 1985. Influence of temperature and food availability on the ecological energetics of the giant scallop *Placopecten magellanicus*. I. Growth rates of shell and somatic tissue. *Marine ecology progress series. Oldendorf* 25, 279–294.
- McDermott, F., 2004. Palaeo-climate reconstruction from stable isotope variations in speleothems: a review. *Quaternary Science Reviews* 23, 901–918.
- Palmer, R.E., Carriker, M.R., 1979. Effects of cultural conditions on morphology of the shell of the oyster *Crassostrea virginica*, in: *Proceedings of the National Shellfisheries Association*. pp. 58–72.
- R Core Team, 2013. R: A language and environment for statistical computing. R Foundation for Statistical Computing. Vienna, Austria.
- Reinfelder, J.R., Wang, W.-X., Luoma, S.N., Fisher, N.S., 1997. Assimilation efficiencies and turnover rates of trace elements in marine bivalves: a comparison of oysters, clams and mussels. *Marine Biology* 129, 443–452.
- Richardson, C.A., Peharda, M., Kennedy, H., Kennedy, P., Onofri, V., 2004a. Age, growth rate and season of recruitment of *Pinna nobilis* (L) in the Croatian Adriatic determined from Mg: Ca and Sr: Ca shell profiles. *Journal of Experimental Marine Biology and Ecology* 299, 1–16.
- Richardson, C.A., Peharda, M., Kennedy, H., Kennedy, P., Onofri, V., 2004b. Age, growth rate and season of recruitment of *Pinna nobilis* (L) in the Croatian Adriatic determined from Mg: Ca and Sr: Ca shell profiles. *Journal of Experimental Marine Biology and Ecology* 299, 1–16.
- Schöne, B., Tanabe, K., Dettman, D.L., Sato, S., 2003. Environmental controls on shell growth rates and  $\delta^{18}\text{O}$  of the shallow-marine bivalve mollusk *Phacosoma japonicum* in Japan. *Marine Biology* 142, 473–485.
- Schöne, B.R., Zhang, Z., Radermacher, P., Thébaud, J., Jacob, D.E., Nunn, E.V., Maurer, A.-F., 2011. Sr/Ca and Mg/Ca ratios of ontogenetically old, long-lived bivalve shells (Arctica islandica) and their function as paleotemperature proxies. *Palaeogeography, Palaeoclimatology, Palaeoecology* 302, 52–64.
- Sluijs, A., Schouten, S., Pagani, M., Woltering, M., Brinkhuis, H., Damsté, J.S.S., Dickens, G.R., Huber, M., Reichert, G.-J., Stein, R., others, 2006. Subtropical Arctic Ocean temperatures during the Palaeocene/Eocene thermal maximum. *Nature* 441, 610–613.



- Stenzel, H.B., 1963. Aragonite and calcite as constituents of adult oyster shells. *Science* 142, 232–233.
- Steuber, T., 1999. Isotopic and chemical intra-shell variations in low-Mg calcite of rudist bivalves (Mollusca-Hippuritacea): disequilibrium fractionations and late Cretaceous seasonality. *International Journal of Earth Sciences* 88, 551–570.
- Surge, D., Lohmann, K.C., 2008. Evaluating Mg/Ca ratios as a temperature proxy in the estuarine oyster, *Crassostrea virginica*. *Journal of Geophysical Research: Biogeosciences* 113.
- Surge, D., Lohmann, K.C., Dettman, D.L., 2001. Controls on isotopic chemistry of the American oyster, *Crassostrea virginica*: implications for growth patterns. *Palaeogeography, Palaeoclimatology, Palaeoecology* 172, 283–296.
- Takesue, R.K., van Geen, A., 2004. Mg/Ca, Sr/Ca, and stable isotopes in modern and Holocene *Protothaca staminea* shells from a northern California coastal upwelling region. *Geochimica et Cosmochimica Acta* 68, 3845–3861.
- Titschack, J., Zuschin, M., Spötl, C., Baal, C., 2010. The giant oyster *Hyotissa hyotis* from the northern Red Sea as a decadal-scale archive for seasonal environmental fluctuations in coral reef habitats. *Coral Reefs* 29, 1061–1075.
- Ullmann, C.V., Böhm, F., Rickaby, R.E., Wiechert, U., Korte, C., 2013. The Giant Pacific Oyster (*Crassostrea gigas*) as a modern analog for fossil ostreoids: isotopic (Ca, O, C) and elemental (Mg/Ca, Sr/Ca, Mn/Ca) proxies. *Geochemistry, Geophysics, Geosystems* 14, 4109–4120.
- Ullmann, C.V., Frei, R., Korte, C., Lüter, C., 2017. Element/Ca, C and O isotope ratios in modern brachiopods: Species-specific signals of biomineralization. *Chemical Geology*.
- Ullmann, C.V., Wiechert, U., Korte, C., 2010. Oxygen isotope fluctuations in a modern North Sea oyster (*Crassostrea gigas*) compared with annual variations in seawater temperature: Implications for palaeoclimate studies. *Chemical Geology* 277, 160–166.
- Urban, H.-J., 2000. Culture potential of the pearl oyster (*Pinctada imbricata*) from the Caribbean.: I. Gametogenic activity, growth, mortality and production of a natural population. *Aquaculture* 189, 361–373.
- Vermeij, G.J., 2014. The oyster enigma variations: a hypothesis of microbial calcification. *Paleobiology* 40, 1–13.
- Vihtakari, M., Ambrose, W.G., Renaud, P.E., Locke, W.L., Carroll, M.L., Berge, J., Clarke, L.J., Cottier, F., Hop, H., 2017. A key to the past? Element ratios as environmental proxies in two Arctic bivalves. *Palaeogeography, Palaeoclimatology, Palaeoecology* 465, 316–332.
- Von Bertalanffy, L., 1957. Quantitative laws in metabolism and growth. *The quarterly review of biology* 32, 217–231.
- Wanamaker Jr, A.D., Kreutz, K.J., Wilson, T., Borns Jr, H.W., Introne, D.S., Feindel, S., 2008. Experimentally determined Mg/Ca and Sr/Ca ratios in juvenile bivalve calcite for *Mytilus edulis*: implications for paleotemperature reconstructions. *Geo-Marine Letters* 28, 359–368.
- Wang, W.-X., Fisher, N.S., 1996. Assimilation of trace elements and carbon by the mussel *Mytilus edulis*: effects of food composition. *Limnology and Oceanography* 41, 1.
- Wang, Y.-J., Cheng, H., Edwards, R.L., An, Z.S., Wu, J.Y., Shen, C.-C., Dorale, J.A., 2001. A high-resolution absolute-dated late Pleistocene monsoon record from Hulu Cave, China. *Science* 294, 2345–2348.
- Wiltshire, K.H., Malzahn, A., Greve, W., Wirtz, K., Janisch, S., Mangelsdorf, P., Manly, B.F., Boersma, M., 2008. Resilience of North Sea phytoplankton spring blooms dynamics: an analysis of long term data at Helgoland Roads. *Limnology and Oceanography* 53, 1294–1302.
- Zachos, J., Pagani, M., Sloan, L., Thomas, E., Billups, K., 2001. Trends, rhythms, and aberrations in global climate 65 Ma to present. *Science* 292, 686–693.
- Zachos, J.C., Schouten, S., Bohaty, S., Quattlebaum, T., Sluijs, A., Brinkhuis, H., Gibbs, S.J., Bralower, T.J., 2006. Extreme warming of mid-latitude coastal ocean during the Paleocene-Eocene Thermal Maximum: Inferences from TEX86 and isotope data. *Geology* 34, 737–740.

602

**FIGURE 1:** Schematic overview of the growth model and trace element module described in this paper. Yellow boxes indicate the modelling steps (Step 1–6) described in chapter 3, diamond-shaped elements represent data packages, ellipses represent model input parameters and boxes represent functions in the model. Elements are connected by arrows if they interact (i.e. if data packages and/or model parameters serve as input or output of model functions). Coloured data packages are the output of the model.

**FIGURE 2:** Example of a colour scan (top), a  $\mu$ XRF phase map (middle) and the digitized increments (bottom) of a *Crassostrea gigas* shell. Phase maps show the distribution of Araldite® 2020 resin (red), foliated calcite (green), chalky calcite (blue) and detrital material (yellow) in the shell cross sections. The *C. gigas* shell depicted in this figure corresponds to *C. gigas* shell #1 in Table 2.

**FIGURE 3:** Schematic illustration of morphology of a typical bivalve shell including an indication of all parameters used in the growth and trace element models. **Figure 2A** shows a cross section along the shell's major growth axis (XY plane), which is the plane along which the shells were sectioned. This cross section illustrates the parameters used to define shell increments and how interpolation between them is done (see section 3.1, model step 2). **Figure 2B** shows an overview of the shell and a definition of the axes (X, Y and Z) as



well as the parameters defined in the base ellipse of the shell (see section 3.1, model step 3). **Figure 2C** shows a perpendicular cross section through the width of the shell (YZ plane), which illustrates the parameters used in the determination of shell volume (see section 3.1, model step 4).

**FIGURE 4:** Overview of the results of quantitative XRF line scanning on a *C. gigas* shell (#1 in **Table 2**). On top of the figure is a colour scan of a cross section through the shell. The enlarged image on the left hand side shows the shell hinge including digitized growth increments (black lines with increment numbers), annual chronology interpreted from trace element records (yellow and transparent bands with years) and the location of the line scan (A to B, dark blue arrow). The lower right graph shows results of trace element records along the XRF line scan with growth increments (black lines) and annual chronology (yellow and transparent areas) indicated vertically on the graphs. From top to bottom, records of Ca (dark blue), Si (dark red), Zn (magenta), Mn (purple), Fe (orange), S (red), Mg (green) and Sr (blue) are plotted against line scan distance, increment number and time on three separate x-axes at the bottom of the graph.

**FIGURE 5:** Graphs showing modelled evolution of shell length (top left), average shell thickness (bottom left), shell volume and mass (top right) and shell growth rates (bottom right) with shell age. Solid blue lines in different shadings indicate records from individual *C. gigas* shells. Thin dashed blue curves indicate models fitted through the growth curves of *C. gigas* shells, while bold dashed black curves show models fitted through a composite of modelled data from all shells combined. Regression formulae and statistics are given in **Table 2**.

**FIGURE 6:** Plotted results of trace element modelling and line scanning in one of the *C. gigas* shells (#1 in **Table 2**). Shaded areas indicate the evolution of modelled accumulation rates (in mg/yr) of major and trace elements with shell age. Solid coloured lines indicate the change in modelled total shell concentrations with shell age. Coloured points connected by black lines indicate measured elemental concentrations in the hinge of the shells plotted against shell age (see also **Figure 4**)

**TABLE 1:** Table listing all functions used in the growth and trace element models (see chapter 3) and the variables used in these functions. Function names and names of data packages are also found in **Figure 1** and in the text.

**TABLE 2:** Overview of measured shell dimensions (top left), dimensions of XRF maps of all shells used in this study (top right), shell dimensions at the end of the model run (bottom left) and parameters of growth curves fitted through the modelled data (bottom right). Average density of shells was calculated from the averages of shell mass and volume.

**TABLE 3:** Table listing concentrations of all elements used in this study in both chalky and foliated calcite phases of *C. gigas* and *O. figari* shells. The “% of map”-column shows the amount of pixels the mineral phases take up relative to the total cross section area (not including resin mapped in the XRF mapping, see **Figure 3**).





649 **Supplementary data 1:** SVG and CSV files of cross sections through all the *Crassostrea gigas* shells used in this  
650 study

651 **Supplementary data 2:** Complete R-script used to model growth and trace element uptake as described in this  
652 study

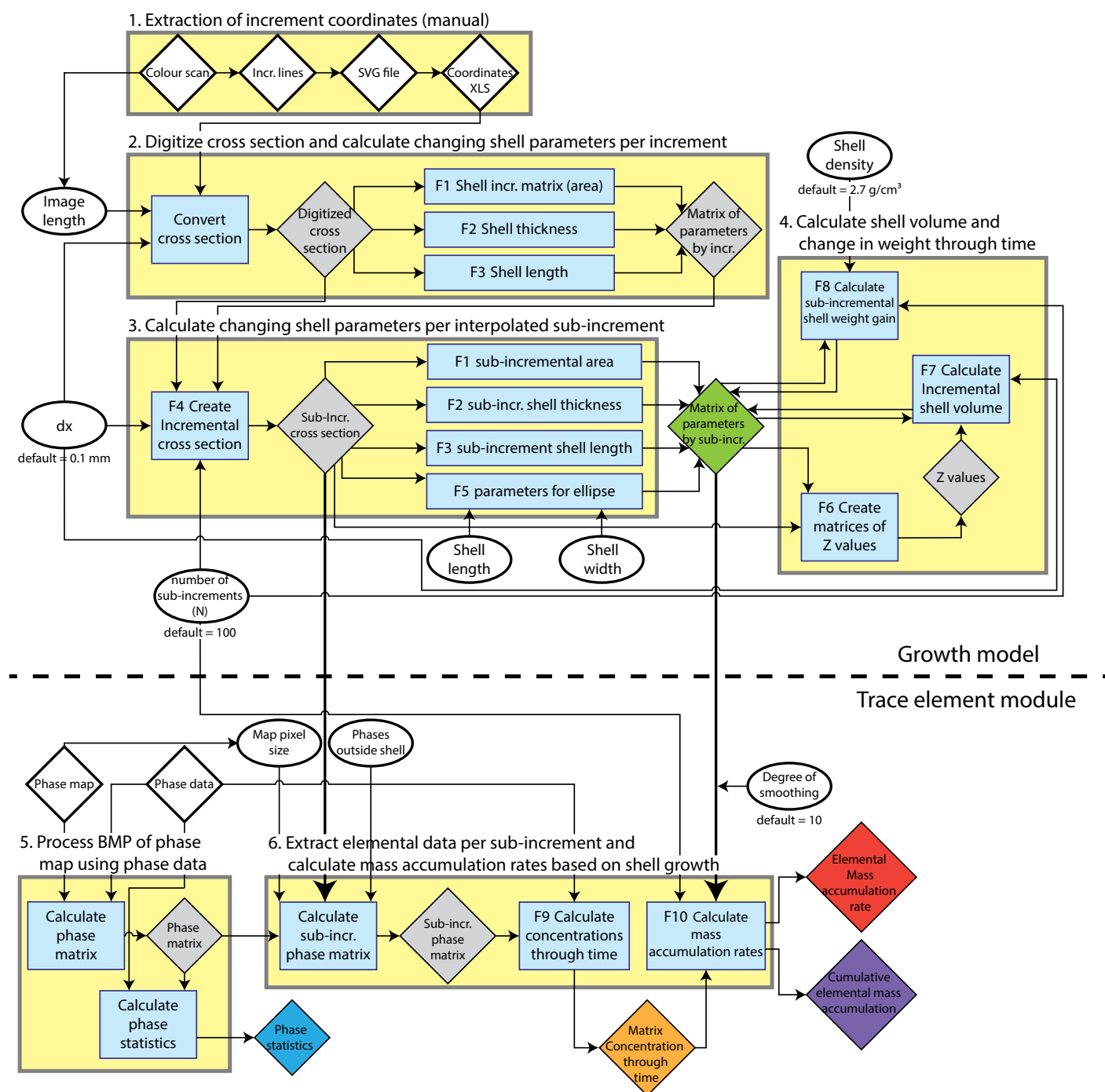
653 **Supplementary data 3:** Data of phase analysis of all trace element XRF maps including RGB colour values and  
654 trace element concentrations of all phases.

655 **Supplementary data 4:** Repository containing all data matrices generated by the model ran on all shells  
656 featuring in this study.

657 **Supplementary data 5:** BMP images of phase maps of all shell cross sections used as input of the trace  
658 element model in this study.



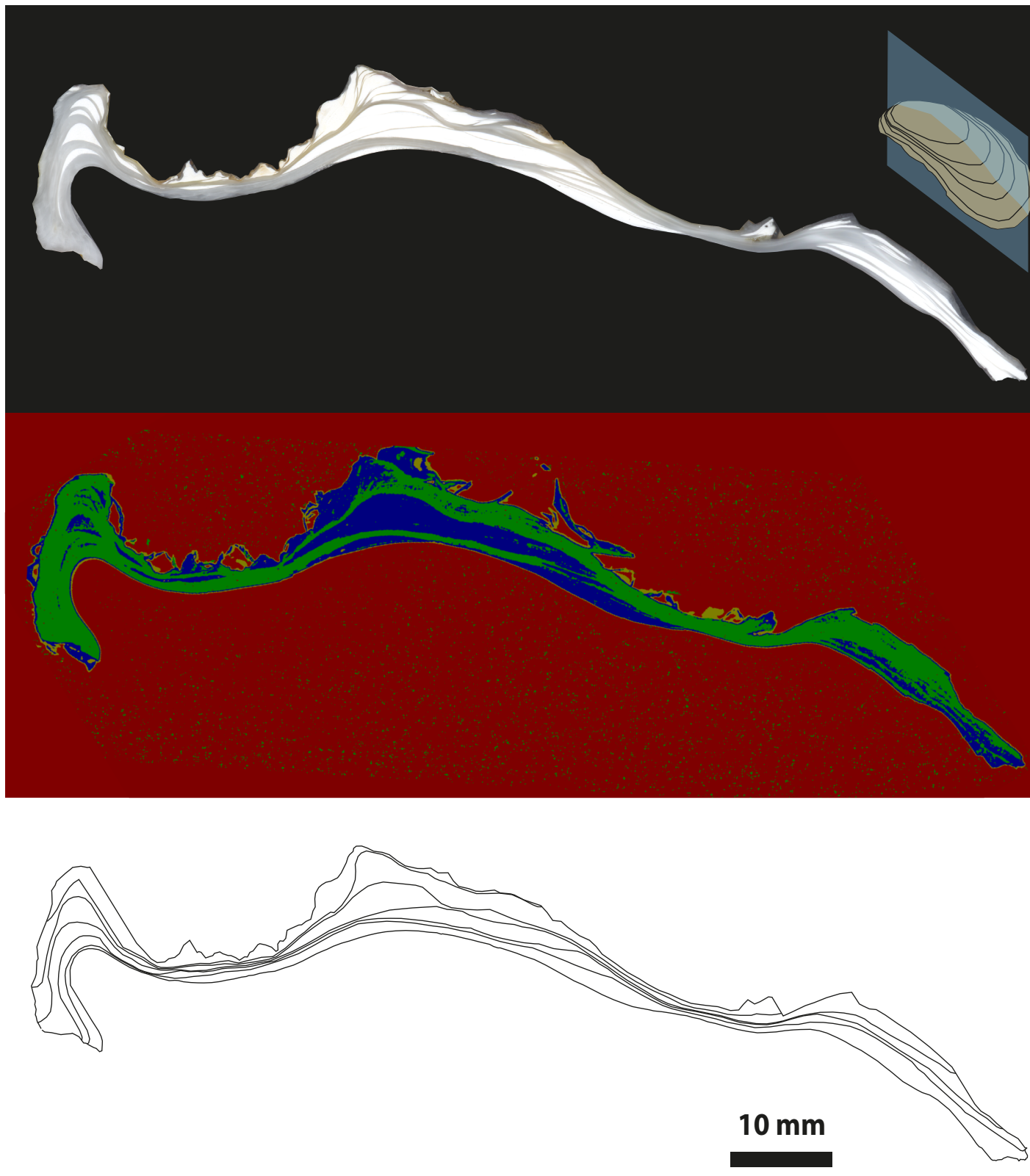
# Figure 1







## Figure 2



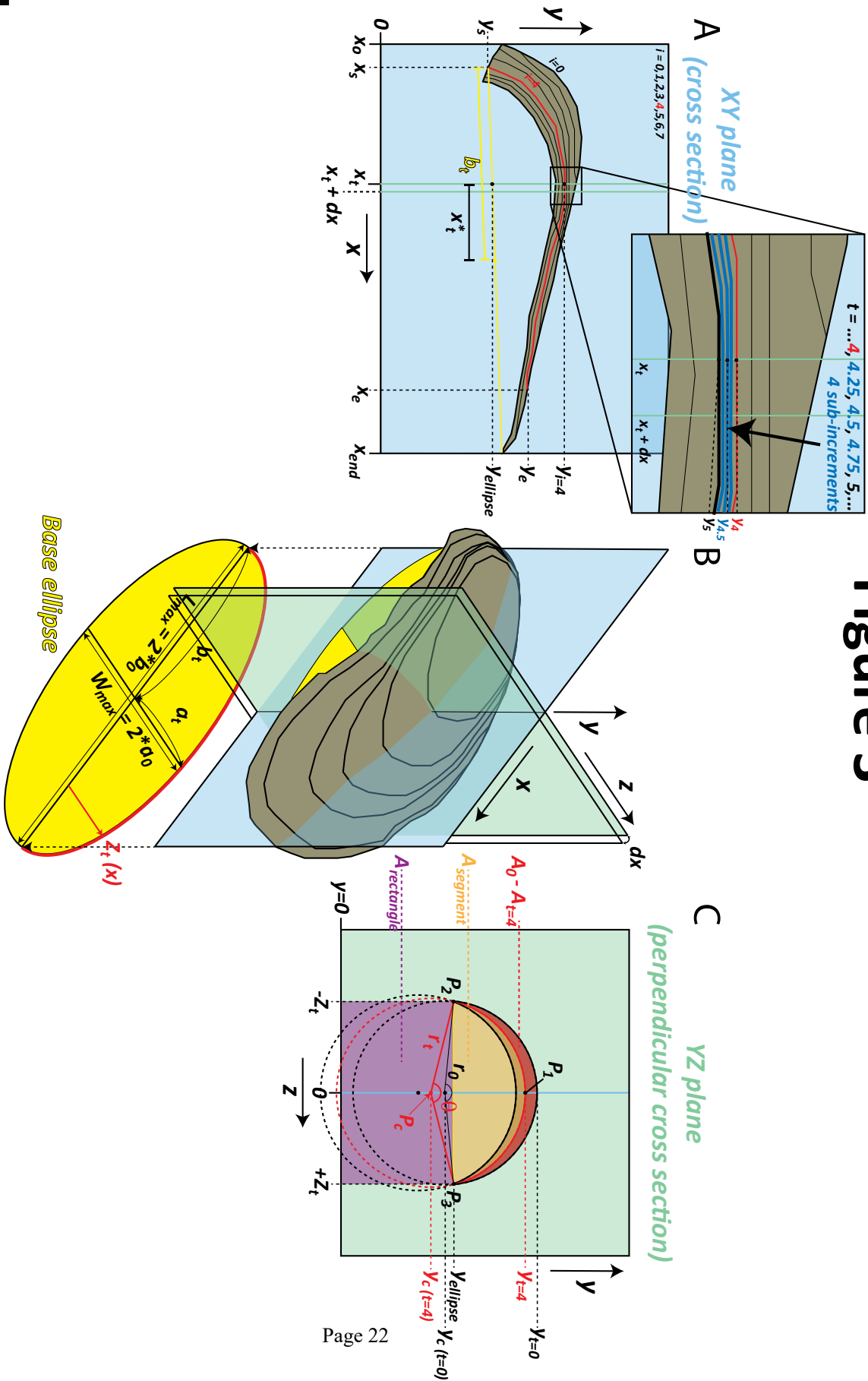
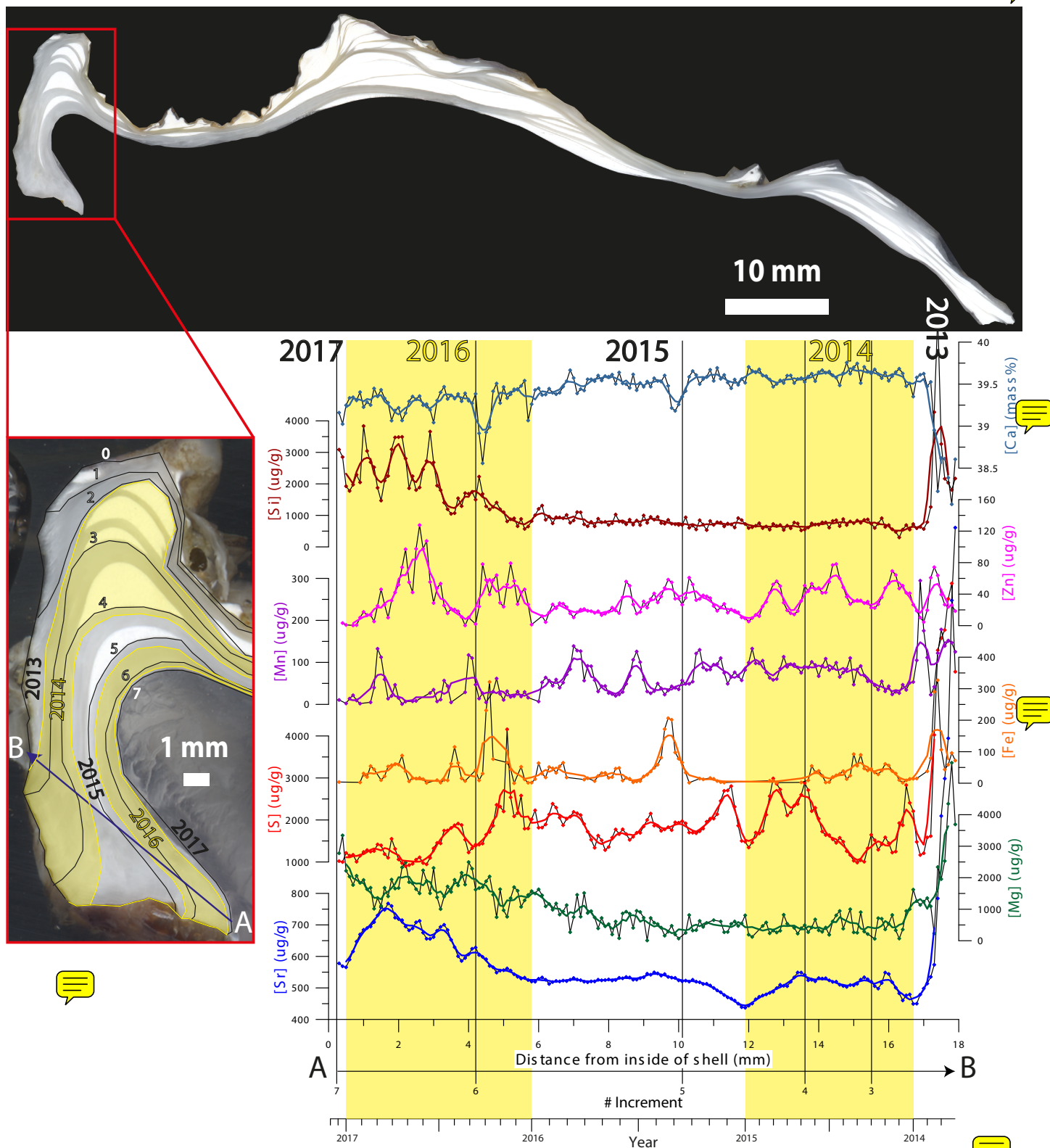


Figure 3

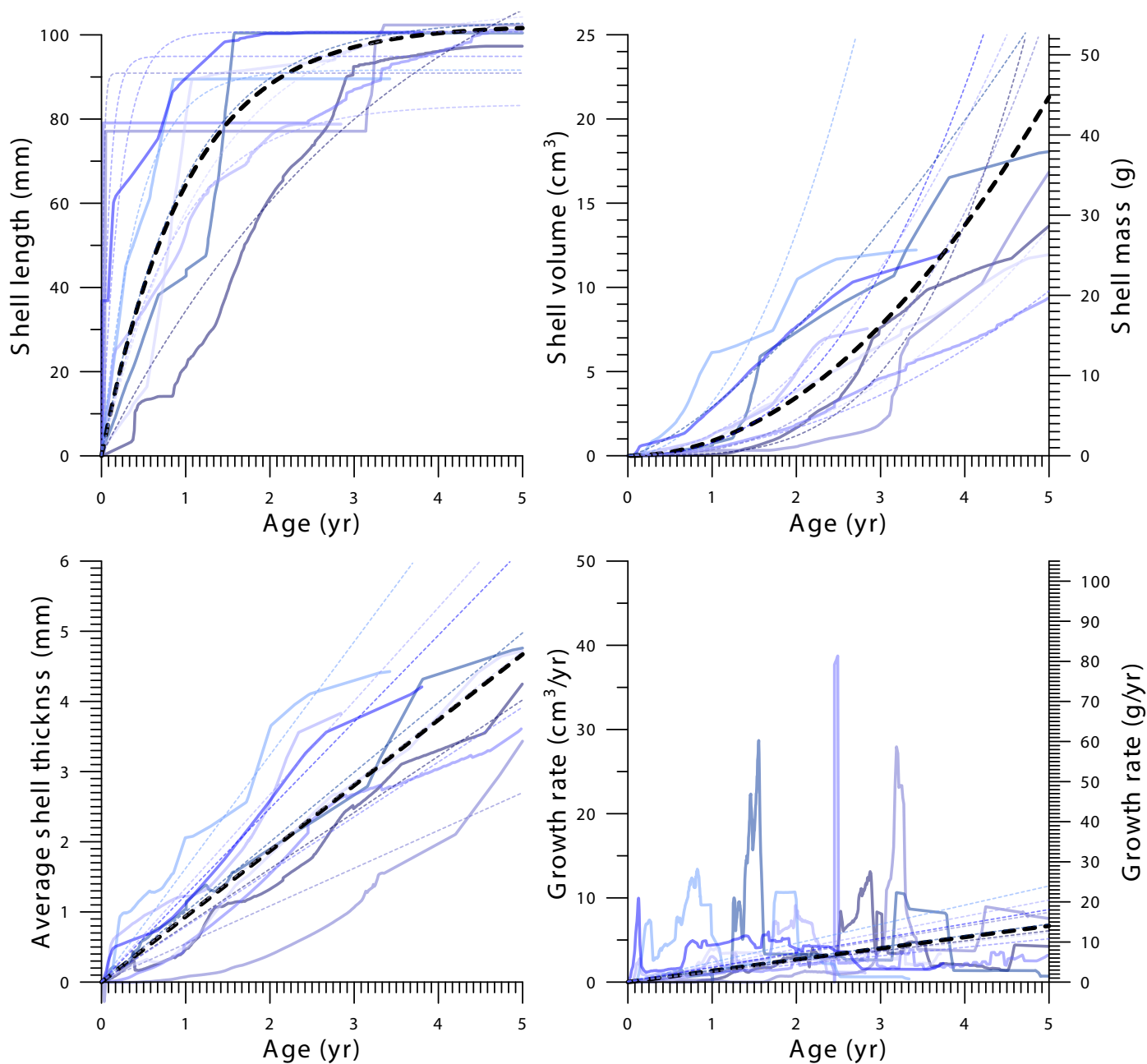


# Figure 4



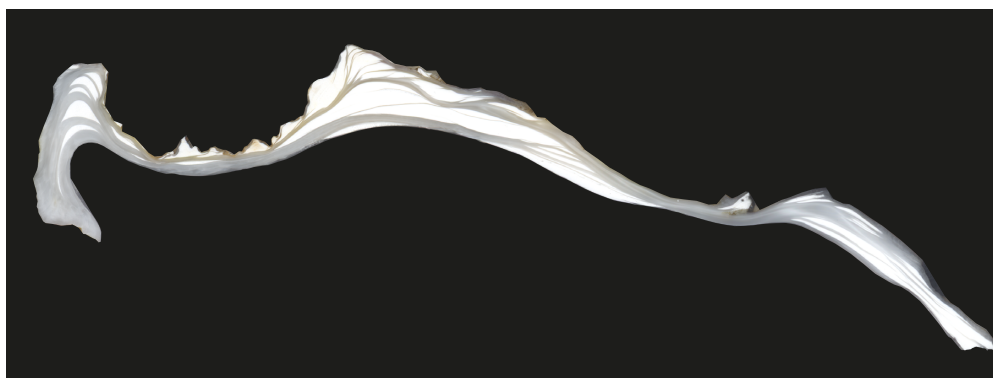
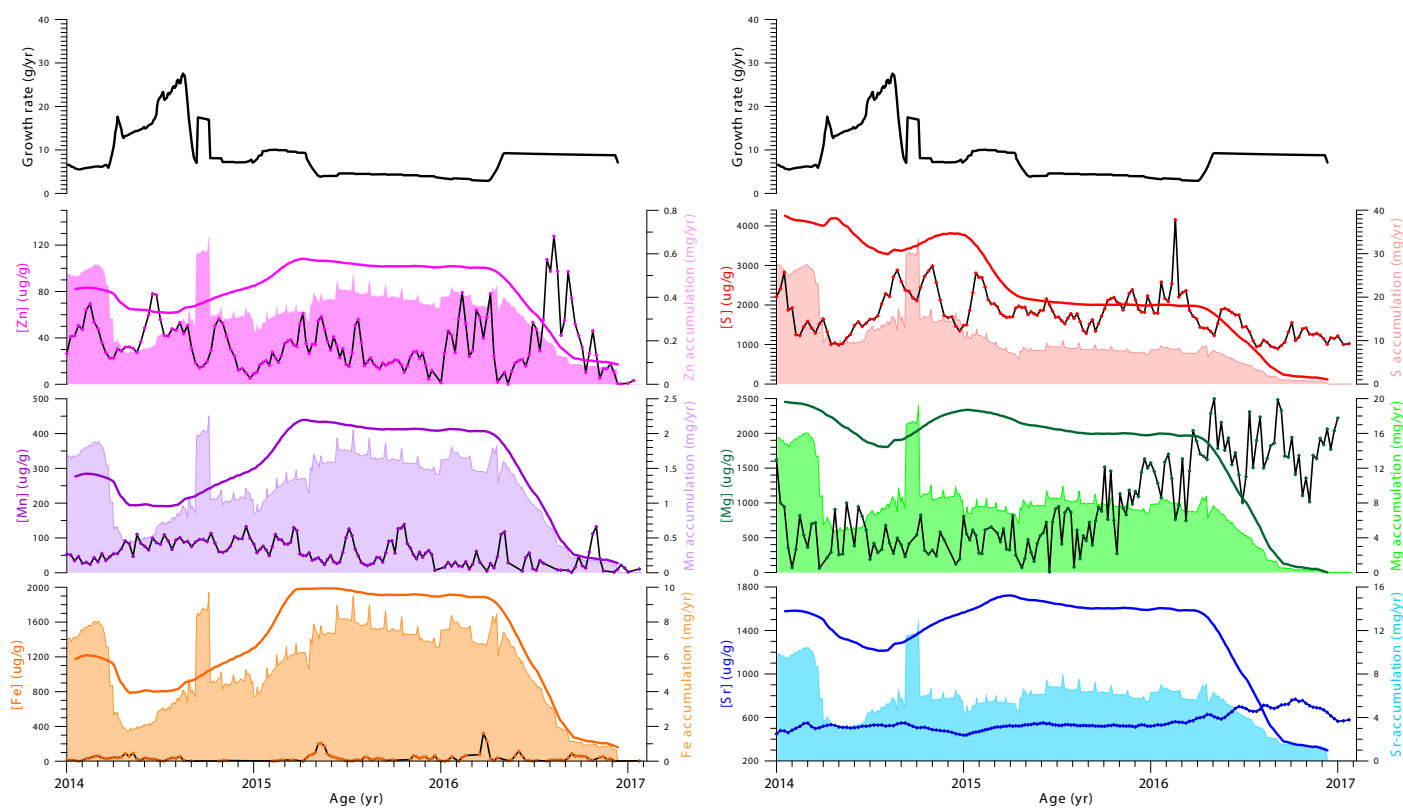


# Figure 5





# Figure 6





# Table 1

Model	Model step	Function name	Equation	variables IN	source of variables	variables OUT	variables stored in:
Growth model	Step 2	F1	$O_i = \int_{x_0}^{x_{end}} Y_{i-1}(x) - Y_i(x) dx$	$x_0$ $x_{end}$ $Y_i(x)$ $Y_{i-1}(x)$	Digitized cross section Digitized cross section Digitized cross section Digitized cross section	$A_i$	Matrix of parameters by increment
		F2	$T_i = \frac{\sum_{x_0}^{x_{end}} Y_i(x) - Y_i(x)}{x_{end} - x_0}$	$x_0$ $x_{end}$ $Y_i(x)$ $Y_{i-1}(x)$	Digitized cross section Digitized cross section Digitized cross section Digitized cross section	$T_i$	Matrix of parameters by increment
		F3	$L_i = \sqrt{(x_e - x_s)^2 + (Y_e - Y_s)^2}$	$x_e$ $x_s$ $Y_e$ $Y_s$	Digitized cross section Digitized cross section Digitized cross section Digitized cross section	$L_i$	Matrix of parameters by increment
	Step 3	F4	$F4.1 \quad \left[ \left[ Y_i(x) - Y_{i-1}(x) - \frac{n}{N} * (Y_{i-1}(x) - Y_i(x)) \right]_{n=0}^{N-1} \right]_{x=x_0}^{x=x_{end}}$	$Y_i(x)$ $Y_{i-1}(x)$	Digitized cross section Digitized cross section	$Y_i(x)$	Sub-incremental cross section
			$F4.2 \quad t = i - 1 + \frac{n}{N}$	$N$	Number of sub-increments		
		F5	$F5.1 \quad a_t = \frac{1}{2} * \frac{W_{max}}{L_{max}} * (x_e - x_s)$	$W_{max}$ $L_{max}$ $x_s$ $x_e$	Shell width Shell length Sub-incremental cross section Sub-incremental cross section	$a_t$	Matrix of parameters by sub-increment
			$F5.2 \quad b_t = \frac{1}{2} * (x_e - x_s)$	$x_s$ $x_e$	Sub-incremental cross section Sub-incremental cross section	$b_t$	Matrix of parameters by sub-increment
		F6	$F6.1 \quad \left( \frac{Z_t(x)}{a_t} \right)^2 + \left( \frac{x' - b_t}{b_t} \right)^2 = 1$	$a_t$	Matrix of parameters by sub-increment	$Z_t(x)$	Z-values
			$F6.2 \quad Z_t(x) = \left( \frac{a_t}{b_t} \right) * \sqrt{b_t^2 - x'^2}$	$b_t$	Matrix of parameters by sub-increment		
			$F6.3 \quad x' = x - x_e - b_t$	$x$ $x_s$ $x_0$	Sub-incremental cross section Sub-incremental cross section Sub-incremental cross section		
	Step 4	F7	$F7.1 \quad V_t = \int_{x_0}^{x_{end}} (A_0(x) - A_t(x)) dx$	$x_0$ $x_{end}$ $dx$	Sub-incremental cross section Sub-incremental cross section Sub-incremental cross section	$V_t$	Matrix of parameters by sub-increment
			$F7.2 \quad Y_{ellipse} = Y_e + \left( \frac{Y_e - Y_t}{x_e - x_s} \right) * (x - x_s)$	$Y_e$ $Y_s$	Sub-incremental cross section Sub-incremental cross section		
			$F7.3 \quad \Delta y^2 + \Delta x^2 = r^2$	$x_s$	Sub-incremental cross section		
			$F7.4 \quad (y_1 - y_e)^2 + (x_1 - x_e)^2 = (y_3 - y_e)^2 + (x_3 - x_e)^2$	$x_e$	Sub-incremental cross section		
			$F7.5 \quad (Y_i(x) - Y_e(x))^2 + (0 - 0)^2 = (Y_{ellipse}(x) - Y_e(x))^2 + (Z_t(x) - 0)^2$	$x_e$	Sub-incremental cross section		
			$F7.6 \quad Y_e(x) = \frac{(Y_t(x)^2 - Y_{ellipse}(x)^2 - Z_t(x)^2)}{2 * (Y_t(x) - Y_{ellipse}(x))}$	$x$	Matrix of parameters by sub-increment		
			$F7.7 \quad r_t(x) = \sqrt{Z_t(x)^2 + (Y_{ellipse}(x) - Y_e(x))^2}$	$Y_t(x)$	Matrix of parameters by sub-increment		
			$F7.8 \quad \theta = 2 * \sin^{-1} \left( \frac{Z_t(x)}{r_t(x)} \right)$	$Z_t(x)$	Z-values		
			$F7.9 \quad A_t = \begin{cases} A_{segment} + A_{rectangle} = \frac{1}{2} (r_t(x))^2 * (\theta - \sin \theta) + 2 * Z_t(x) * Y_{ellipse}(x), & Y_t(x) \geq Y_{ellipse}(x) \\ A_{rectangle} - A_{segment} = 2 * Z_t(x) * Y_{ellipse}(x) - \frac{1}{2} (r_t(x))^2 * (\theta - \sin \theta), & Y_t(x) < Y_{ellipse}(x) \end{cases}$				
		F8	$F8.1 \quad \Delta M_t = \rho * (V_{t-1} - V_t)$	$V_t$ $V_{t-1}$	Matrix of parameters by sub-increment Matrix of parameters by sub-increment	$\Delta M_t$	Matrix of parameters by sub-increment
			$F8.2 \quad M_t = \rho * V_t$	$\rho$	Shell density		
	Step 6	F9	$C_t^E = \sum_{p=p_1}^{p=p_n} \frac{S_p}{S_{tot}} * C_p^E \quad p \in [p_1, p_2, p_3, \dots, p_n]$	$S_p$ $S_{tot}$ $C_p^E$	Sub-increment phase matrix Sub-increment phase matrix Phase data	$C_t^E$	Matrix of concentration through time
		F10.1	$M_t^E = C_t^E * M_t$	$C_t^E$	Matrix of concentration through time		
		F10	$\left[ \frac{\partial M}{\partial t} \right]_t^E = C_t^E * N * \Delta M_t$	$M_t$	Matrix of parameters by sub-increment		
				$\Delta M_t$ $N$	Matrix of parameters by sub-increment Number of sub-increments		



# Table 2

Oyster dimensions							Oyster XRF map dimensions						
#	Shell length (mm)	Shell width (mm)	Maximum shell thickness (mm)	Shell mass (g)	Shell volume (cm3)	Density	map width (mm)	map length (mm)	#pixels in X-direction	#pixels in Y-direction	Total # of pixels	Spatial resolution (um)	
<i>Crassostrea gigas</i> #1	93.06	56.78	9.26	45.00	21.38	2.10	78.10	24.93	3124	997	3114628	25	
<i>Crassostrea gigas</i> #2	100.82	59.18	7.22	50.46	21.38	2.10	101.01	33.30	3367	1110	3737370	30	
<i>Crassostrea gigas</i> #3	101.94	43.07	8.84	40.25	19.13	2.10	102.00	39.00	3400	1300	4420000	30	
<i>Crassostrea gigas</i> #4	101.46	53.37	10.62	39.48	18.76	2.10	106.55	26.55	4262	1062	4582444	25	
<i>Crassostrea gigas</i> #5	86.47	60.46	7.71	54.93	26.10	2.10	88.30	29.50	3532	1180	4167760	25	
<i>Crassostrea gigas</i> #6	83.52	53.95	3.39	23.88	11.35	2.10	85.68	31.53	3427	1261	4321447	25	
<i>Crassostrea gigas</i> #7	100.86	51.60	7.05	42.95	20.41	2.10	103.00	20.10	4120	804	3312480	25	
<i>Crassostrea gigas</i> #8	101.50	57.77	4.54	39.38	18.71	2.10	102.66	34.80	3422	1160	3969520	30	
Average	96.20	54.52	7.33	42.04	19.98						3946181	26.88	
standard deviation	7.53	5.53	2.40	9.21	4.38						519354	2.59	

Model results						
#	Shell length (mm)	Shell width (mm)	Average shell thickness (mm)	Shell mass (g)	Shell volume (cm3)	Density
<i>Crassostrea gigas</i> #1	97.31	59.37	4.61	30.66	14.60	2.10
<i>Crassostrea gigas</i> #2	100.45	58.96	4.89	38.65	18.40	2.10
<i>Crassostrea gigas</i> #3	102.34	43.24	5.88	58.26	27.74	2.10
<i>Crassostrea gigas</i> #4	101.08	53.17	6.41	36.57	17.41	2.10
<i>Crassostrea gigas</i> #5	88.56	62.62	4.42	25.66	12.21	2.10
<i>Crassostrea gigas</i> #6	80.90	78.80	3.83	15.85	7.55	2.10
<i>Crassostrea gigas</i> #7	100.74	51.54	4.84	25.78	12.28	2.10
<i>Crassostrea gigas</i> #8	100.99	57.25	5.16	32.97	15.70	2.10
Average	96.36	54.63	5.01	33.05	15.74	
standard deviation	8.17	6.18	0.82	12.47	5.94	

Growth curve fits											
Fit equation	N	L0	k	R2	a	R2	a	e	R2	a	R2
Fit equation											
Fit parameter											
<i>Crassostrea gigas</i> #1	700	140.40	0.28	0.95	0.80	0.99	0.20	3.58	0.98	2.53	0.62
<i>Crassostrea gigas</i> #2	900	106.19	0.79	0.91	0.99	0.99	1.53	2.45	0.97	2.86	0.29
<i>Crassostrea gigas</i> #3	600	94.76	11.09	0.25	0.54	0.93	0.87	1.97	0.88	3.45	0.70
<i>Crassostrea gigas</i> #4	700	90.77	44.65	0.17	0.78	0.99	0.78	2.33	0.99	2.13	0.19
<i>Crassostrea gigas</i> #5	700	91.50	2.29	0.98	1.61	0.98	6.58	2.08	0.88	4.77	0.32
<i>Crassostrea gigas</i> #6	800	83.39	1.13	0.98	1.33	0.99	0.66	2.00	0.93	4.06	0.66
<i>Crassostrea gigas</i> #7	600	103.28	1.01	0.91	0.94	0.99	2.66	2.76	0.96	2.44	0.76
<i>Crassostrea gigas</i> #8	482	100.59	4.47	0.76	1.23	0.99	5.99	1.40	0.97	3.59	0.44
<i>Crassostrea gigas</i> composite	4655	102.34	0.99	0.60	0.94	0.91	1.83	1.98	0.76	2.81	0.21



Table 3

	Reproducibility error (1σ)	species	#	% of map	Mg	Si	P	S	Ca	Mn	Fe	Zn	Sr
					(μg/g) ± 422	(μg/g) ± 148	(μg/g) ± 4	(μg/g) ± 121	(μg/g) ± 395	(μg/g) ± 4	(μg/g) ± 6	(μg/g) ± 11	(μg/g) ± 3
Chalky calcite		<i>C. gigas</i>	1	62.53%	-	2384	2	116	383304	27	162	17	297
		<i>C. gigas</i>	2	68.87%	4001	4259	137	5341	374559	98	49	57	1877
		<i>C. gigas</i>	3	59.25%	4393	4934	151	5267	353857	33	7	31	1135
		<i>C. gigas</i>	4	49.55%	2368	4239	44	2891	378539	47	14	23	644
		<i>C. gigas</i>	5	83.91%	1643	3108	73	1627	390939	51	30	57	703
		<i>C. gigas</i>	6	79.12%	1802	2903	81	1923	388243	66	43	27	611
		<i>C. gigas</i>	7	40.40%	1847	4231	70	3097	384371	23	125	20	699
		<i>C. gigas</i>	8	71.97%	4267	5175	182	5921	357400	92	71	52	1358
Foliated calcite		<i>C. gigas</i>	1	37.47%	2744	4732	114	6399	351759	103	261	54	1433
		<i>C. gigas</i>	2	31.13%	1153	3644	78	3920	388370	70	13	75	1432
		<i>C. gigas</i>	3	40.75%	1801	3279	71	3069	381022	30	0	25	1382
		<i>C. gigas</i>	4	50.45%	1304	1455	484	2173	387183	61	39	26	694
		<i>C. gigas</i>	5	16.09%	2112	3837	48	1907	389176	65	138	62	1690
		<i>C. gigas</i>	6	20.88%	1903	6047	81	2040	386690	58	31	22	1510
		<i>C. gigas</i>	7	59.60%	1062	2161	31	1691	392954	38	1	17	728
		<i>C. gigas</i>	8	28.03%	2488	2522	72	4077	380805	78	37	52	1525

